Ocean – Atmosphere Interaction

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Ocean-Atmosphere Interaction

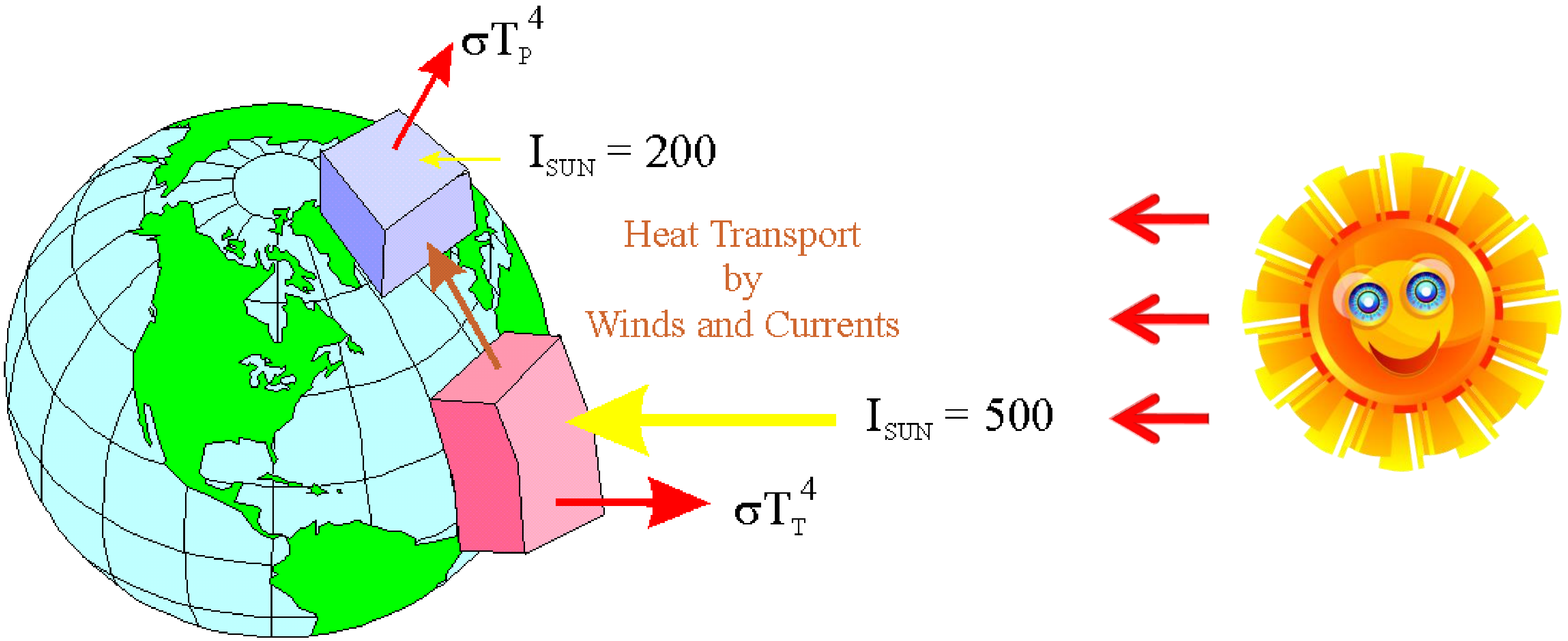
1. Tropical-Extratropical, Interhemispheric Climate Interaction : Atmospheric Bridge and Oceanic Tunnel
2. Dynamics of Decadal Climate Variability and Tropical Decadal Variability
3. Ocean-Atmosphere Interaction: A Global Scale, Coupled Climate Dynamics and Bjerknes Compensation
4. Timescale and Reversibility of Climate Change

Outline

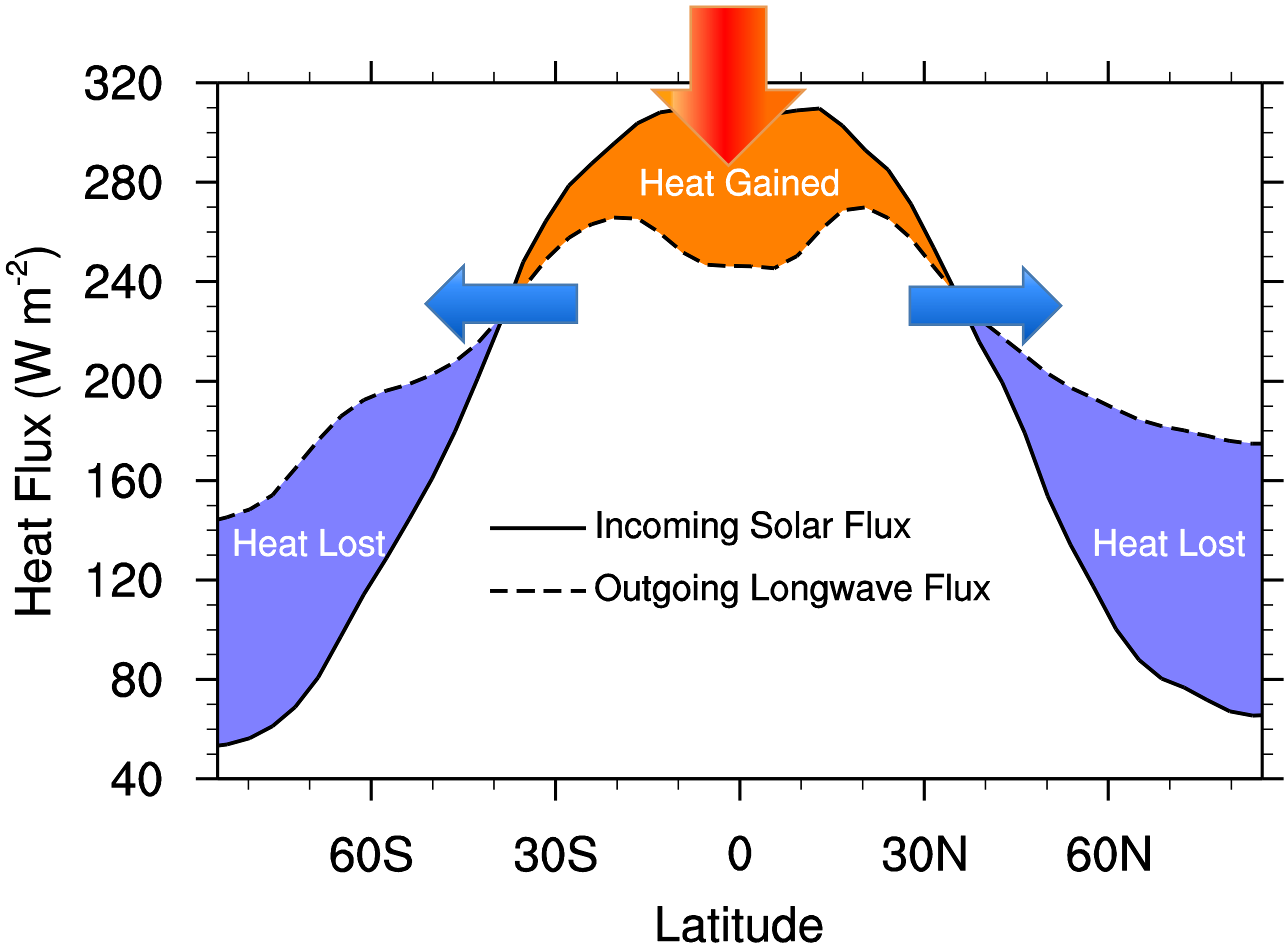
Bjerknes补偿：海气耦合系统本征模

- Fundamentals
- Questions
- Hypothesis and Theory
- CGCM results
- Aquaplanet
- Summary

Fundamentals



Heat Budget at the TOA



Fundamental Questions

Energy

$$Energy = c_p T + L_v q + gz + \frac{(u^2 + v^2 + w^2)}{2}$$

- **Sensible Heat / Latent heat**
- **Potential energy / Kinetic energy**

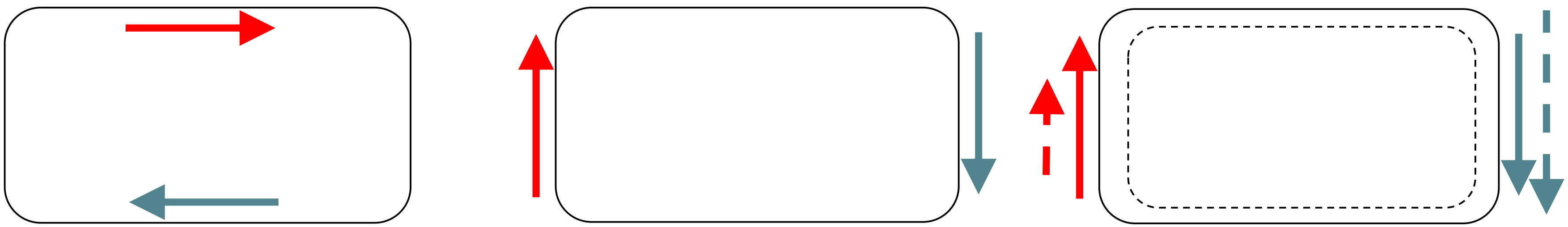
1. **Kinetic energy transport is small**
2. **In the ocean only sensible heat transport**

Fundamental Questions

Energy Transport Mechanisms

Energy multiply by meridional mass transport

$$F_E = \Psi \, C \, \Delta E$$



- **Hadley Cell**
- **Ferrell Cell**
- **AMOC**
- **Wind-driven STC**

- **Monsoon Circulation**
- **Oceanic Gyres**

Eddies:

- **Standing Waves**
- **Transient Storms**
- **Instability Waves**
- **Gulf Stream Rings…**

Ocean Heat Transport Calculation

Temperature equation

$$\frac{\partial T}{\partial t}+\frac{\partial (u+u^*+u')T}{\partial x}+\frac{\partial (v+v^*+v')T}{\partial y}+\frac{\partial (w+w^*+w')T}{\partial z}=R(T)+D_V(T)+\frac{\partial Q}{\partial z}.$$

Zonal and Vertical Integration

$$\iiint \frac{\partial T}{\partial t}dxdydz+\iint (v+v^*+v')T\,dxdz=\iiint R(T)\,dxdydz+\iint Q_0\,dxdy,$$

Storage　　OHT_{vT}　　OHT_{diff}　　OHT_{HF}

$OHT_{eul} + OHT_{bol} + OHT_{sbm}$

Ocean Heat Transport Calculation

No Submesoscale Eddies Equilibrium

$$OHT = OHT_{vT} = OHT_{eul} + \cancel{OHT_{bol}} + \cancel{OHT_{sbm}} = OHT_{HF} + \cancel{OHT_{diff}} - \cancel{Storage}$$

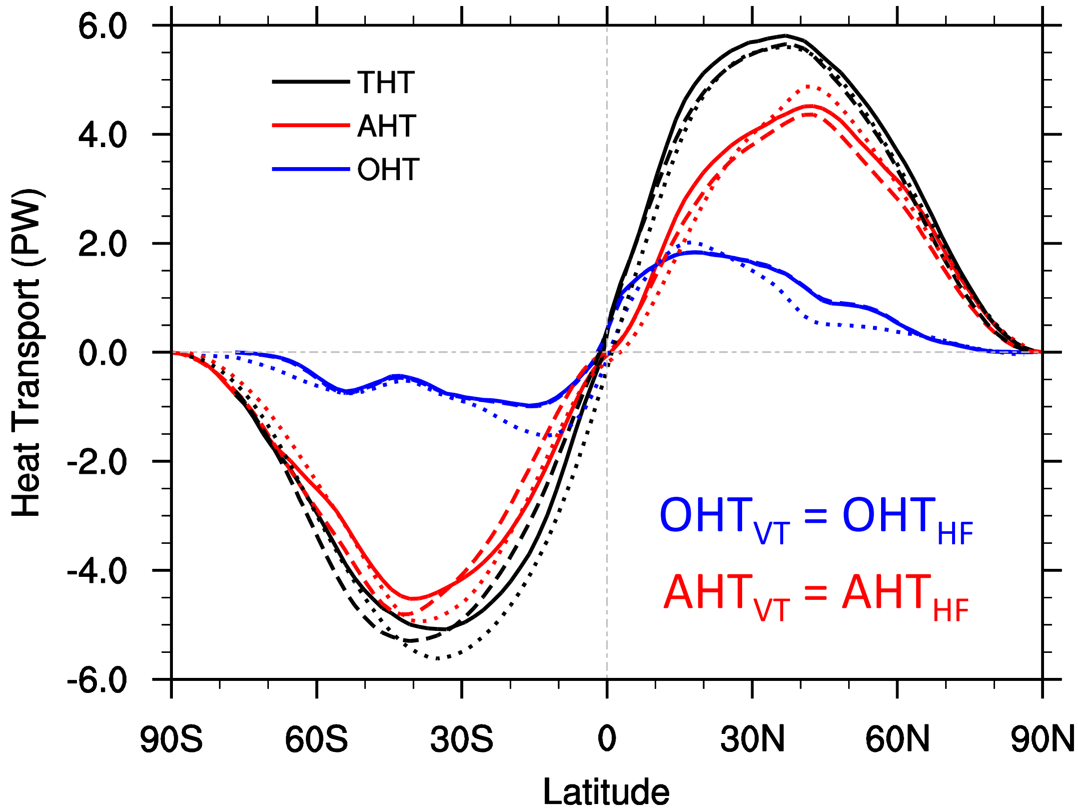
No Mesoscale Eddies No Mixing

$$OHT_{vT} \approx OHT_{HF} \approx OHT_{eul}$$

Calculation in CGCM

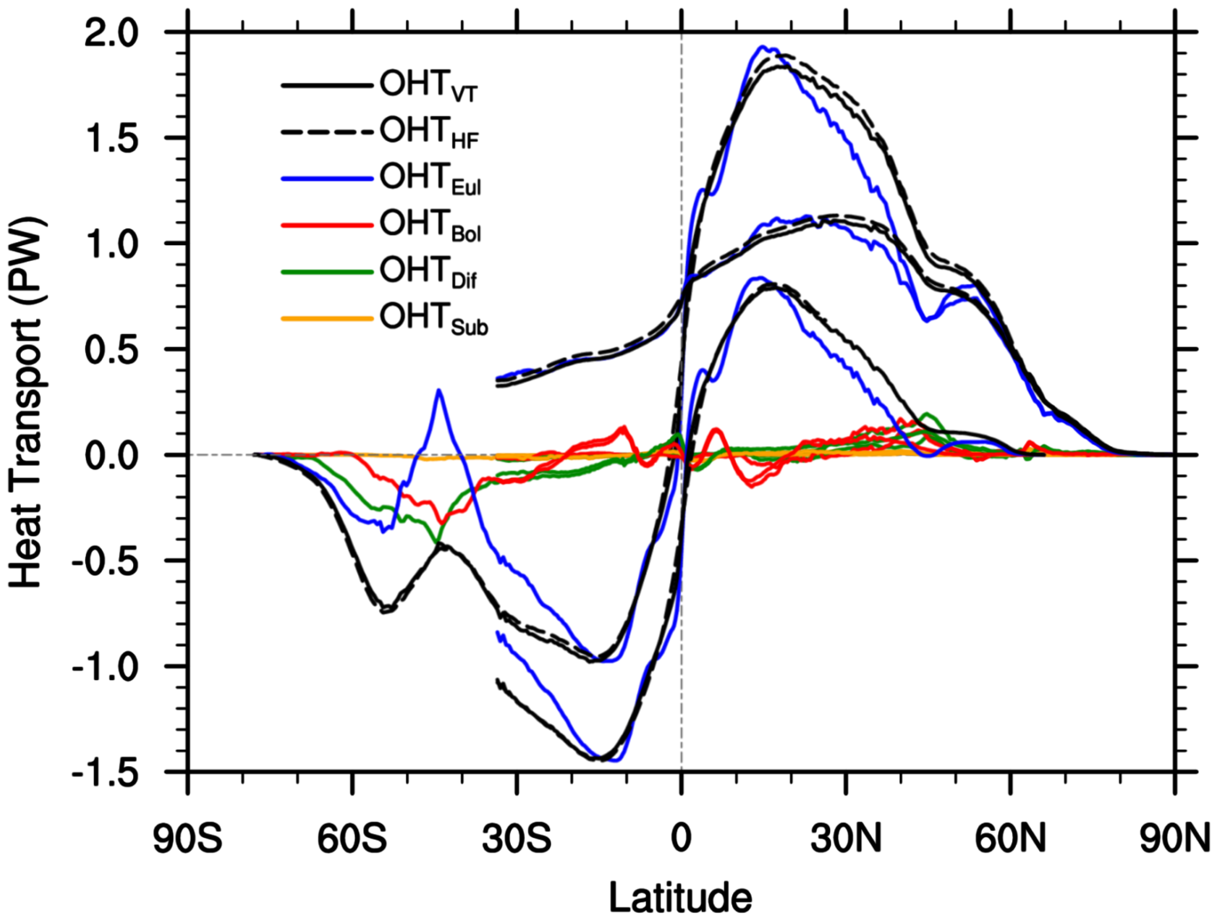
- **Real** heat transport by Atmosphere and Ocean
- OHT components:
  - Basin contribution / Wind-driven / Thermohaline
- AHT components:
  - Dry air VS. Water vapor / Mean VS. Eddies

Meridional Heat Transport



Dotted lines from Obs. Heat Flux (Trenberth and Caron, 2001)

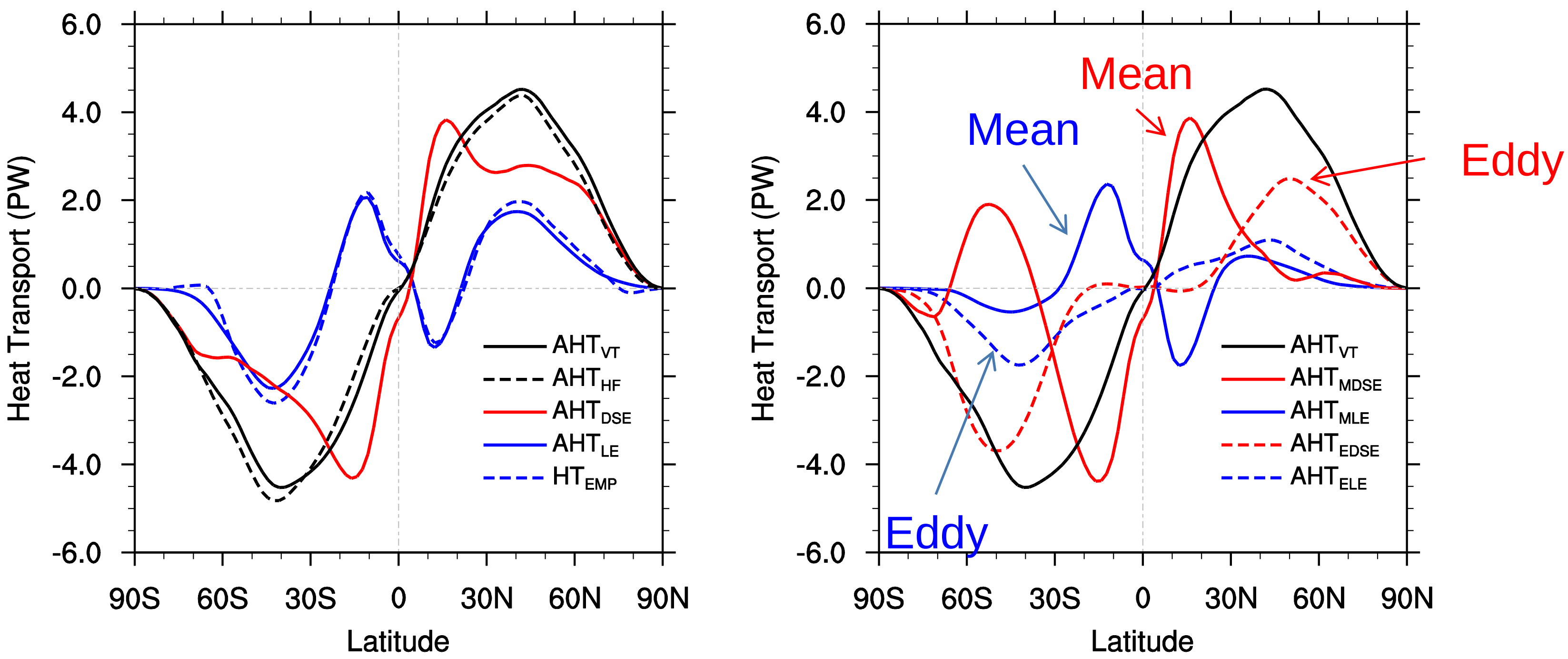
Decomposing OHT_{VT}



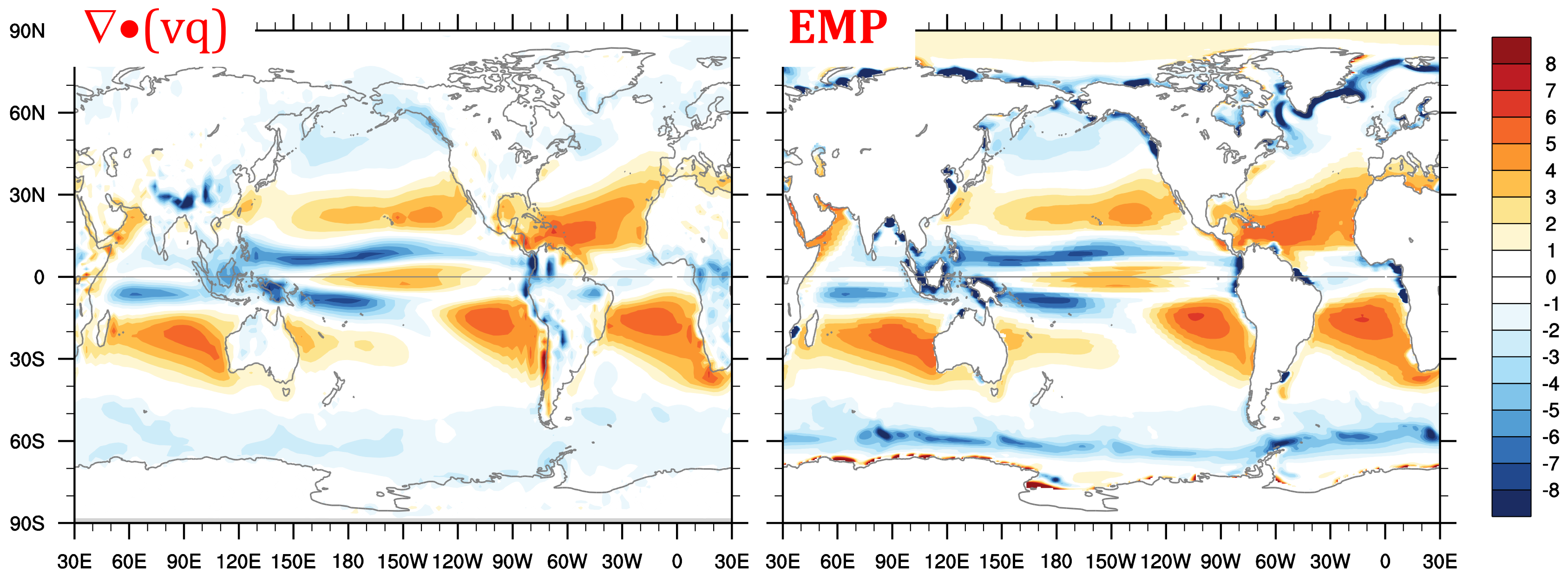
Consistency between OHT_{VT} and OHT_{HF}
Important role of Bolus and dissipation in ACC

Yang and Li (2015)

Decomposing AHT$_{VT}$



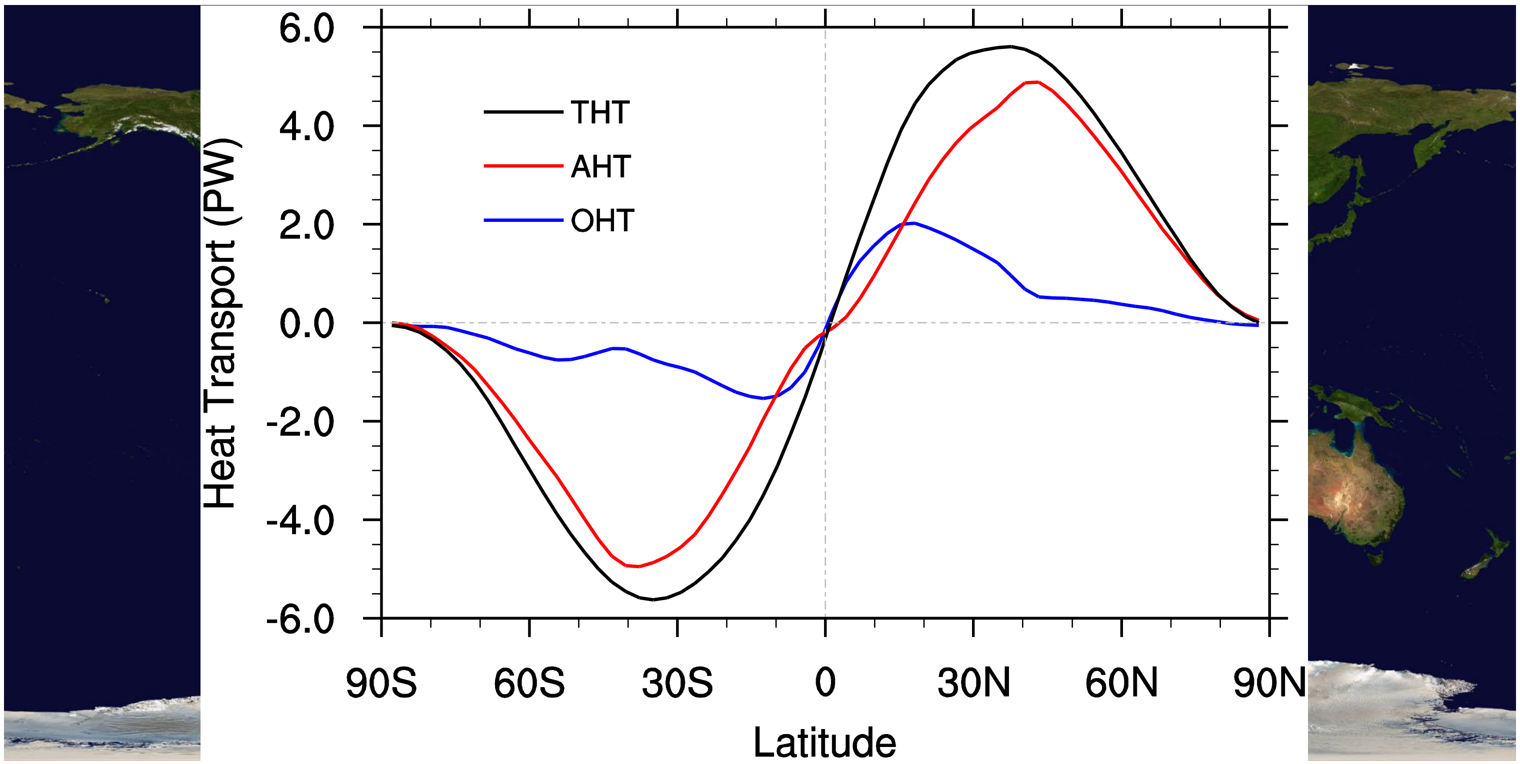
Divergence of Water Vapor Transport



- Vertically integrated water vapor divergence $\nabla \bullet (vq)$
- Comes from EMP over the global ocean (unit: 10^{-5} $kgm^{-2}s^{-1}$)

Fundamental Questions

1. Antisymmetric MHT?

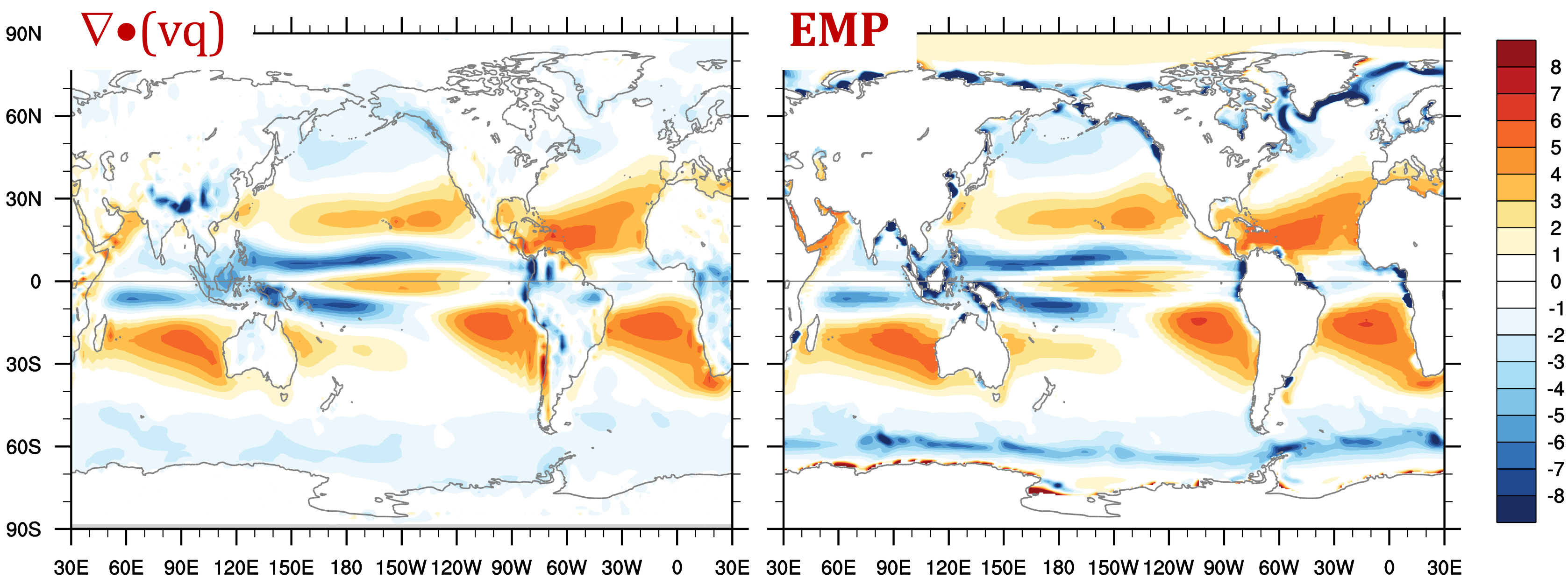


Aquaplanet ➔ Real Earth

Trenberth and Caron (2001)

Fundamental Questions

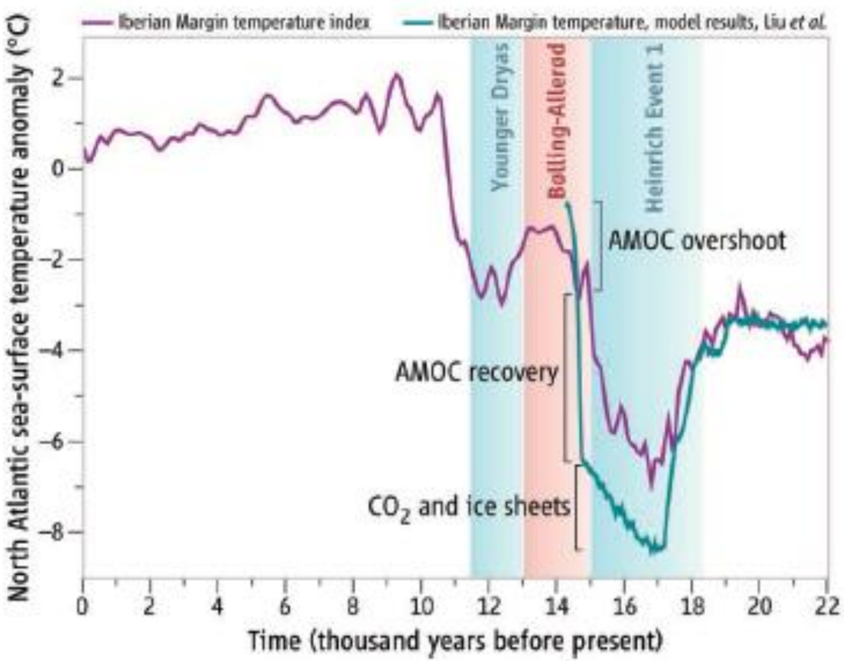
2. “Real” Oceanic Contribution?



Yang and Li (2015)

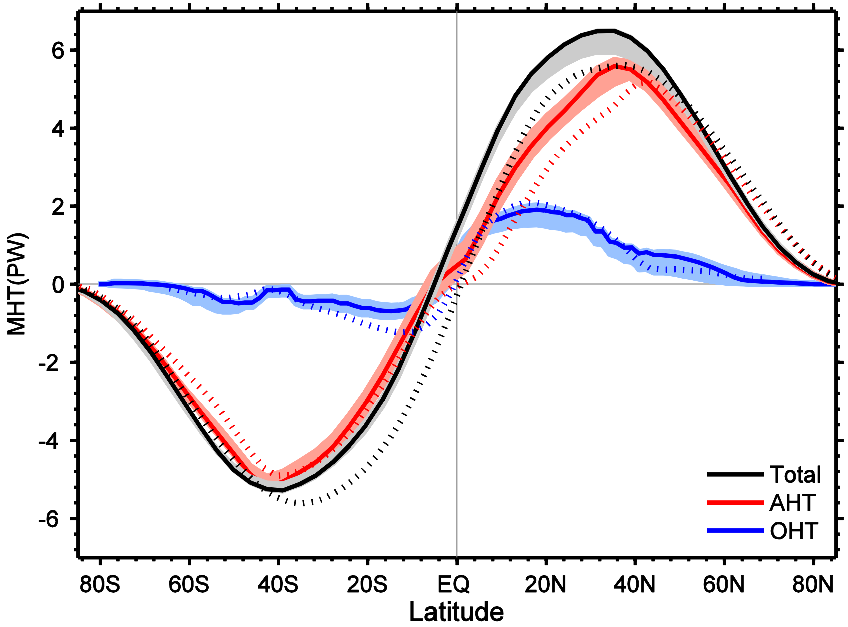
3. Relationship between OHT and AHT Changes?

Earth Climate Stability Mechanism



Climate Change during Past 22 kyr

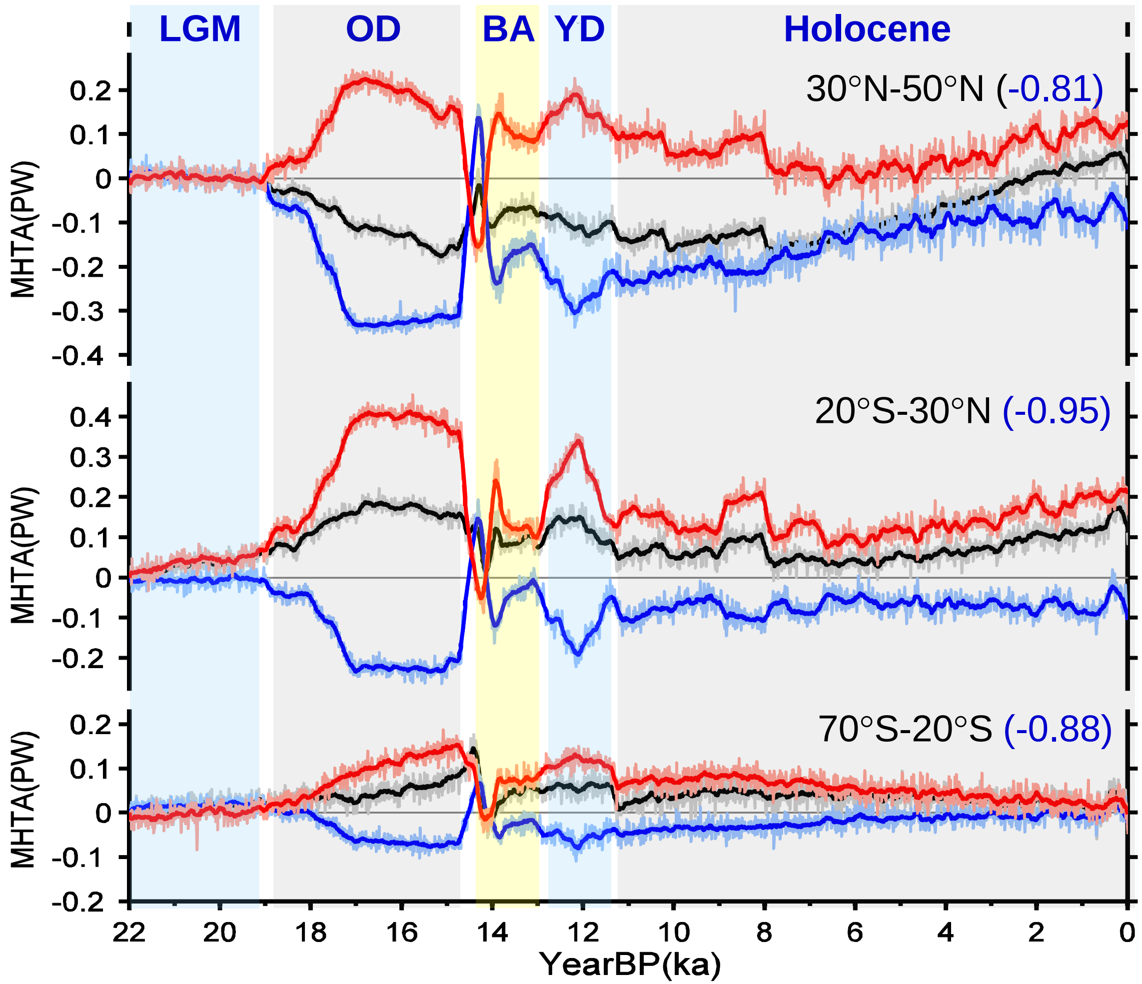
Timmermann (2009), Science



MHT from CCSM3 simulation TraCE-21K, From LGM to present

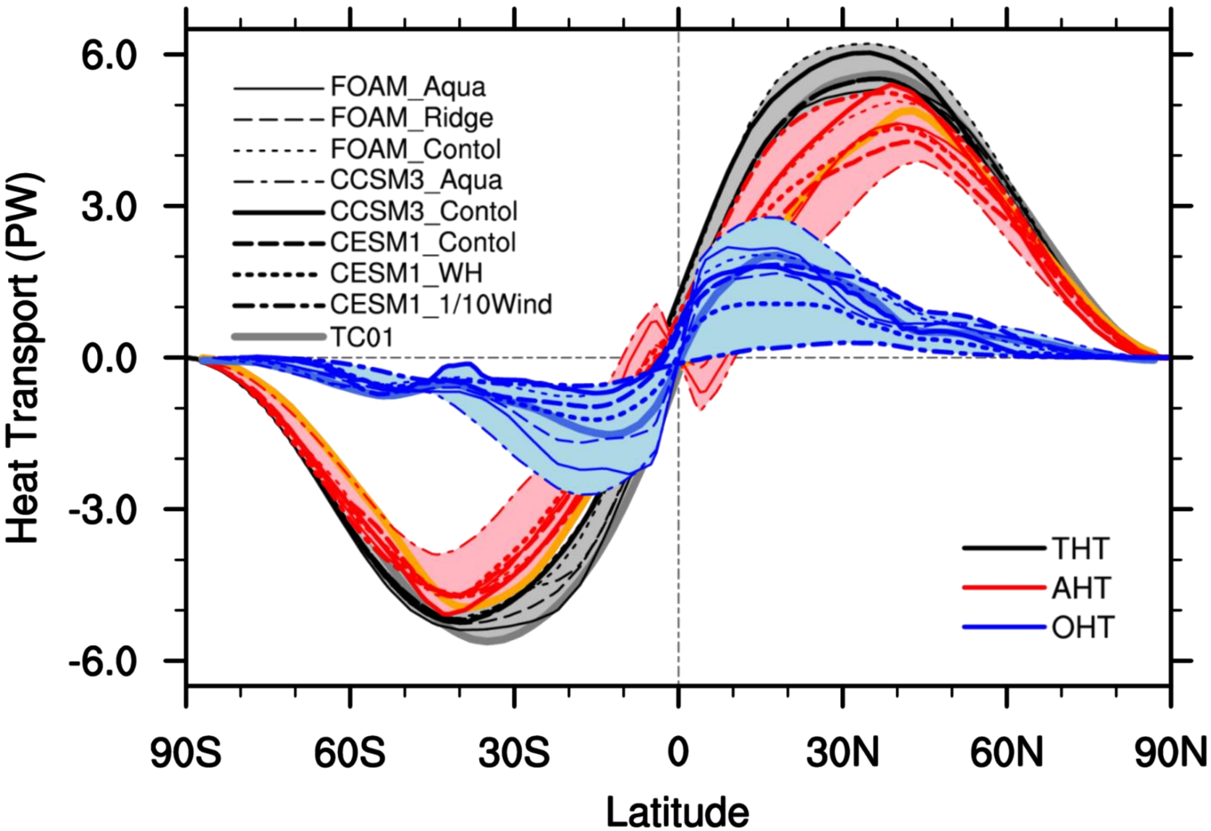
Liu et al. (2009); He (2011)

MHT Change Since LGM



Yang et al. (2015)

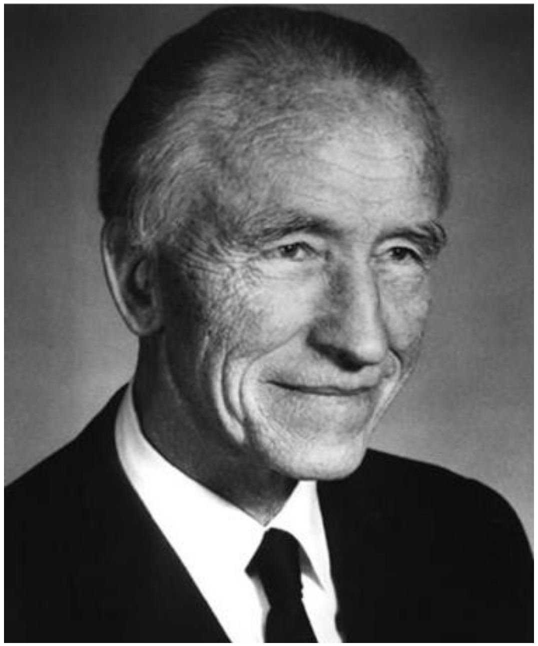
Compensation between AHT and OHT



Note: TC01 is from Trenberth and Caron (2001)

Hypothesis: Bjerknes Compensation

Jacob Aaal Bonnevie Bjerknes
1897-1975



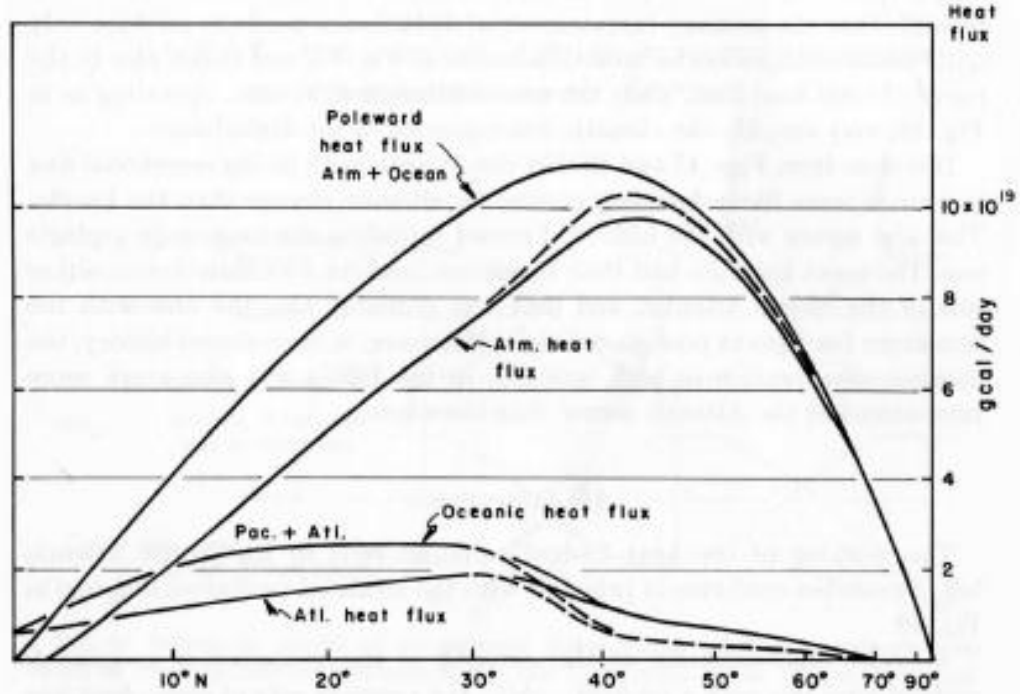
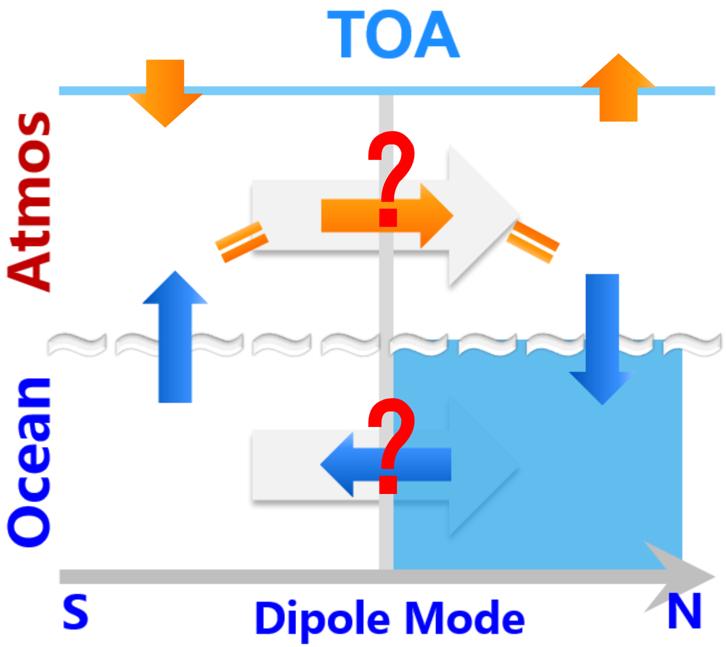


Fig. 48. Solid lines: flux data from Fig. 47 pertaining to present climatic conditions. Dashed lines refer to a sketchy model of the conditions around 1800 A.D. and show qualitative estimate of curtailed Atlantic and total oceanic heat flux as well as increased heat flux by low index atmospheric circulation. The anomalies of heat flux in oceans and atmosphere are assumed to cancel, leaving total heat flux and radiation budget unchanged. Actually, some change in the radiation budget is also likely to have taken place, but it could well have been quite small.

Bjerknes, 1964: Atlantic Air-Sea Interaction, Advances in Geophysics, Vol. 10, P77

Hypothesis: Bjerknes Compensation

Question: How Climate Feedback Determines BJC?



A + B = 0 → A = -B

but A + B + C = 0

C: climate feedback

Then A = -(B+C)

Energy Conserved

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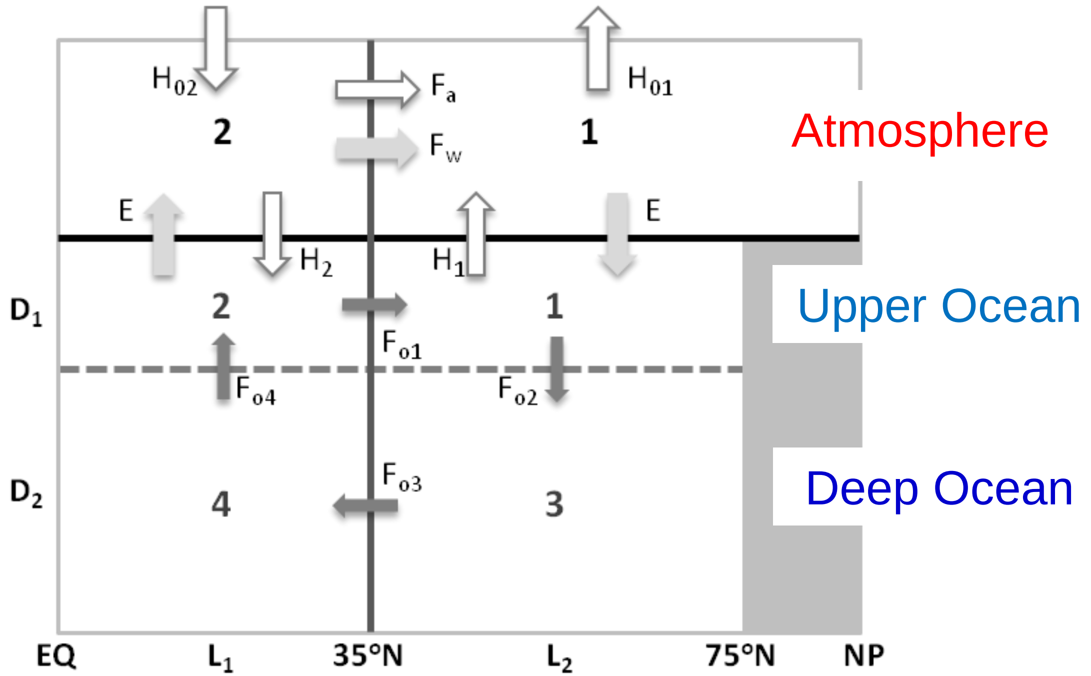
Theory for Equilibrium Change

1. **Coupled Multi-Box Model**
2. **1-D Energy Balance Model (EBM)**

Go to Final Equations

Coupled Multi-Box Model

Yang, Zhao and Liu, 2016: Understanding Bjerknes compensation in atmosphere and ocean heat transports using a coupled box model. J. Climate



Stommel (1961); Nakamura et al. (1994); Marotzke and Stone (1995); Tziperman et al. (1994); Tziperman and Ioannou (2002)

Go to Final Equations

Equations and Dynamics

T 1⋯ 2⋯ 3⋯ 4⋯	S 1⋯ 2⋯ 3⋯ 4⋯

Heat Flux at TOA: $H_{1,2} = A_{1,2} - B_{1,2} T_{1,2}$

Fresh Water Flux : $H_{fw} = \gamma (T_2 - T_1)$

Thermohaline : $q = \kappa [\alpha (T_2 - T_1) - \beta (S_2 - S_1)]$

AHT: $H_d \sim (T_2 - T_1)$ | **OHT:** $O_d \sim q^*(T_2 - T_1)$

Yang, H., Y. Zhao, and Z. Liu, 2016: Understanding Bjerknes compensation in atmosphere and ocean heat transports using a coupled box model. *J. Climate*, 29(6), 2145-2160, doi: 10.1175/JCLI-D-15-0281.1.

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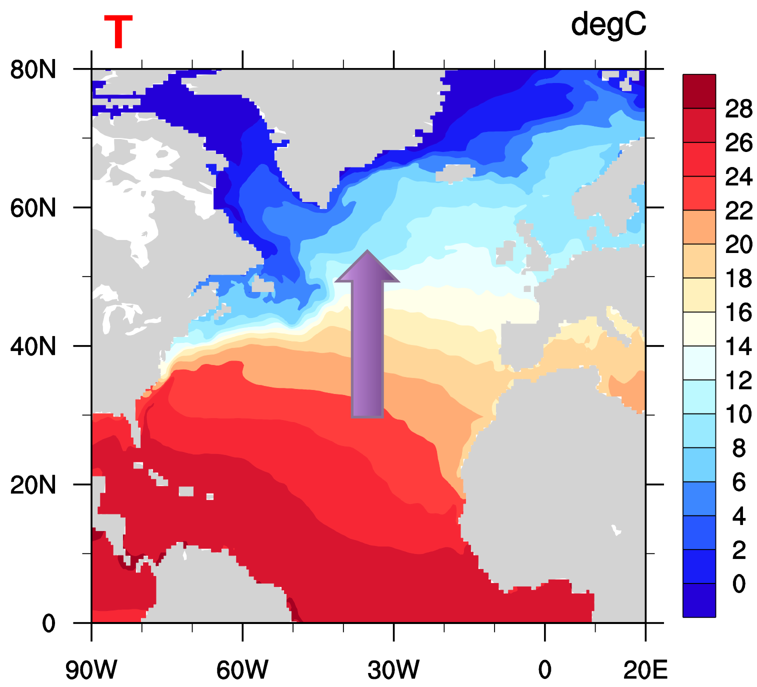
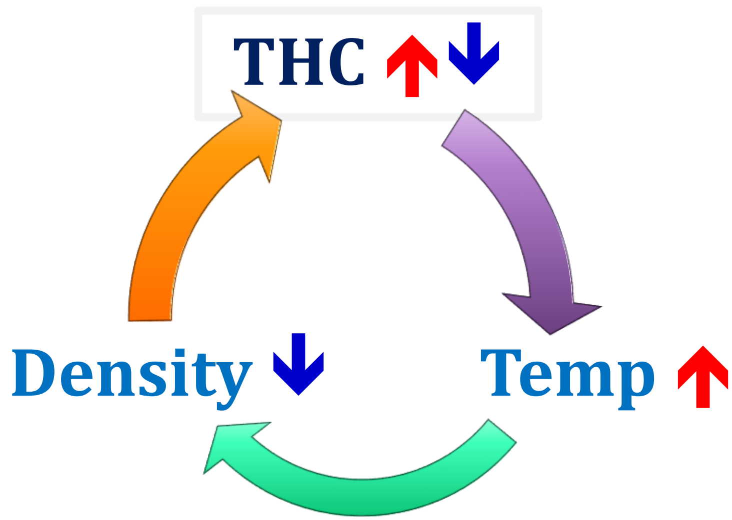
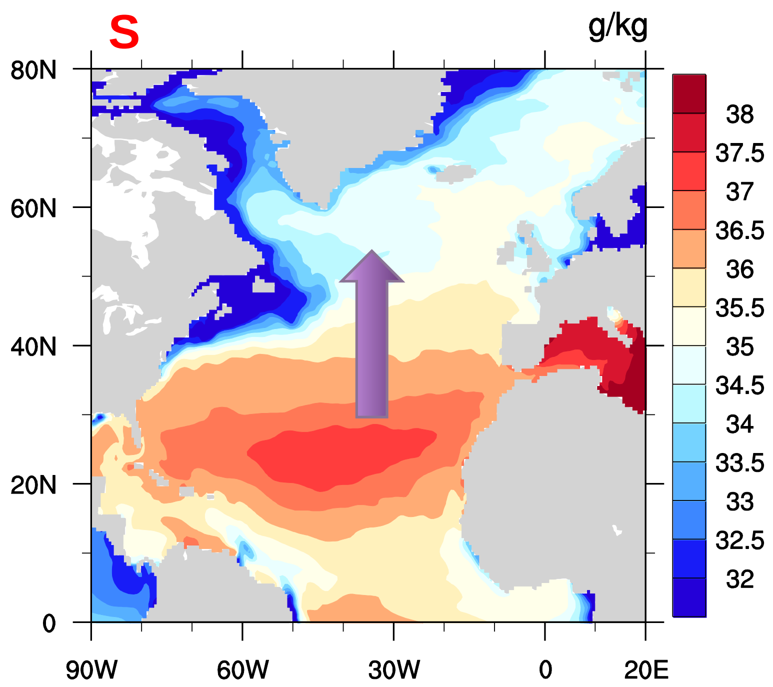
Negative Feedback: *THC* vs *T*

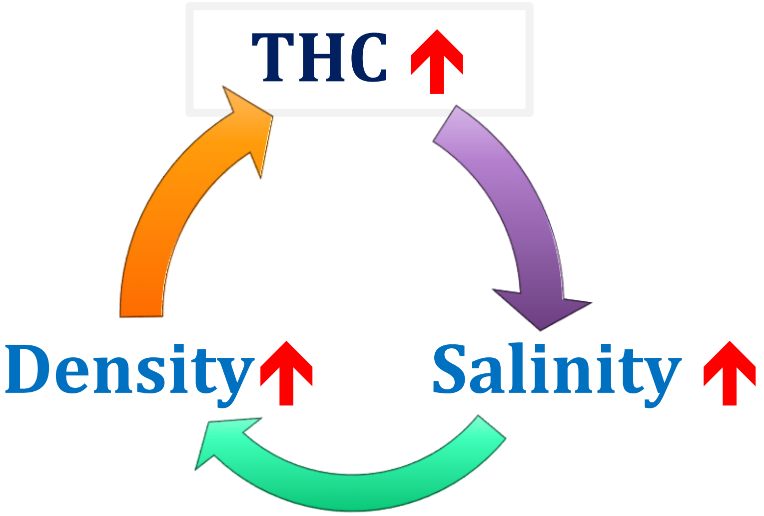




Willebrand (1993)

Positive Feedback: *THC* vs *S*

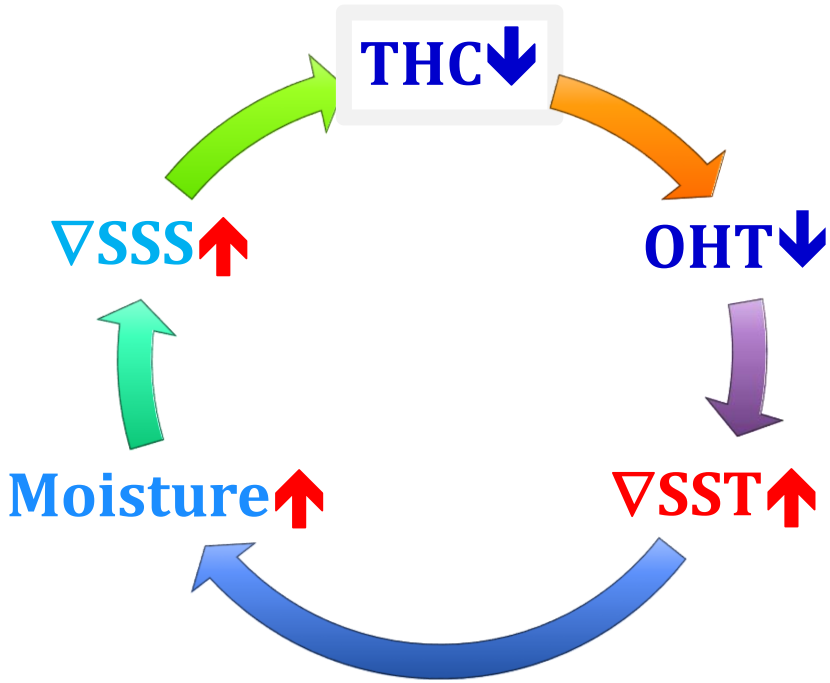




Willebrand (1993)

Positive Feedback: *THC* vs *EMT*

EMT: Atmosphere Eddy Moisture Transport



Nakamura et al. (1994)

Analytical Solution: BJC Rate

$$C_R \equiv \Delta H_a / \Delta H_o = -(B_1+B_2)\chi / \left[B_1 B_2 + (B_1+B_2)\chi\right]$$

if $B_1 = B_2$

$$C_R = -1/\left[1+B/2\chi\right] < 0 \quad \sim \quad \text{Intrinsic Rate}$$

Intrinsic parameters:

- $B_{1,2}$ ~ **Local feedback between *T* and *HF* at TOA**
- χ ~ **Atmospheric heat transport efficiency.**

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Intrinsic Rate

$$|C_R| \begin{cases} <1, \quad \textit{if } -B_1, -B_2 < 0, \quad \text{Undercompensation} \\ =1, \quad \textit{if } B_1 B_2 = 0, \quad \text{Full Compensation} \\ >1, \quad \textit{if either } -B_1 \text{ or } -B_2 > 0, \quad \text{Overcompensation} \end{cases}$$

$B_{1,2}$ ~ Local climate feedback **VITAL!**

1-D Slab-Ocean Energy Balance Model

Liu, Yang, He and Zhao, 2016: A theory for Bjerknes compensation: the role of climate feedback. J. Climate

Non-dimensional Equation

$$\partial_{xx}[M(x)T] - b(x)T + f(x) = 0, \qquad for \;\; 0 < x < 1$$

With Non-dimensional OHT and AHT

$$f(x) = -\partial_x H_O,$$

$$H_A = -\partial_x [M(x)T].$$

Stommel (1961); North, (1975)

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Analytical Solution: BJC Rate

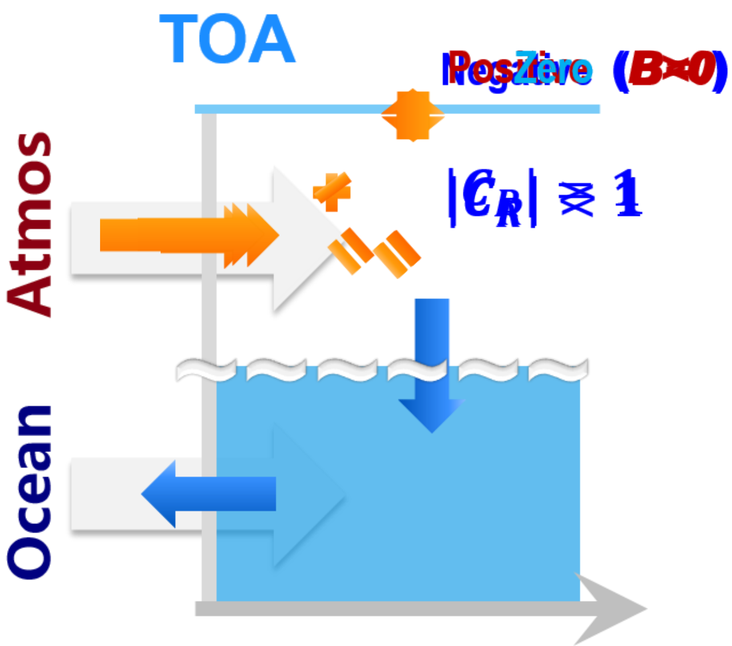
Intrinsic Rate

$$C_R \equiv \Delta H_a / \Delta H_o = -1/\left[1+B/(n\pi)^2\right] < 0,\ n=1,2,\ \ldots$$

- B ~ Intrinsic parameter, Local feedback between T and HF at TOA
- n ~ Forcing scale.

Coupled Intrinsic Mode

Thermohaline-Climate Feedback-Energy Balance



$$C_R \equiv \frac{\Delta F_a}{\Delta F_o} = -\frac{1}{1-B} < 0$$

Local climate feedback $B(y)$

Yang, Zhao and Liu (2016)
Zhao, Yang and Liu (2016)

Climate Feedback + MHT ➔ Earth Energy Balance

Coupled Intrinsic Mode

能量补偿 ⇔ 体重保持

饮食
(海洋)

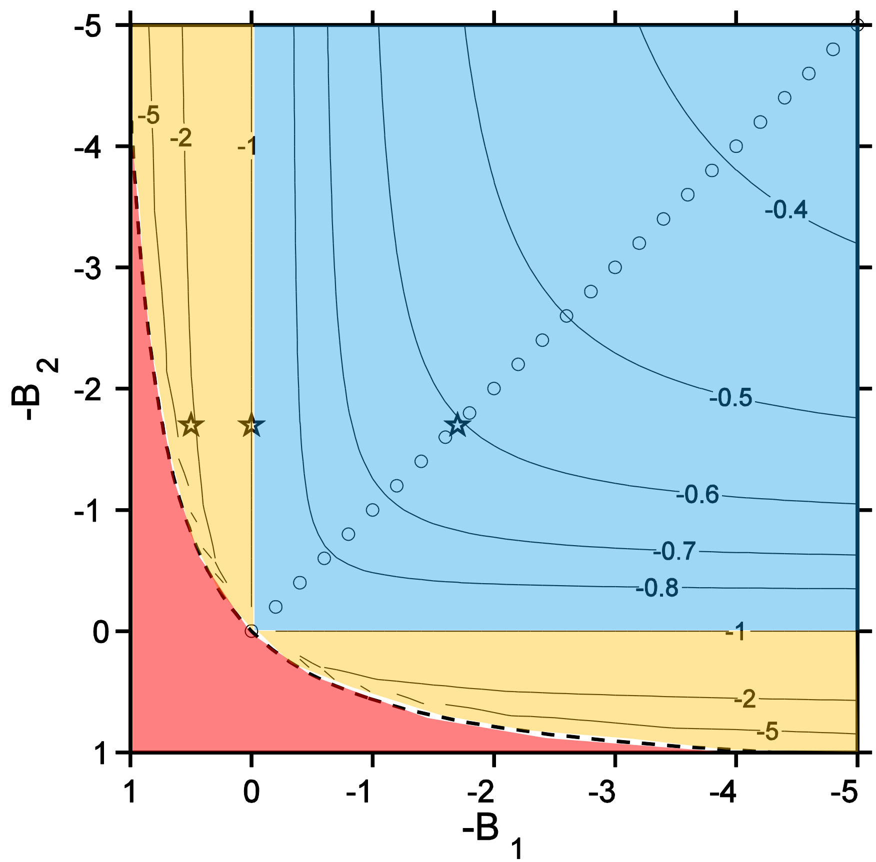
新陈代谢
(气候反馈)

运动
(大气)

$$C_R = -\frac{1}{1-B}$$

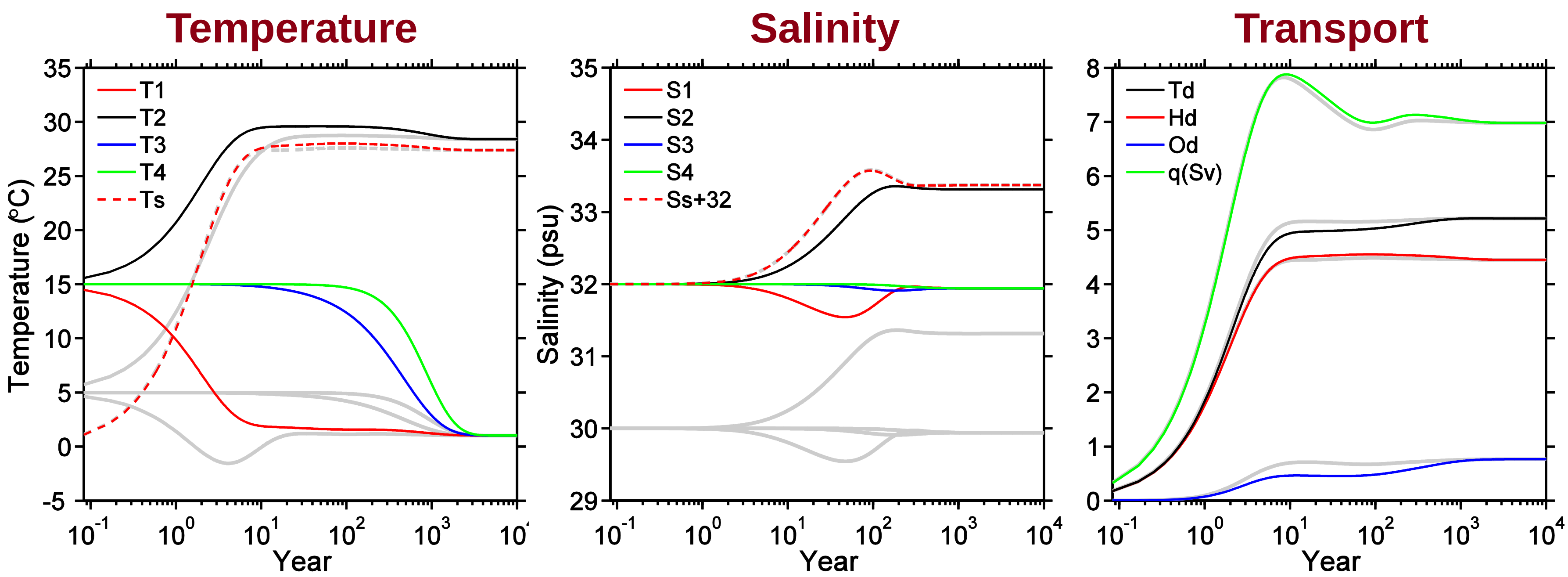
Go to BJC for Climate Variability

Intrinsic Rate

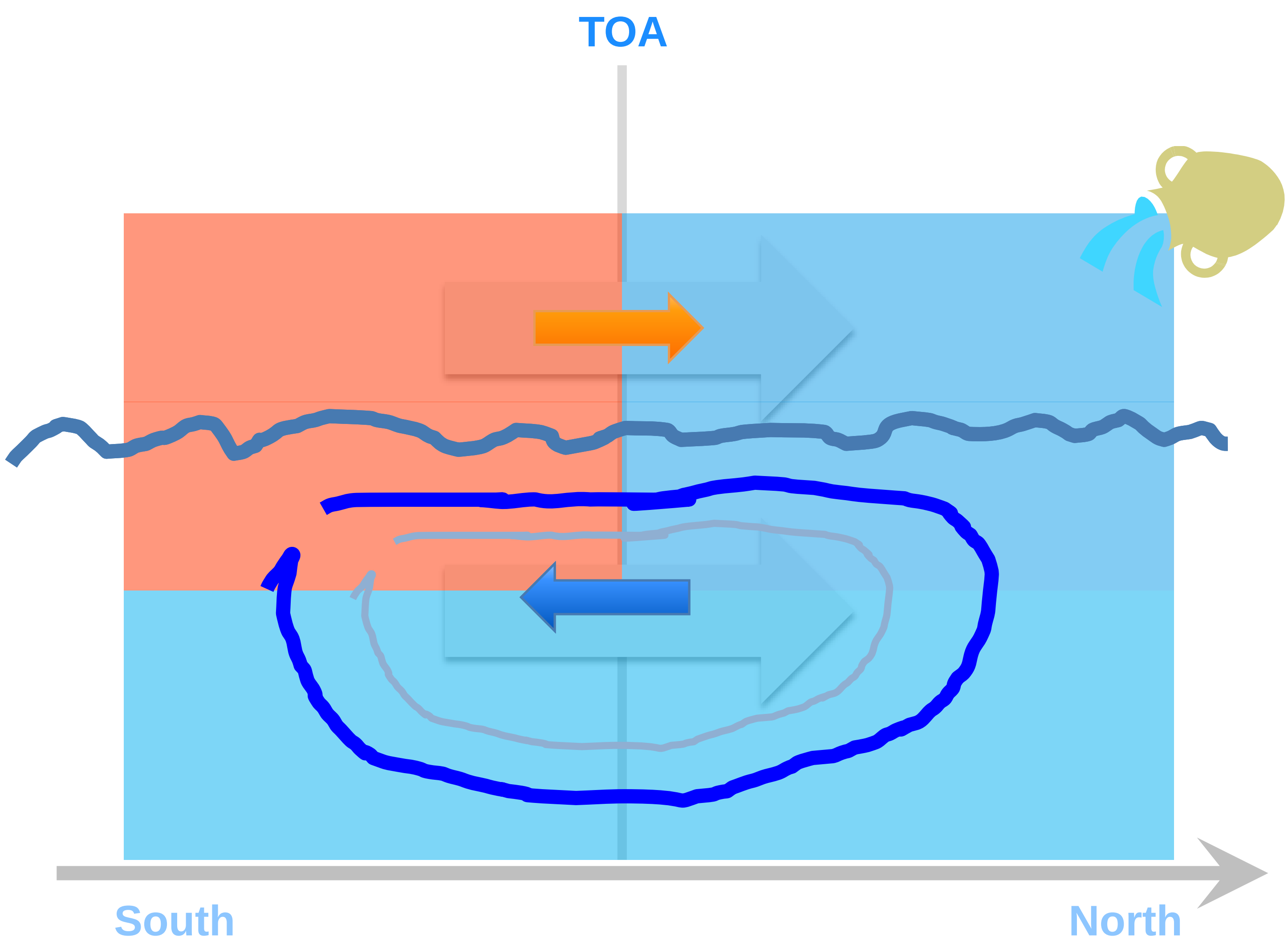


$$C_R = -(B_1+B_2)\chi/[B_1B_2+(B_1+B_2)\chi]$$

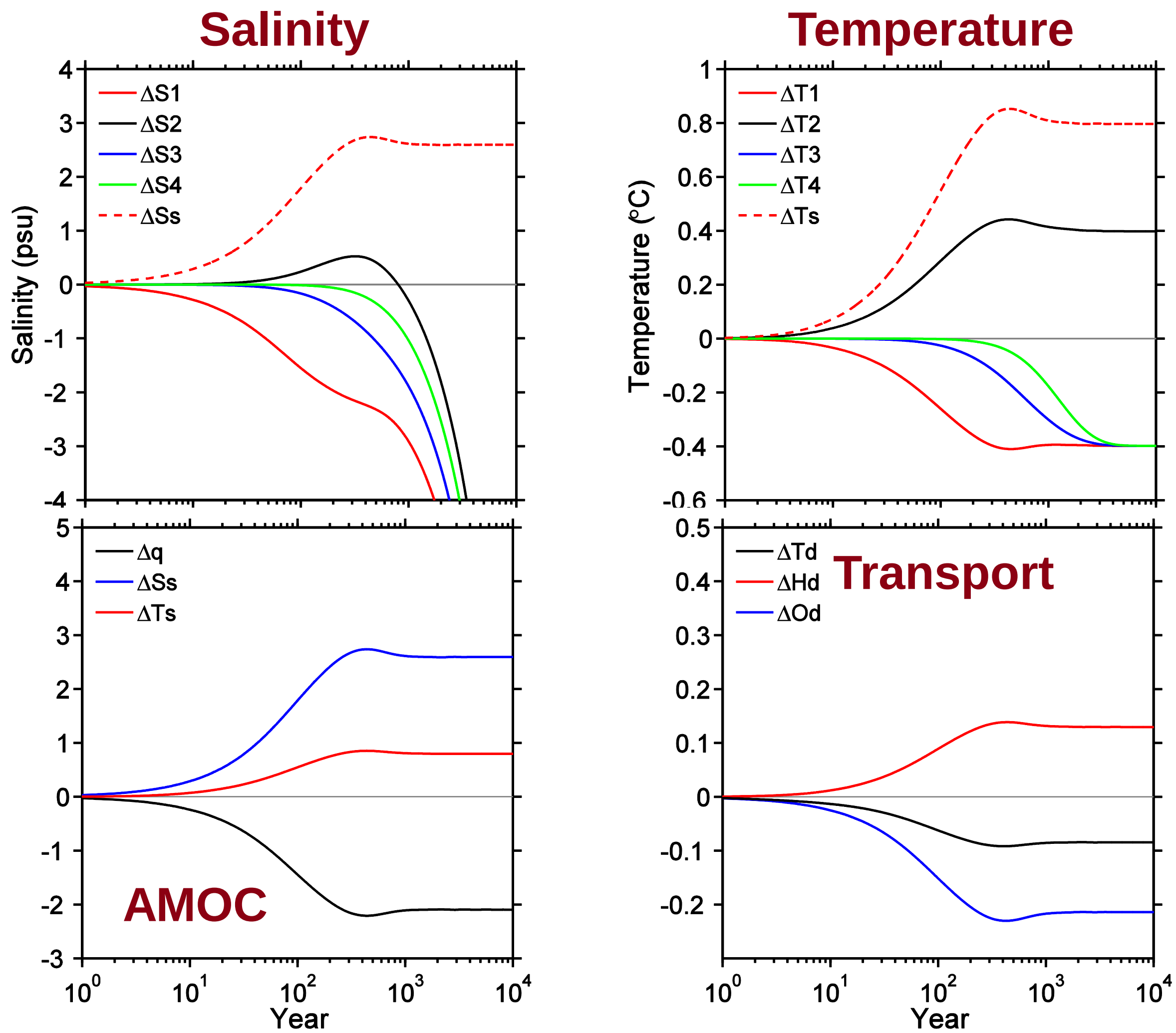
Mean Climate



How BJC Occurs: Fresh Water

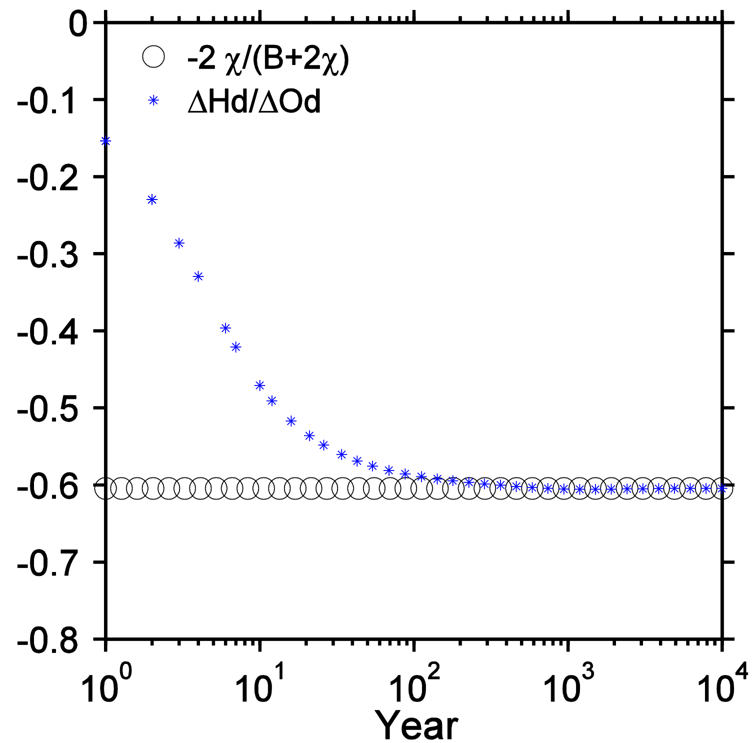


Climate Change under Fresh Water



BJC under Negative Feedback

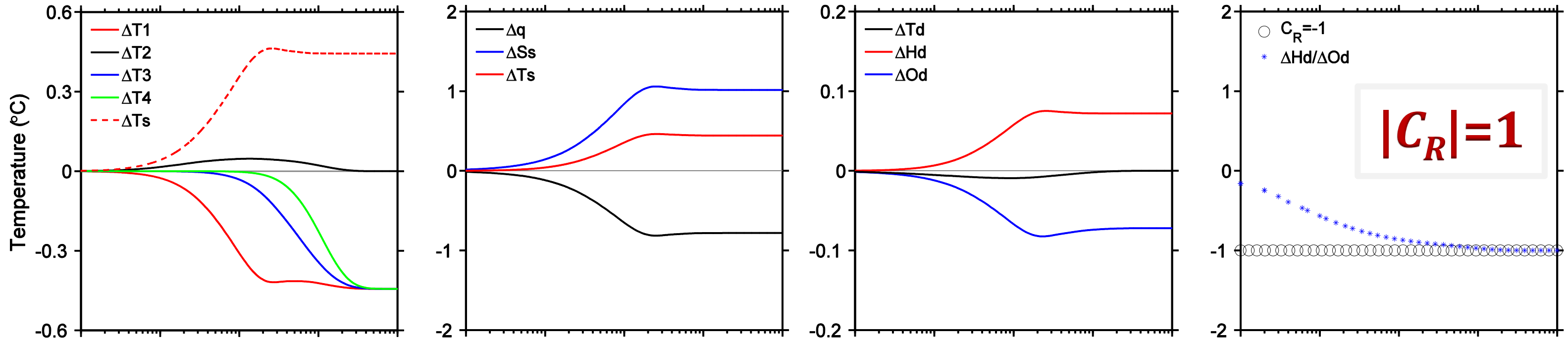
Compensation Rate: Constant!



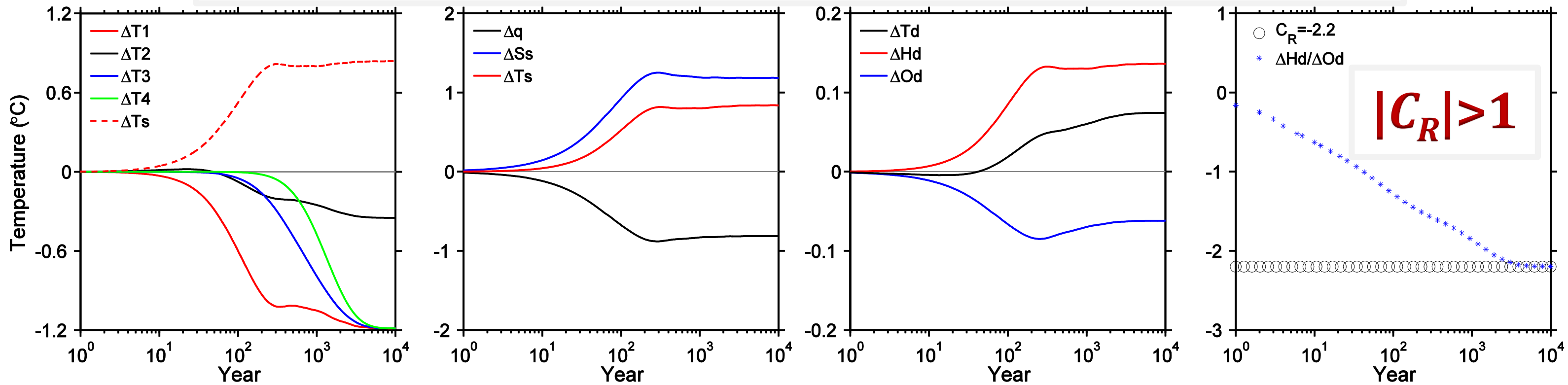
Under-Compensation (-B1,2<0)

Climate Change under Fresh Water

Full Compensation (-B1=0, Zero Feedback)



Overcompensation (-B1=0.5, Positive Feedback)



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Modeling MHT and OHT in Past 22ka

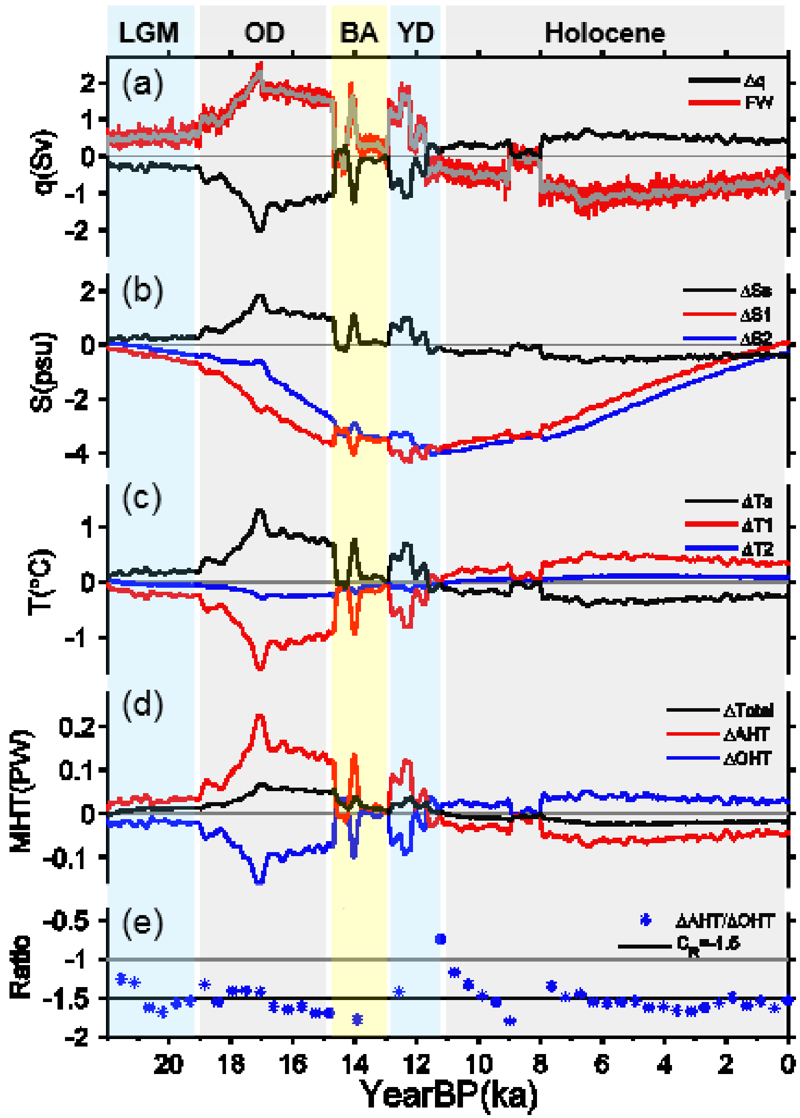
Freshwater Forcing

Salinity Change

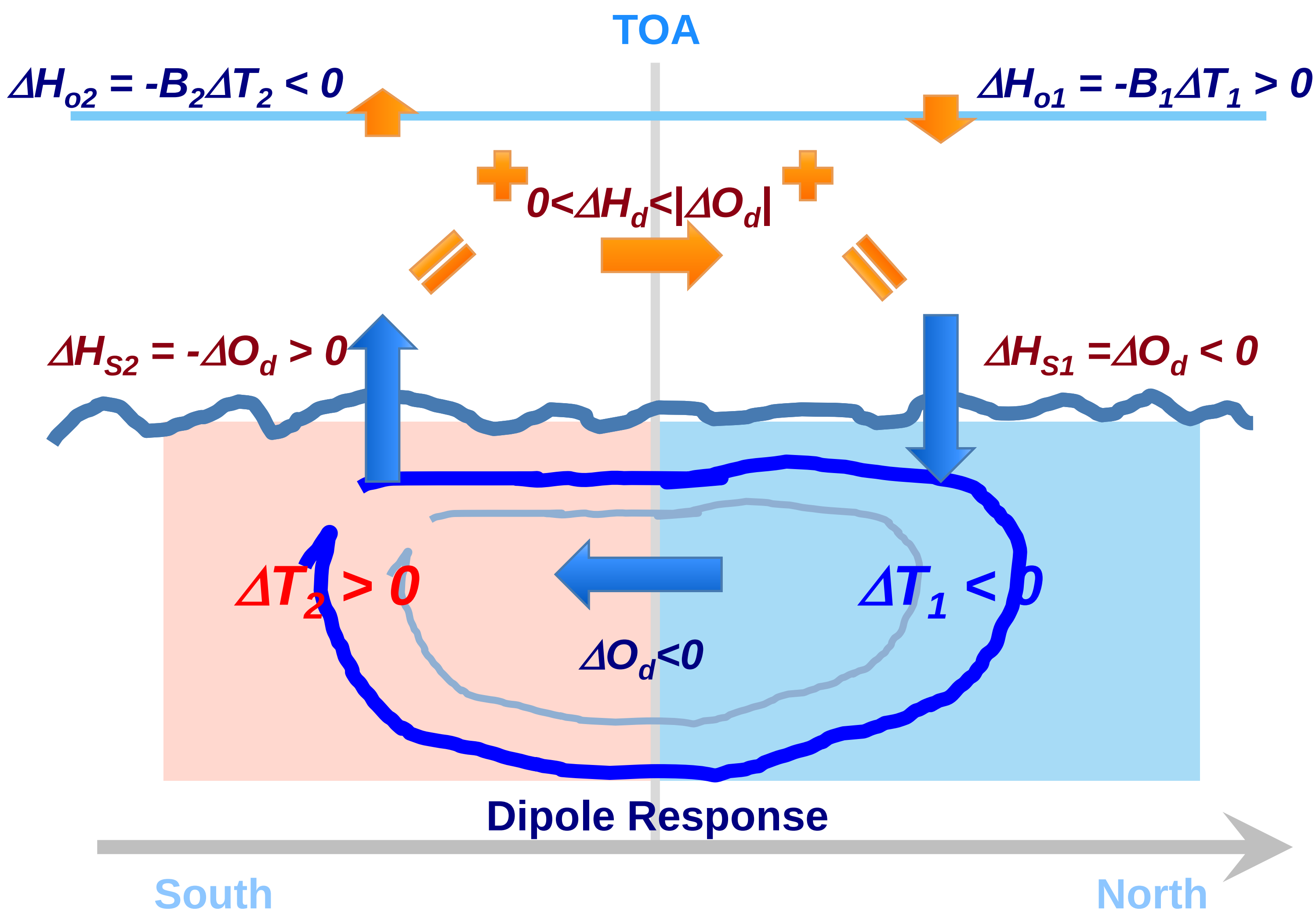
Temperature Change

AHT and OHT

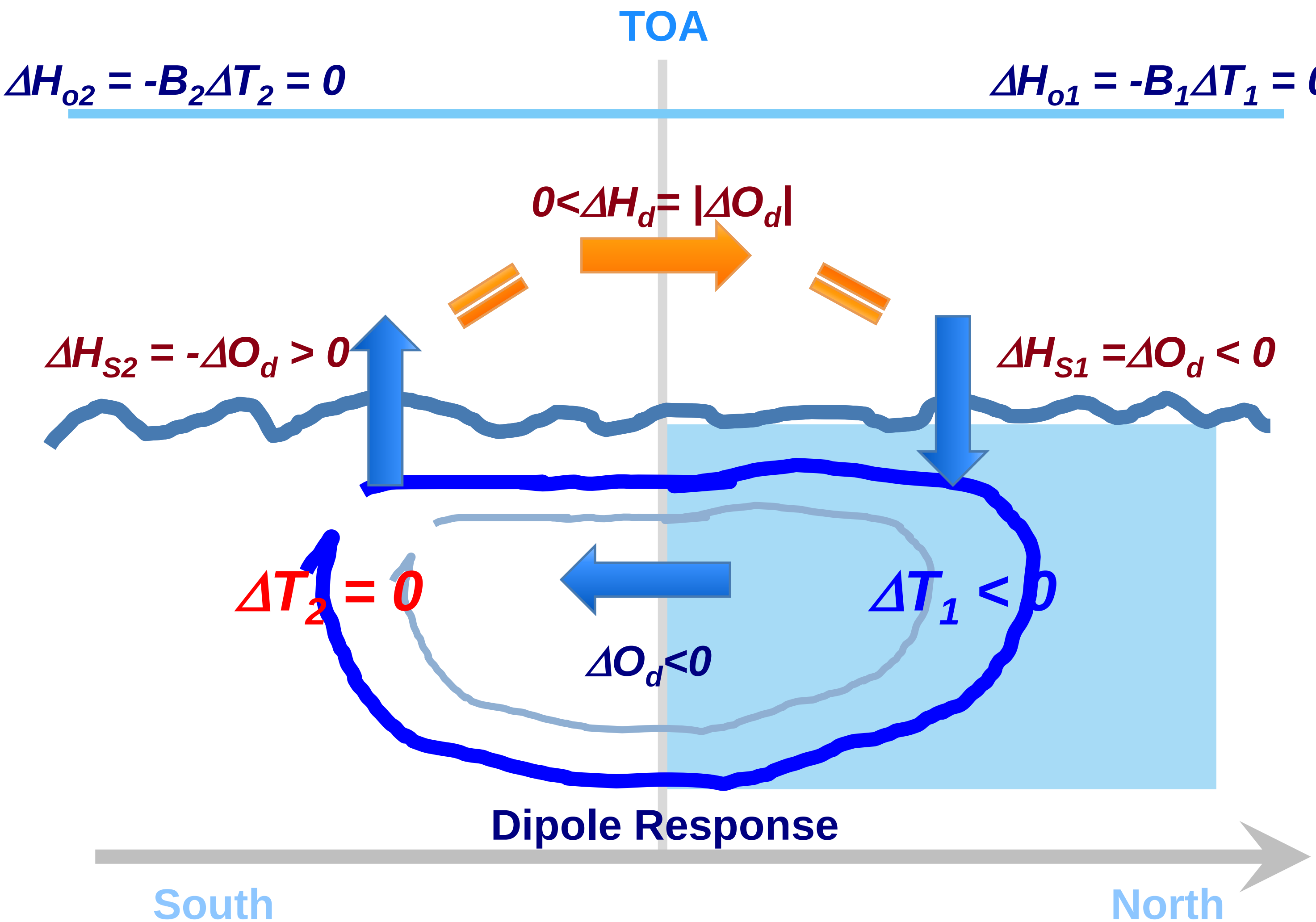
BJC = -1.5



Mechanism for Undercompensation



Mechanism for Full compensation



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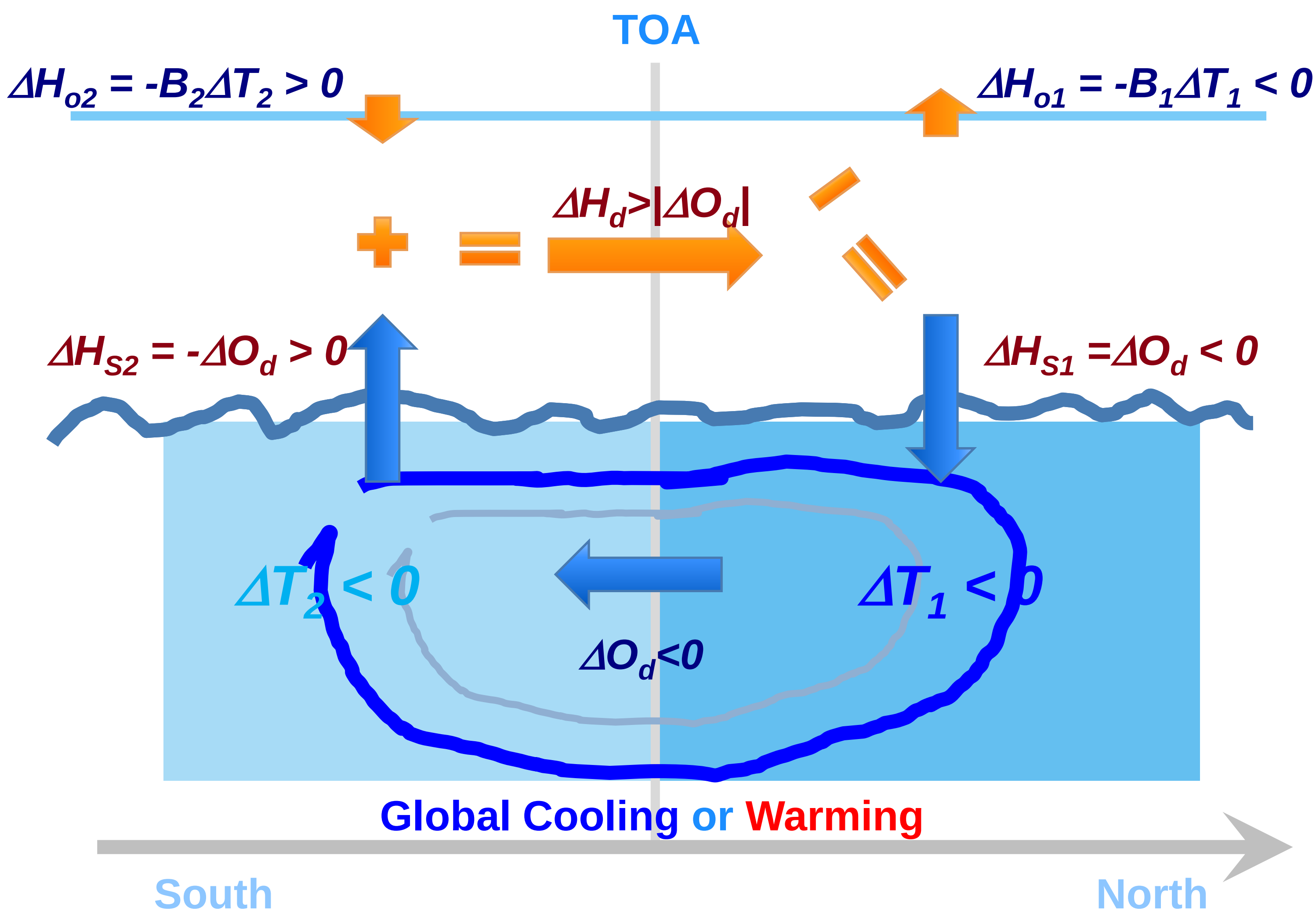
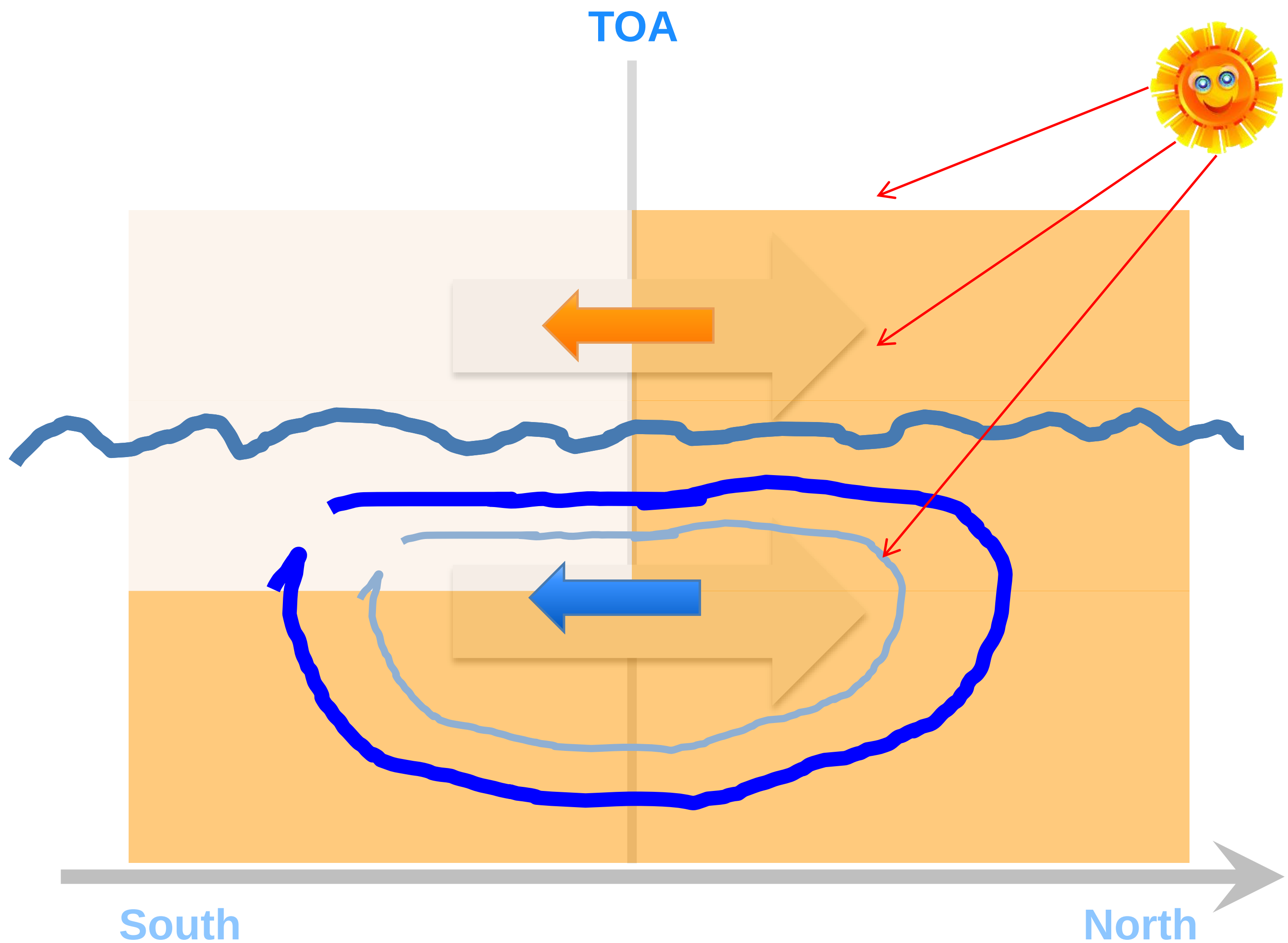
Mechanism for Overcompensation



Will BJC Occur under Heating?



Analytical Solution: BJC Rate under Heating

$$C_h \equiv \Delta H_a/\Delta H_o = \boxed{\{1/[1+(B_2-B_1)\ \Delta A/\Delta T_S\,[B_1B_2+(B_1+B_2)\chi]]\}}\ C_R$$

if $B_1=B_2$ global uniform feedback

$$C_h = C_R < 0$$

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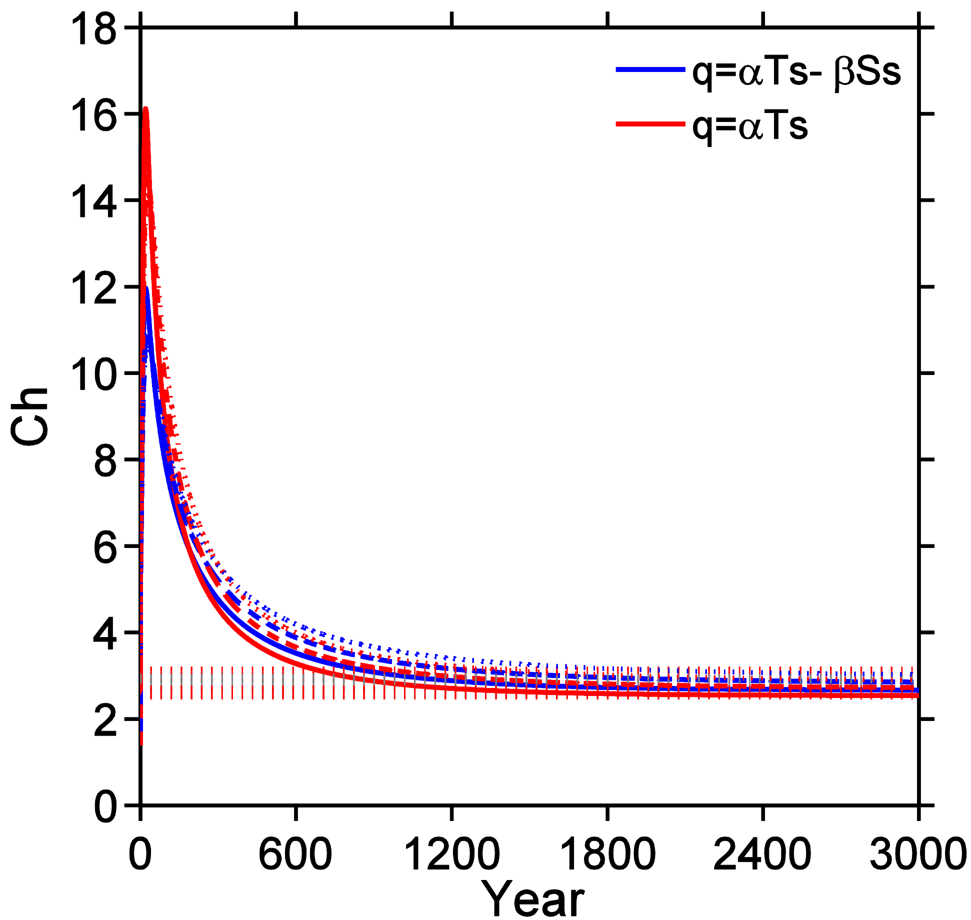
"BJC Rate" under Heating

No thermohaline dynamics ➔ No Compensation!



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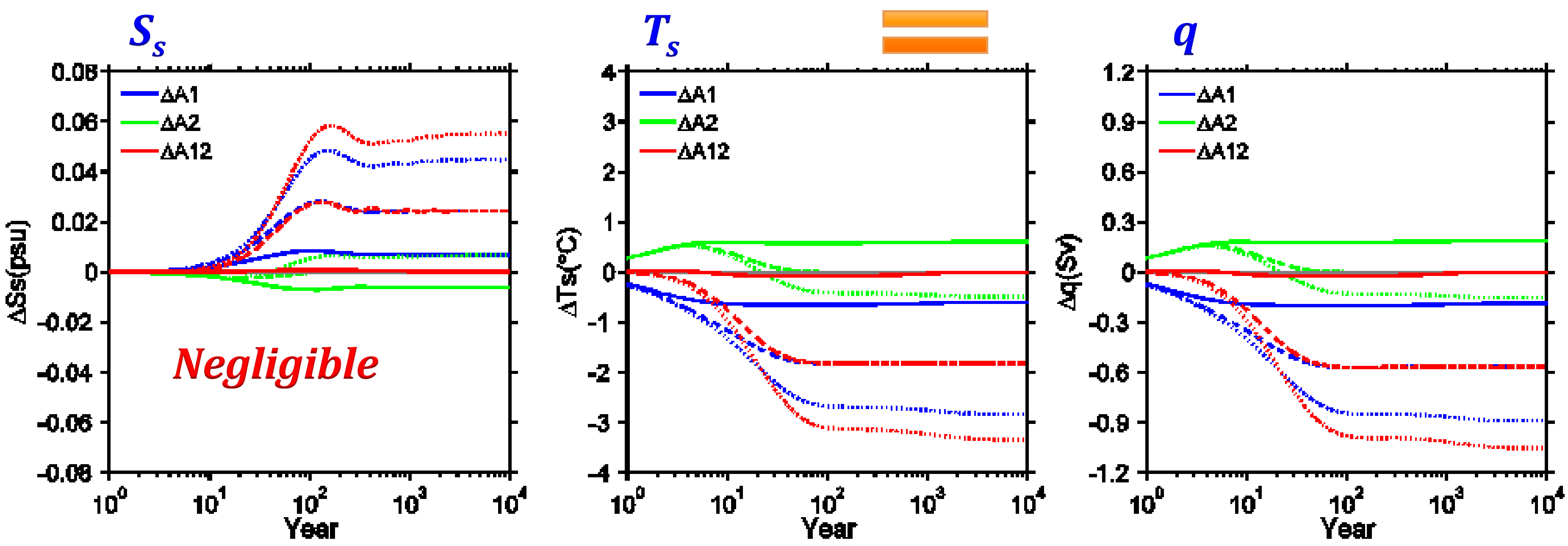
Why *No* BJC under Heating?



$$q \sim T_s \rightarrow O_d \sim q{*}T_s \sim T_s{*}T_s$$

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Why *No* BJC under Heating?

$$q \sim T_s \Rightarrow O_d \sim q^*T_s \sim T_s^*T_s$$

$$C_q \equiv \Delta H_a / \Delta H_o = H_a / 2H_o$$

No Compensation! $C_q > 0$ & $C_q \sim$Constant

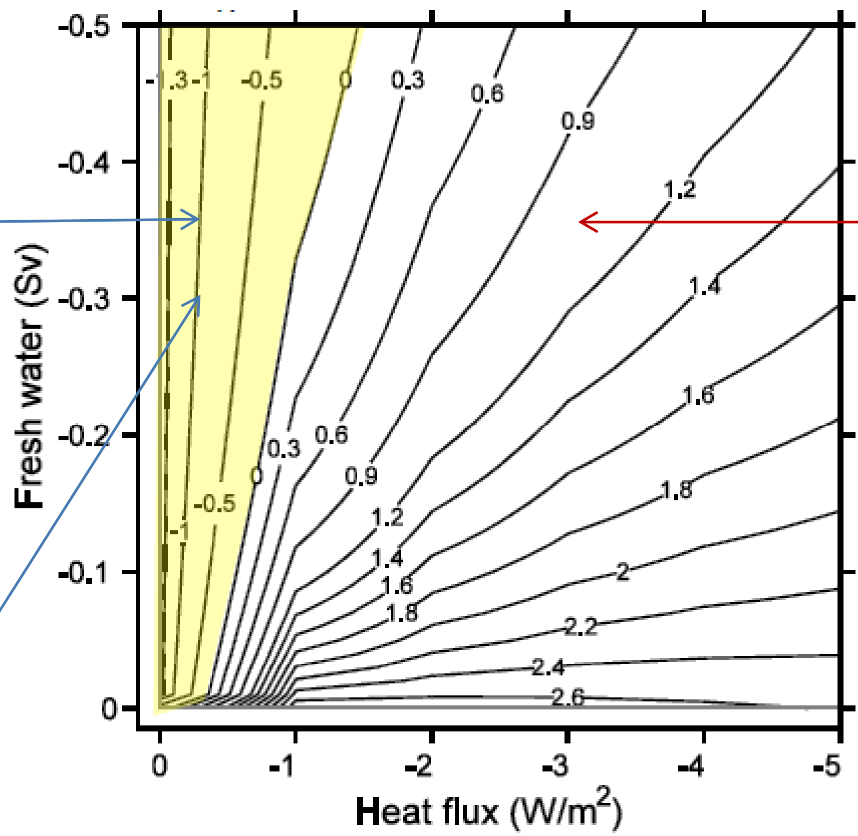
No thermohaline dynamics

BJC Possible under Freshwater + Heating?

Valid for weak heating

BJC

No BJC



Positive Feedback: S1⇑ ➔ THC⇑ ➔ S1⇑ ➔ THC⇑ ➔ Ts⇓

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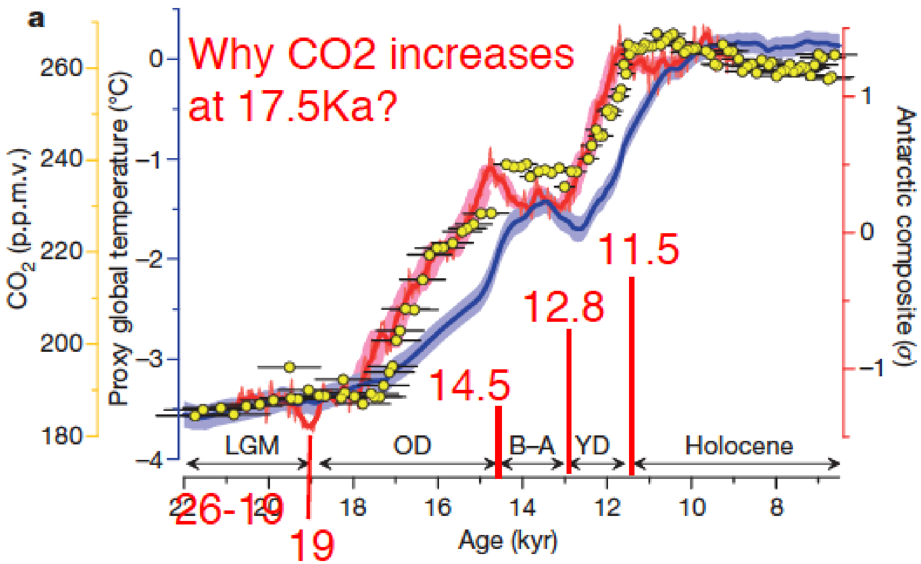
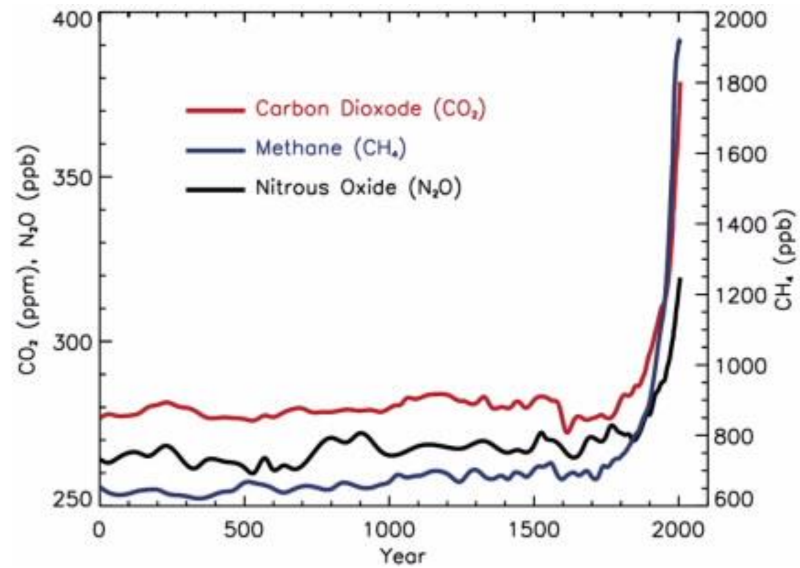
CO2 since the LGM

Since LGM: 1.5×10^{-4} W/m^2/year (Shakun et al., 2012)





Since 1770: 1×10^{-2}W/m^2/year (IPCC, 2007)

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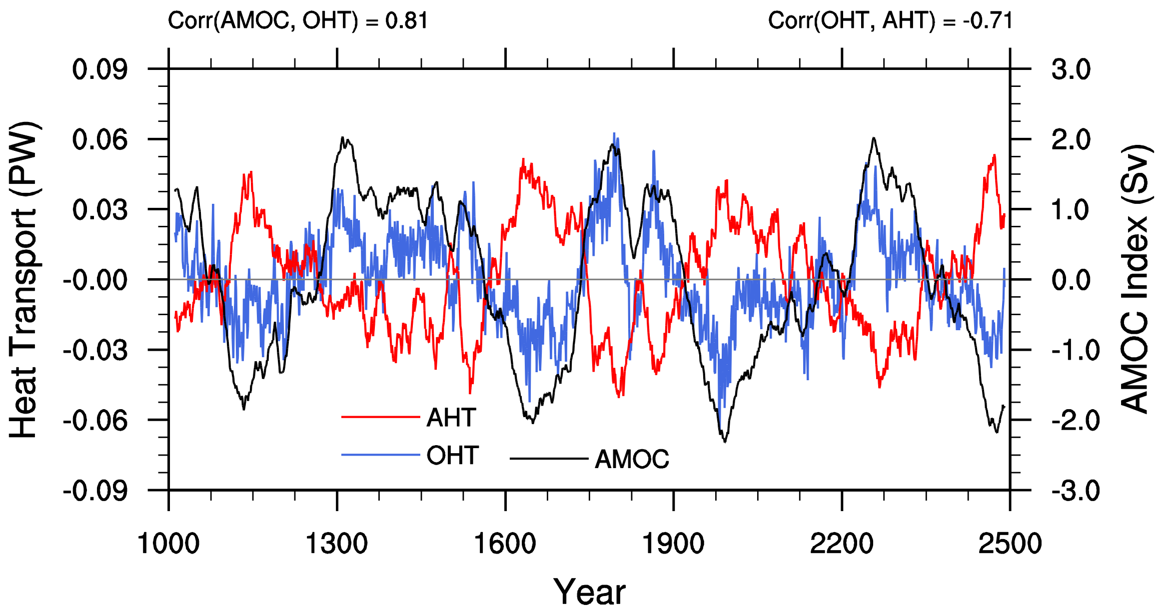
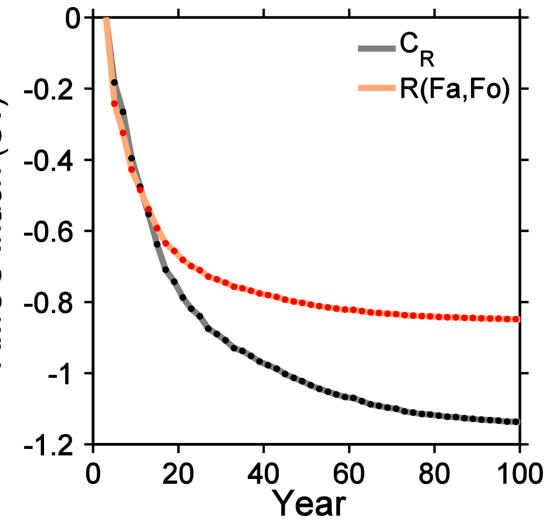
Climate Variability and BJC in CESM





Define: $$C_R \equiv r\frac{\sigma_{F_a}}{\sigma_{F_o}}$$ **(1)**

Theory for Transient Climate Variability

$$\dot{\boldsymbol{T}}_{\boldsymbol{s}} = \frac{1}{\epsilon c \rho_0 D_1}[(A_2 - A_1 - B\boldsymbol{T}_{\boldsymbol{s}}) - 2\chi \boldsymbol{T}_{\boldsymbol{s}}] - 2q\boldsymbol{T}_{\boldsymbol{s}},$$

$$\dot{\boldsymbol{S}}_{\boldsymbol{s}} = \frac{2S_0}{\epsilon_w D_1}\gamma \boldsymbol{T}_{\boldsymbol{s}} - 2q\boldsymbol{S}_{\boldsymbol{s}} + \boldsymbol{h}_{\boldsymbol{fw}}.$$

Linearization: $T = \bar{T} + T'$ and $S = \bar{S} + S'$

$$\frac{\partial}{\partial t}\begin{pmatrix} \boldsymbol{T'_s} \\ \boldsymbol{S'_s} \end{pmatrix} = \boldsymbol{M}\begin{pmatrix} \boldsymbol{T'_s} \\ \boldsymbol{S'_s} \end{pmatrix} + \begin{pmatrix} \boldsymbol{0} \\ \boldsymbol{h_0 e^{i\omega t}} \end{pmatrix}$$

If $\boldsymbol{h}_{\boldsymbol{fw}} = const. \longrightarrow$

$$C_{R0} = -\frac{1}{1+B/2\chi} \qquad \textbf{(2)}$$

BJC for Climate Variability

Zhao, Yang and Liu, 2016: Assessing Bjerknes compensation for climate variability and its timescale dependence. J. Climate

$$C_{Rp} \equiv \frac{F'_a}{F'_o} = Re\left(C_{R0}e^{i\delta}\right) = r_\delta * C_{R\omega}$$

$$r_\delta \equiv cos\delta = -\frac{F}{\sqrt{\omega^2 + F^2}}$$

$$C_{R\omega} = \frac{2\chi}{\epsilon c\rho_0 D_1\sqrt{\omega^2 + F^2}}$$

$$C_R \equiv r\frac{\sigma_{F_a}}{\sigma_{F_o}}$$

$\omega \to 0$

$$C_{R0} = -\frac{1}{1 + B/2\chi}$$

BJC for Climate Variability

$$\boldsymbol{\omega} \to \infty \quad \Rightarrow \; r_\delta \to \mathbf{0};\;\; C_{R0} \to \mathbf{0};\; C_{R\boldsymbol{p}} \to \mathbf{0}$$

No correlation and No BJC

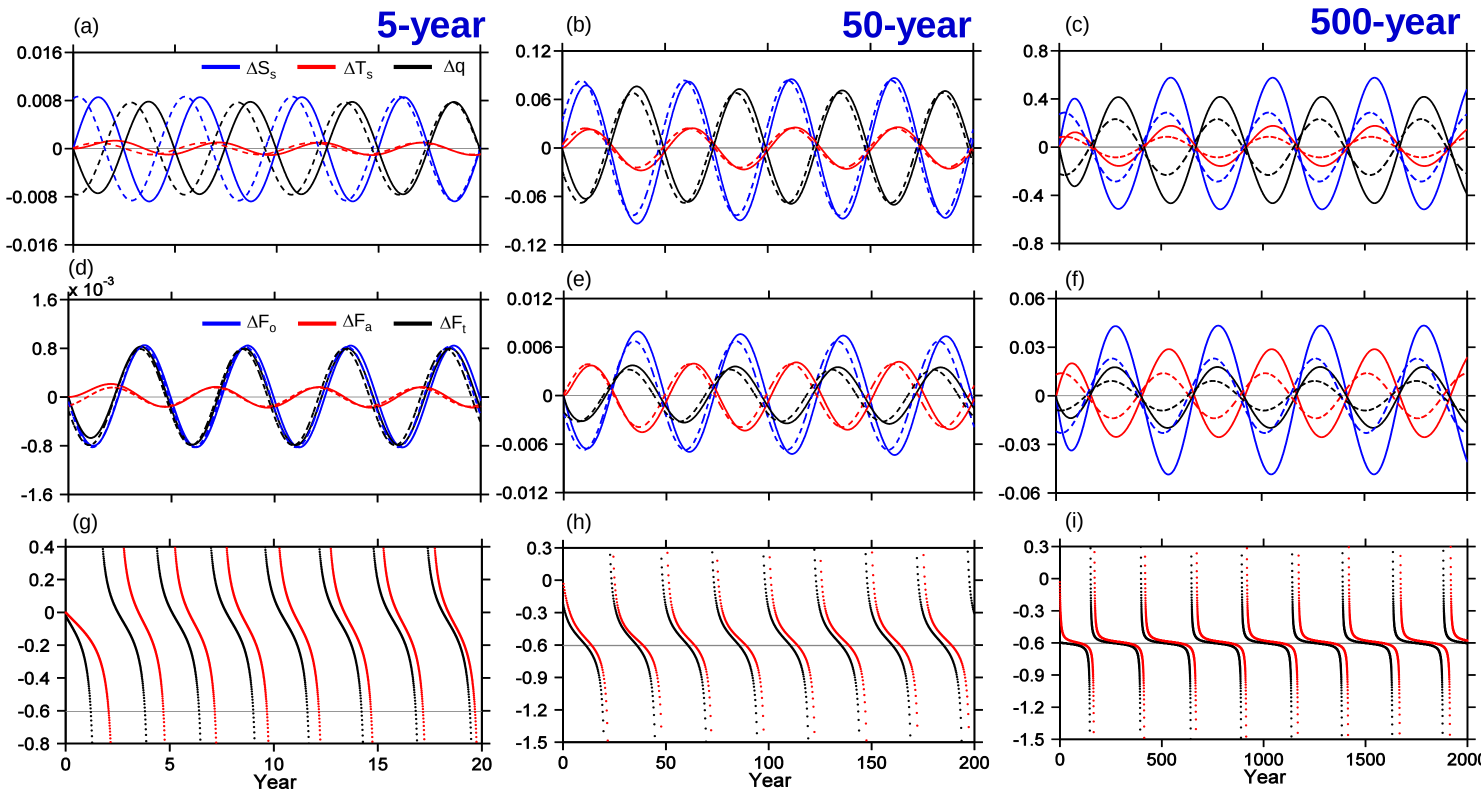
$$\boldsymbol{\omega} \to \mathbf{0} \quad \Rightarrow \; r_\delta \to -\mathbf{1};\;\; C_R(\mathbf{1}) \asymp C_{R\mathbf{0}}(\mathbf{2}) \asymp C_{R\boldsymbol{p}}(\mathbf{3})$$

Full correlation and equilibrium BJC

Go to Climate Variability Validation

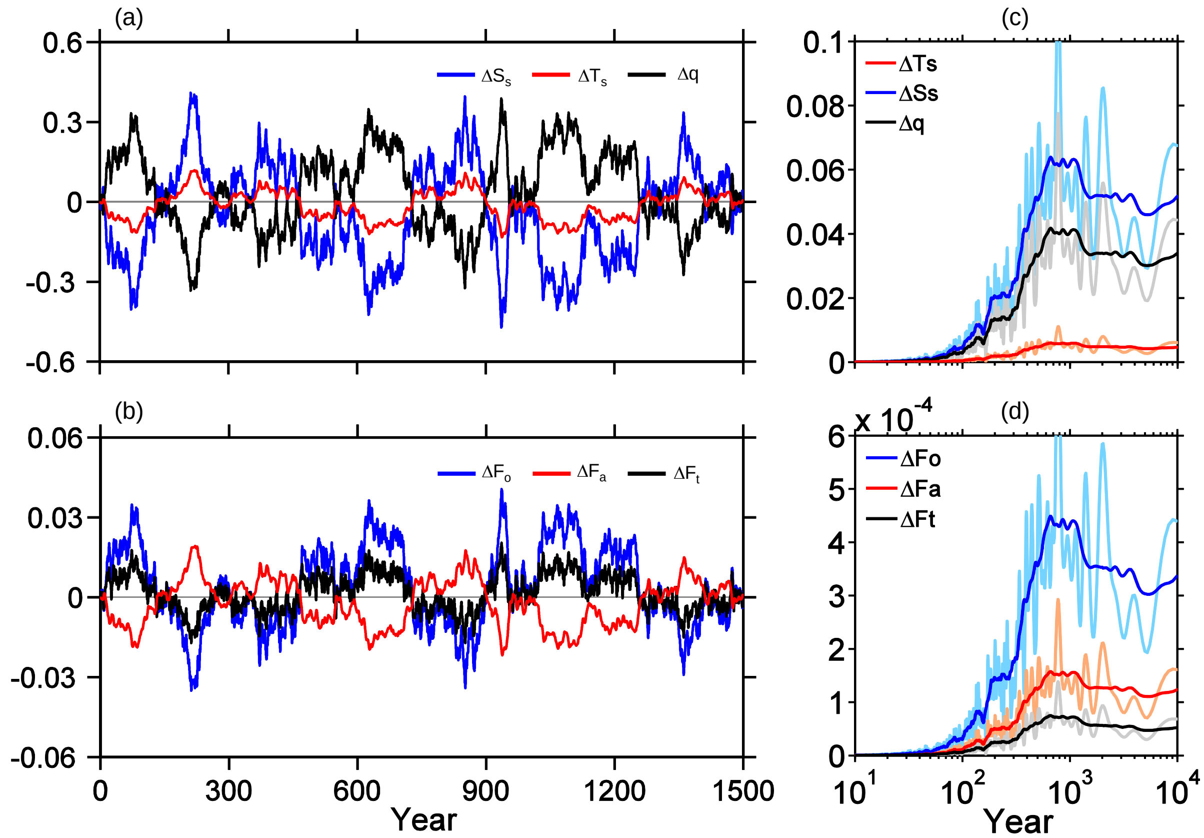
BJC for Climate Variability

Under periodic forcing: $h_{fw} = Re(h_0 e^{i\omega t})$



BJC for Climate Variability

Under stochastic forcing: $h_{fw} = h_0 \int_{\omega=0}^{\omega=\infty} e^{i\omega t}\, d\omega$



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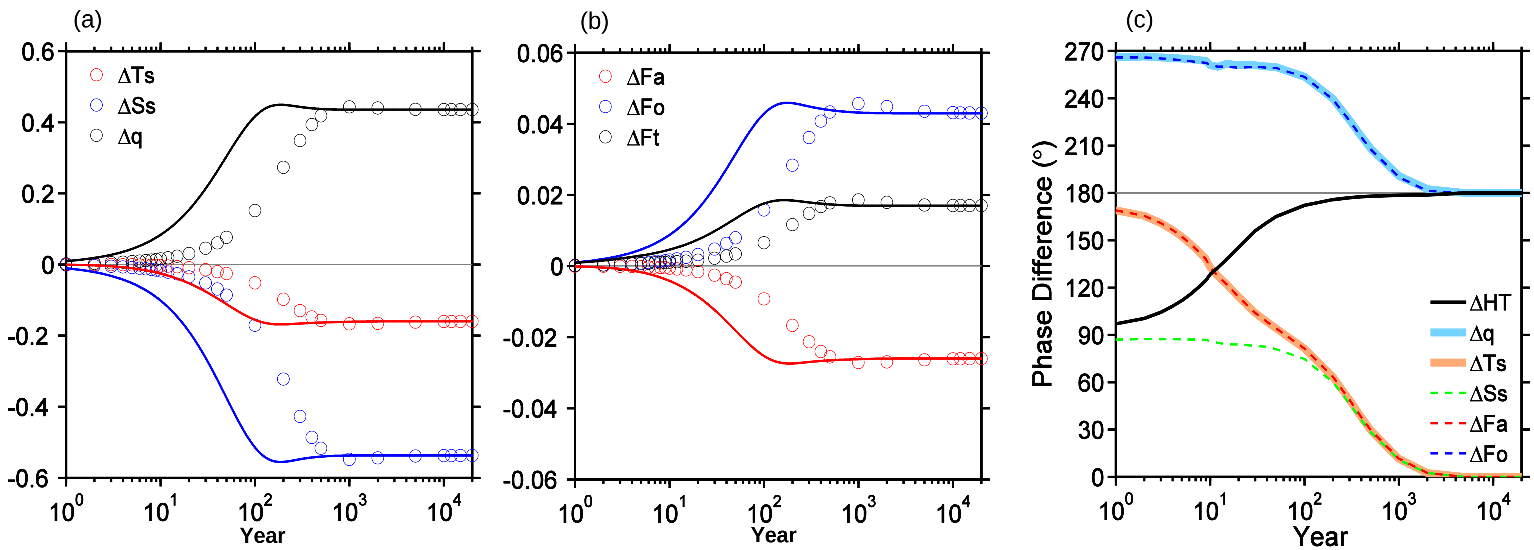
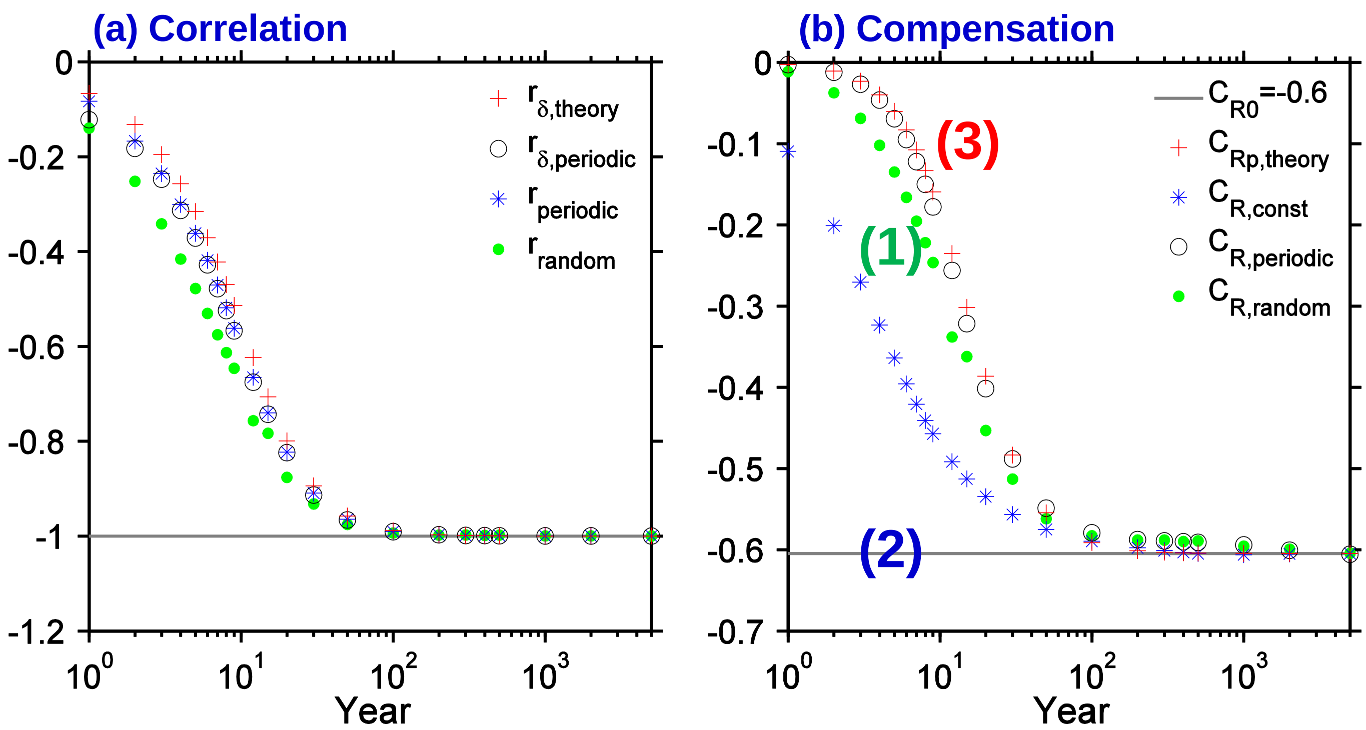
BJC for Climate Variability



Forcing frequency $\omega \rightarrow 0 \;\Rightarrow$ Equilibrium solution

BJC Theory Valid for Climate Variability



Beyond *decadal* timescale, AHT and OHT out of phase, BJC established

Outline

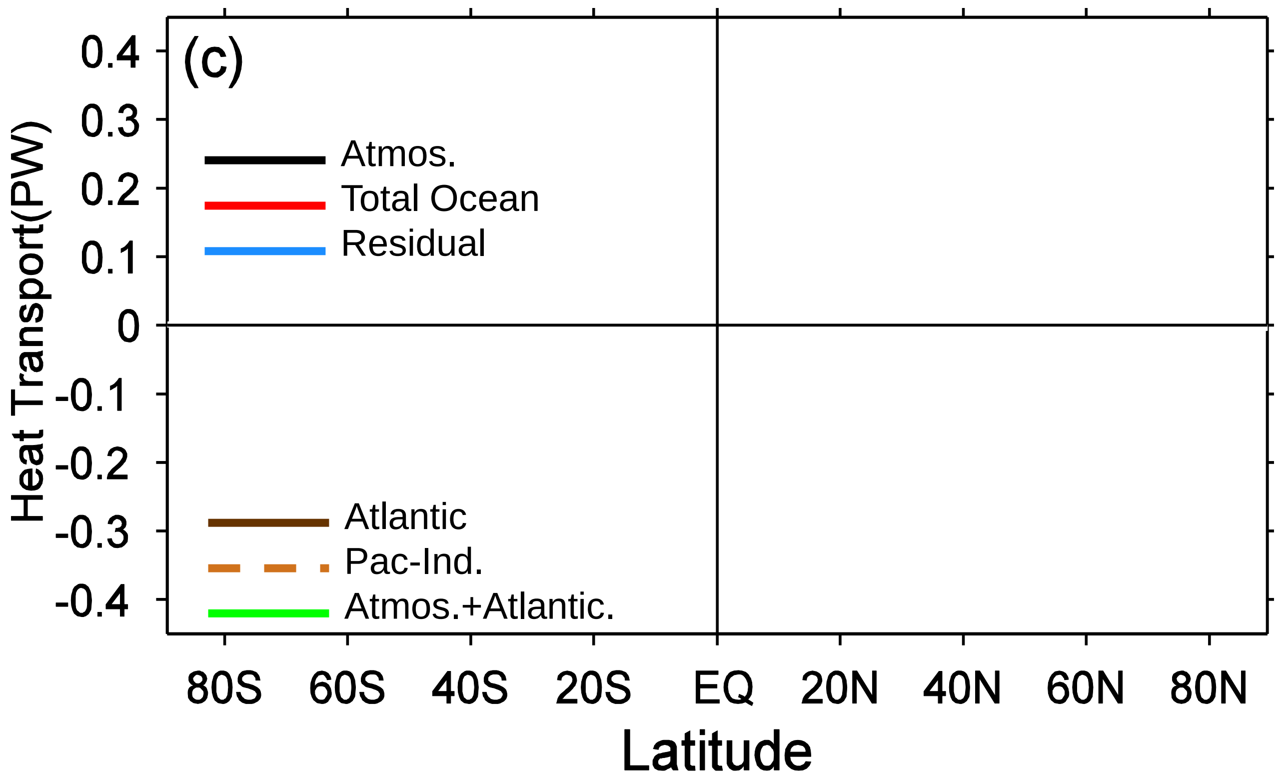
- **Fundamentals**
- **Questions**
- **Hypothesis and Theory**
- **CGCM results**
- **Aquaplanet**
- **Summary**

CESM Experiment and Mechanism

- Fresh-water experiments
- Wind-perturbation experiments
- Global warming experiments
- Internal variability from a long control run

Yang and Dai (2015), Yang et al. (2013, 2016, 2017)

BJC under Freshwater in CESM



80-90%

Atlantic OHT ↓ ⇒ AHT ↑ ⇒ Pac-Ind. OHT ↑

Atlantic OHT ≈ AHT; Pac-Ind. OHT ≈ Overcompensation

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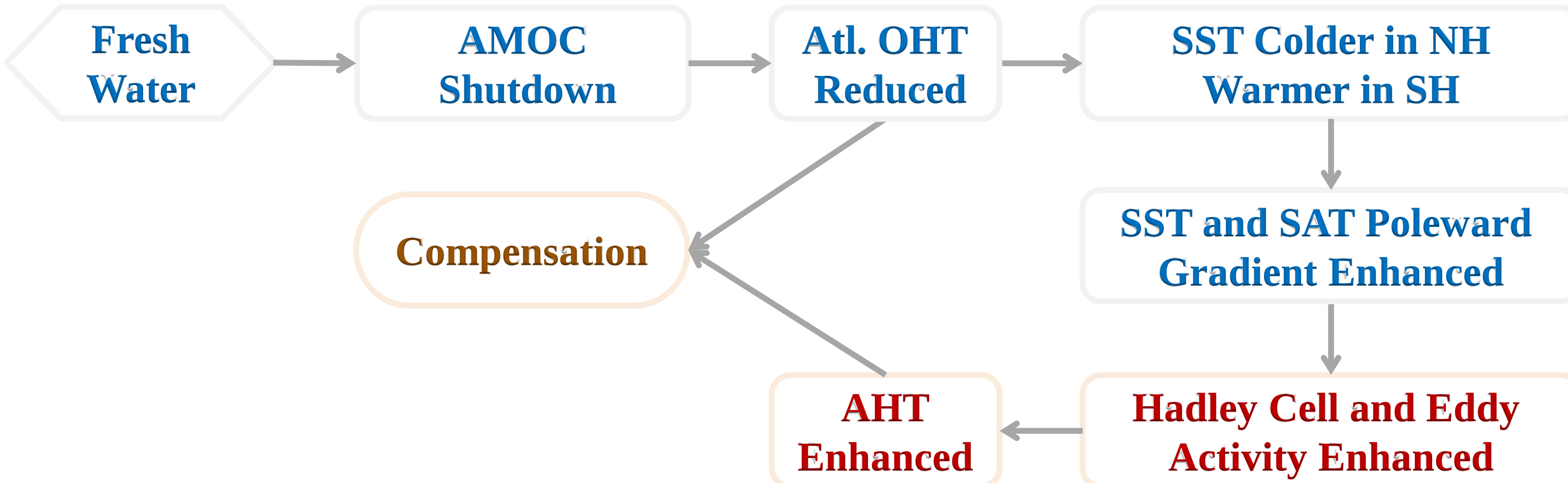
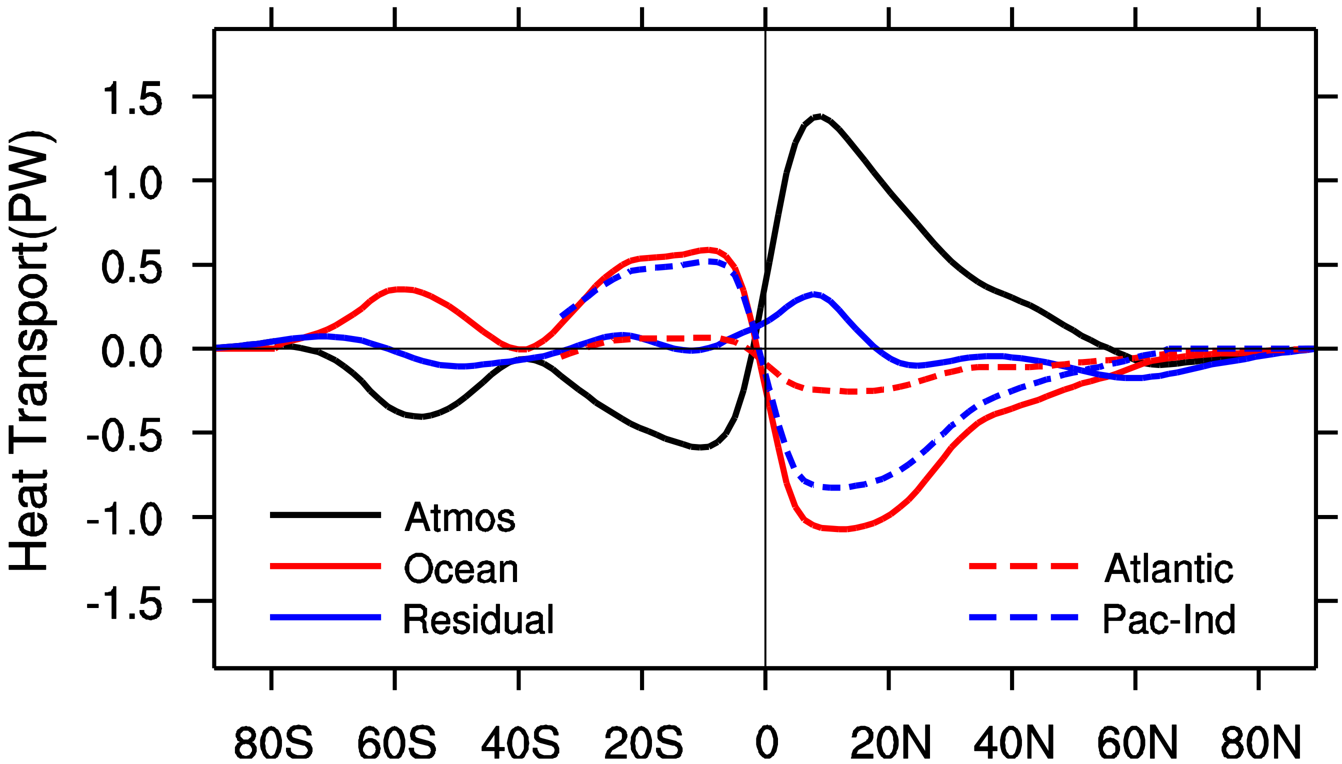
“Mechanism”



Yang et al. (2013, 2017)

BJC under Wind Perturbation in CESM

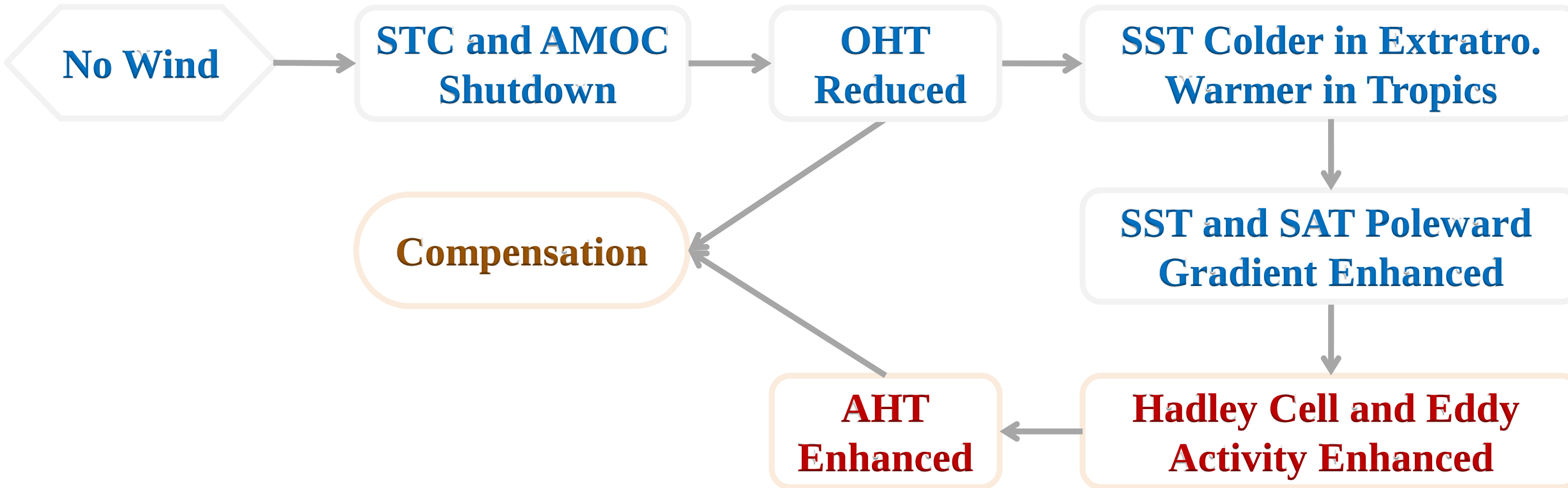


80-90%

Pacific-Indian OHT ↓ ⇒ Ty ↑ ⇒ HC ↑ ⇒ AHT ↑

Nearly Compensation

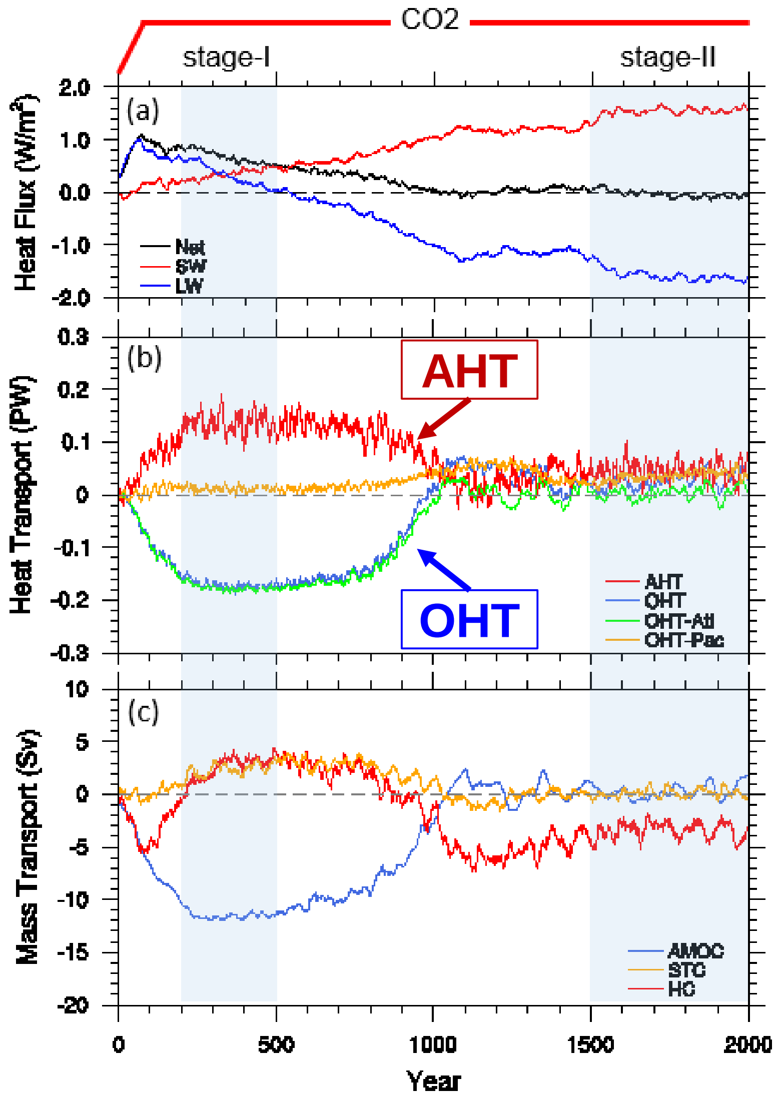
"Mechanism"



Yang and Dai (2015), Dai et al. (2017)

BJC under Global Warming in CESM

85%



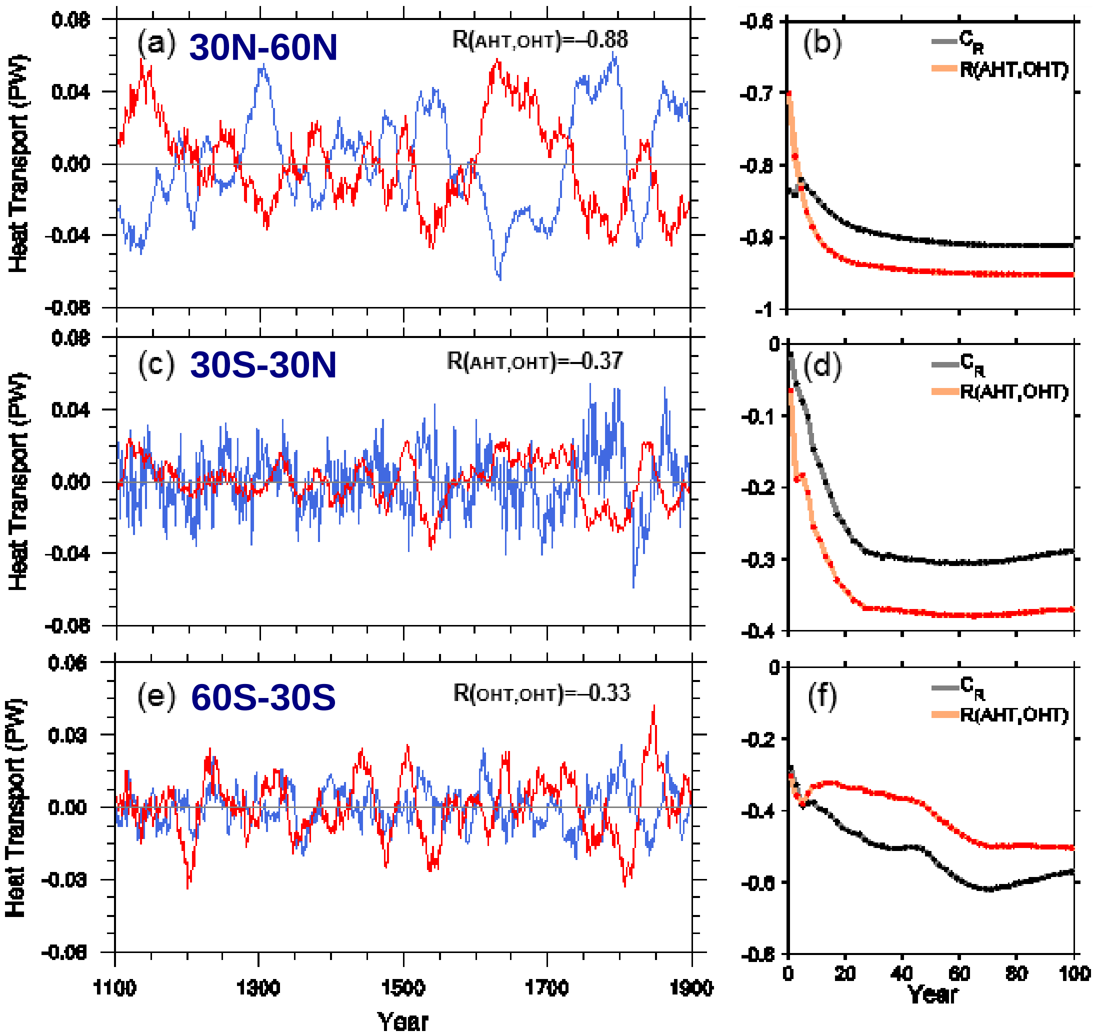
Yang et al. (2017)

Under 2 X CO2 forcing

BJC valid due to *thermohaline dynamics*!

BJC in Natural Variability in CESM

2000-year CESM control run



95%

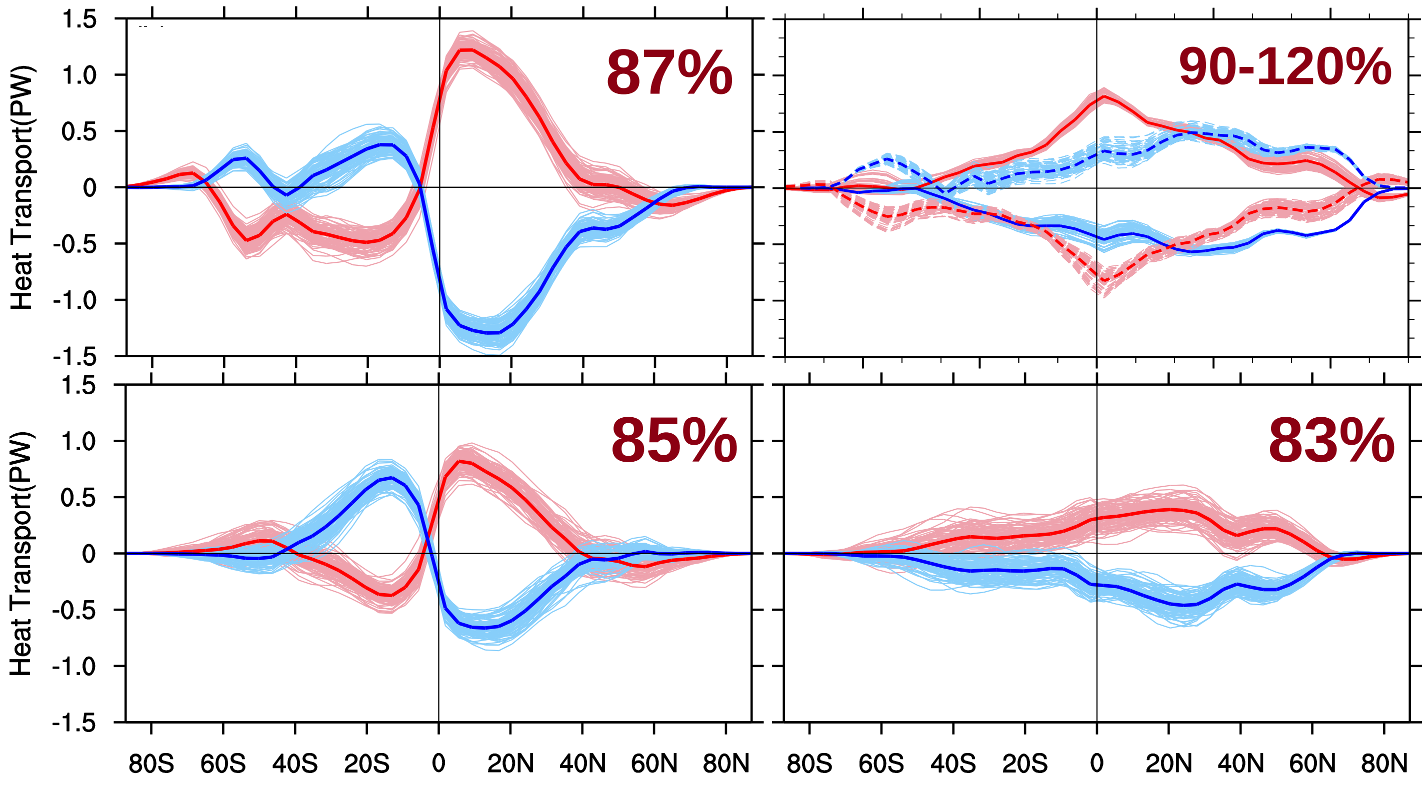
40%

50%

Zhao and Yang (2017)

Summary: BJC in CESM

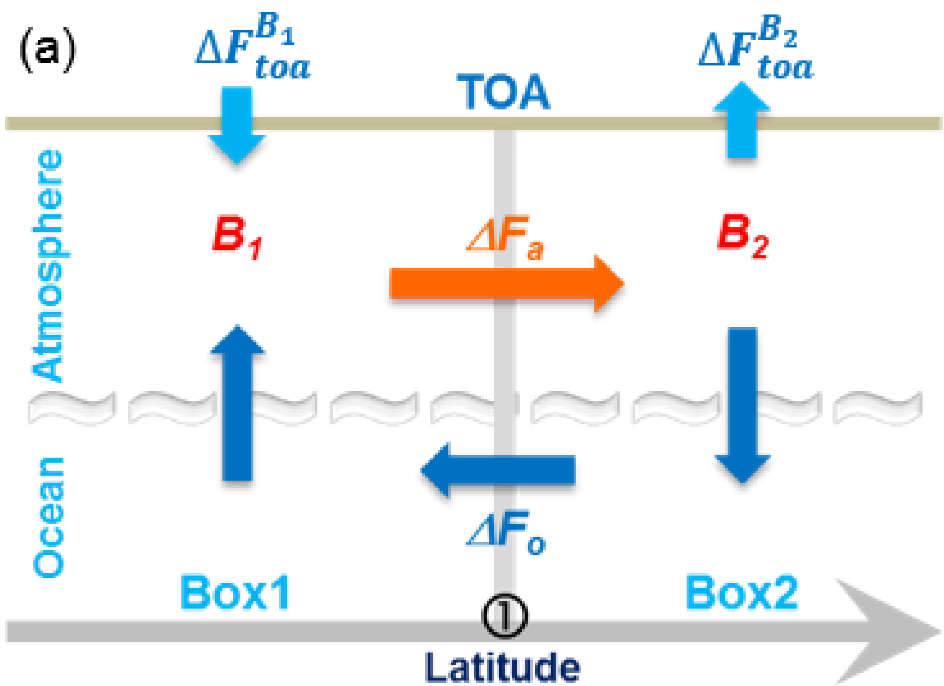
Surface Reason: *Out-of-phase* changes in Atmos-Ocean MOC ➔ BJC



Yang and Dai (2015), Yang et al. (2013, 2016, 2017)

BJC Mechanism

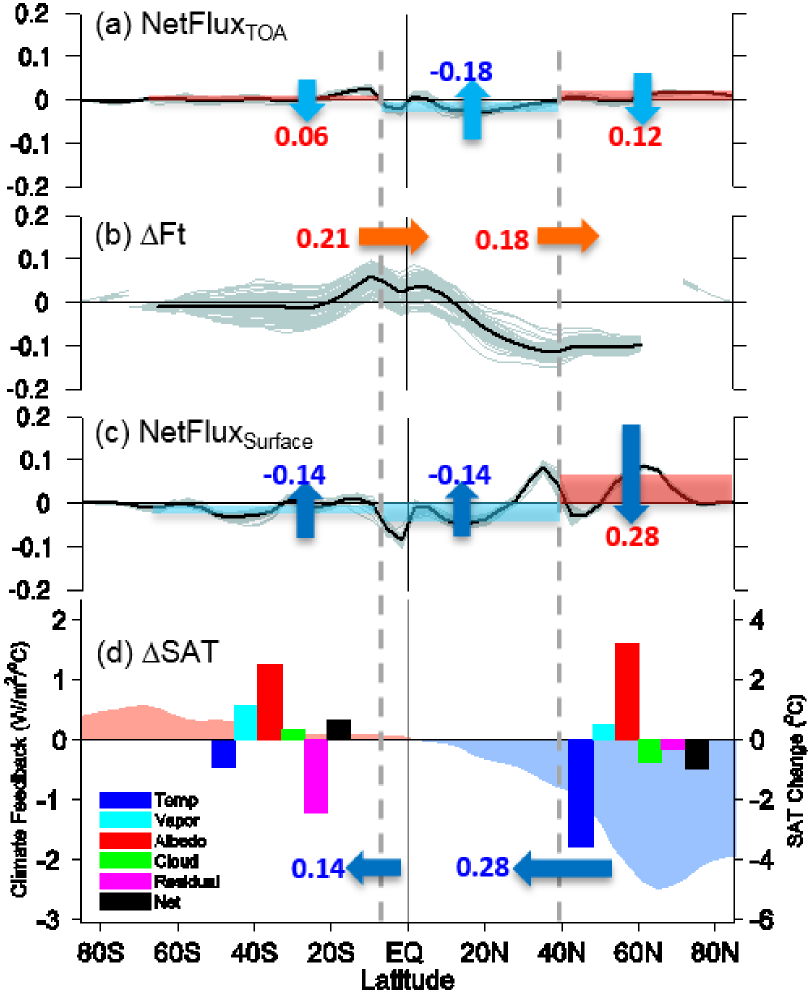
Fundamental Reason: Climate feedback + Energy constrain ➔ BJC



$$C_R = -(B_1+B_2)\chi/[B_1B_2+(B_1+B_2)\chi]$$

BJC: *Theory* vs *CGCM*

CESM Wind Perturbation experiments



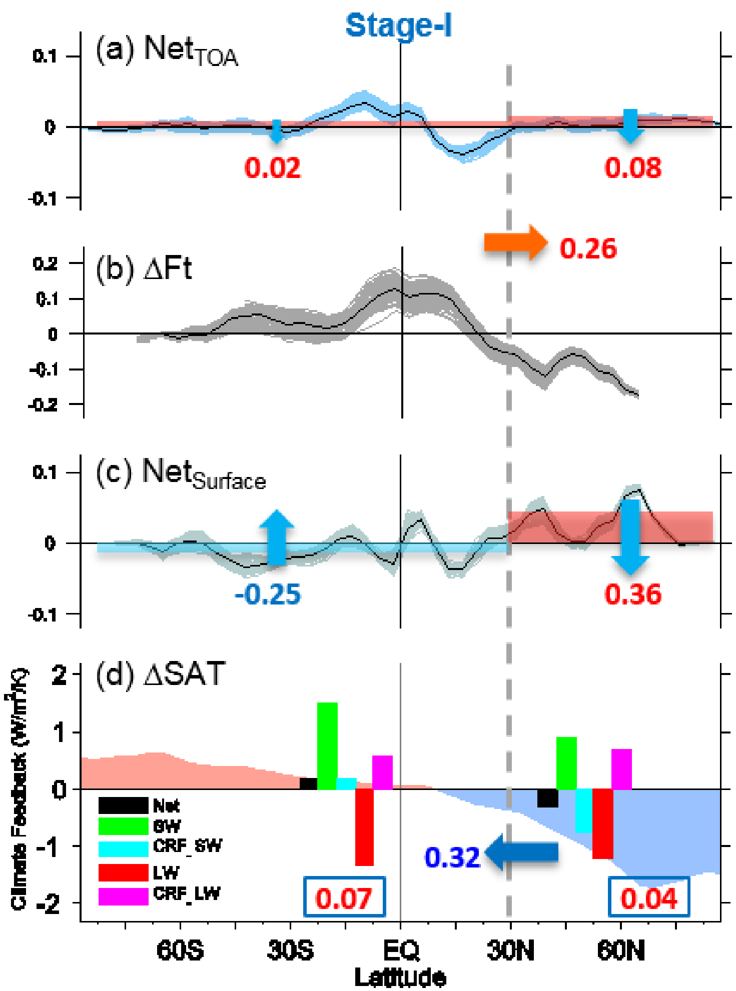
Wind		Temp	Albedo	Cloud	C_{Rmodel}	C_{Rtheor}
0.1A	NH	-1.80	1.30	-0.23	-1.10	-1.09
	SH	-0.44	1.16	-0.07		
	Global	-2.20	1.34	-0.28	-0.83	-0.71
0.1P	NH	-1.87	1.30	0.34	-1.85	-1.61
	SH	-2.14	0.39	0.57		
	Global	-1.88	0.95	0.27	-0.85	-0.77
0.1G	NH	-1.84	1.05	-0.08	-0.95	-0.91
	SH	-1.70	1.96	-0.29		
	Global	-1.77	1.51	-0.18	-0.87	-0.81

$$C_{Rmodel} \approx C_{Rtheory}$$

Dai et al. (2017)

BJC: *Theory* vs *CGCM*

CESM Freshwater experiments



	Freshwater	Global	SH	NH	90°S-30°N	30°N-90°N
Stage-I	Net CF	-0.63	0.89	0.08	0.18	-0.30
	C_{Rmodel}	-0.97	-1.36		-0.88	
	$C_{Rtheory}$	-0.83	-1.05		-1.43	
Stage-II	Net CF	1.96	0.66	0.35	0.16	-0.28
	C_{Rmodel}	-1.11	-1.90		-0.90	
	$C_{Rtheory}$	-2.88	-1.18		-1.33	

$$C_{Rmodel} \approx C_{Rtheory}$$

Yang et al. (2017)

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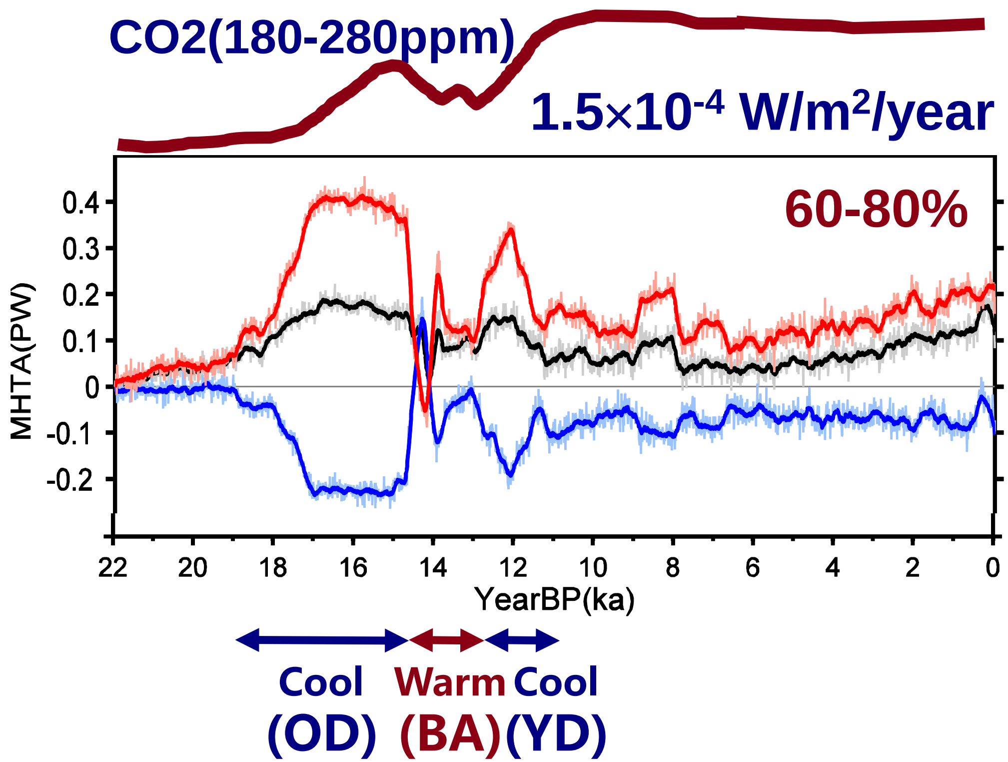
BJC: Coupled *Intrinsic* Mode

3. Relationship between OHT and AHT Changes? **Answered!**

All-in-One **Simulation since LGM**



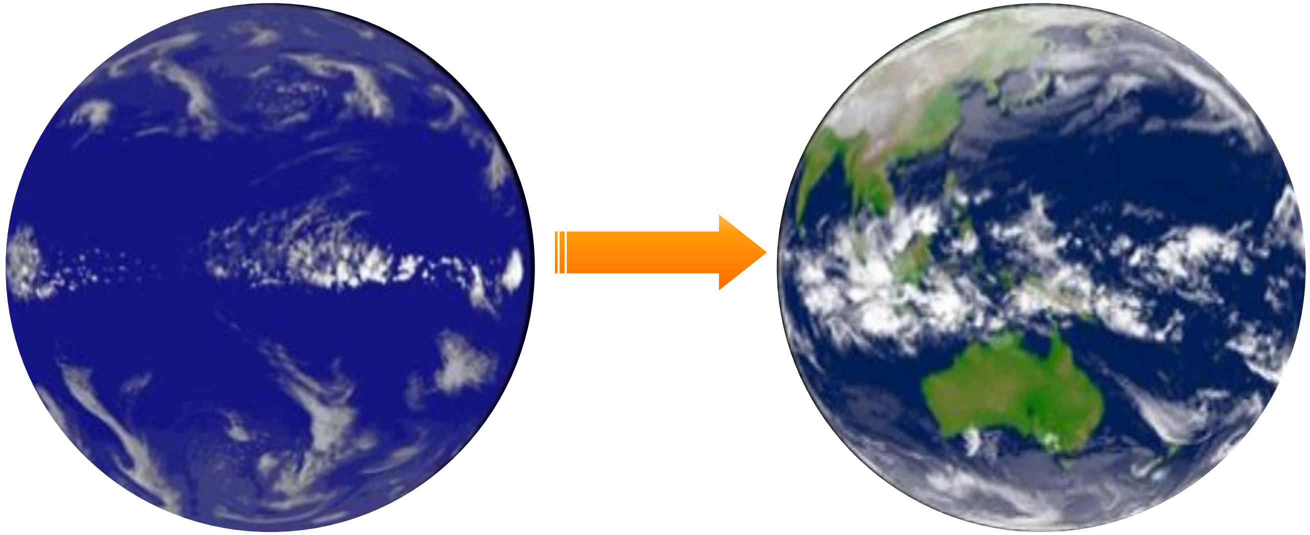
BJC helps to maintain overall Earth climate stability

Yang et al. (2015), Sci. Rep.

Outline

- **Fundamentals**
- **Questions**
- **Hypothesis and Theory**
- **CGCM results**
- **Aquaplanet**
- **Summary**

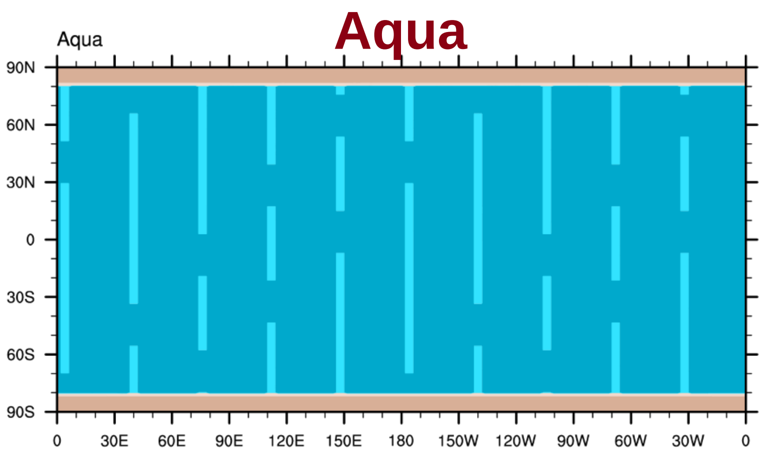
Why anti-symmetric MHT?

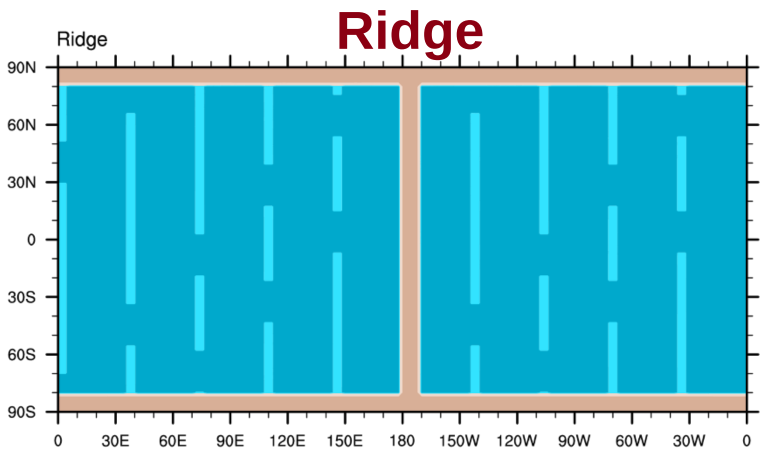


Try to answer this fundamental question

Topo for Aqua and Ridge

Land: 10 m; Ocean: 5000 m; Bottom random ridge: 500 m





Li and Yang (2018)

Go to Summary

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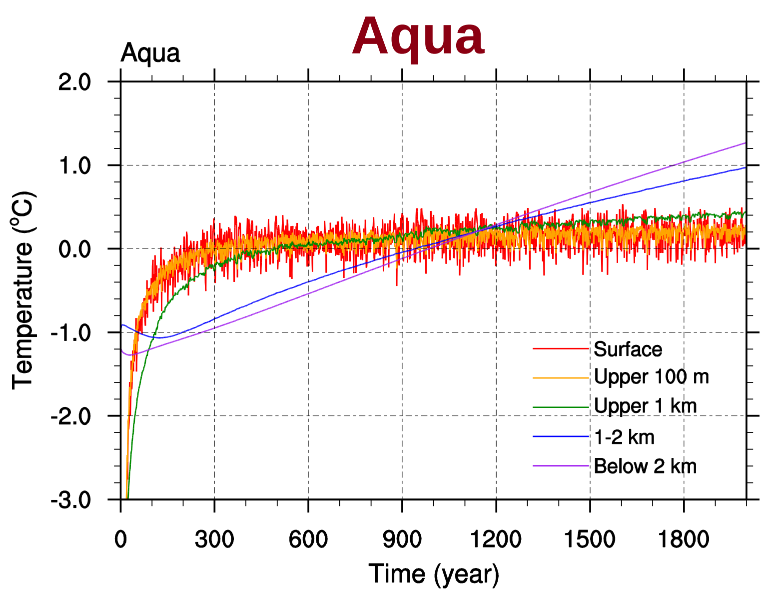
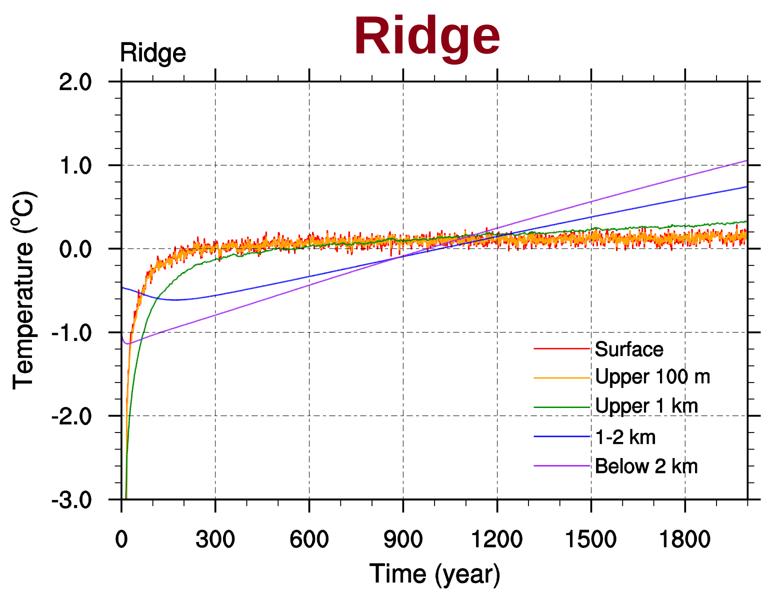
Hard to Reach Equilibrium

2000-year Simulation using CESM





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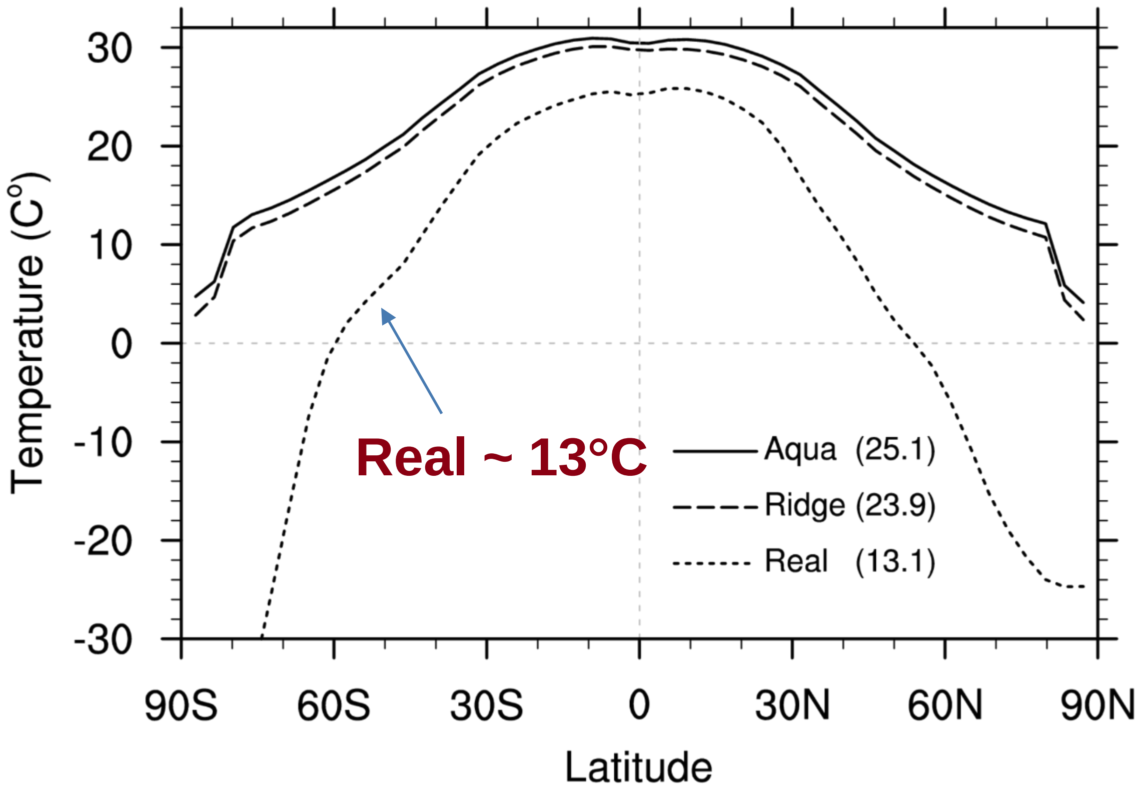
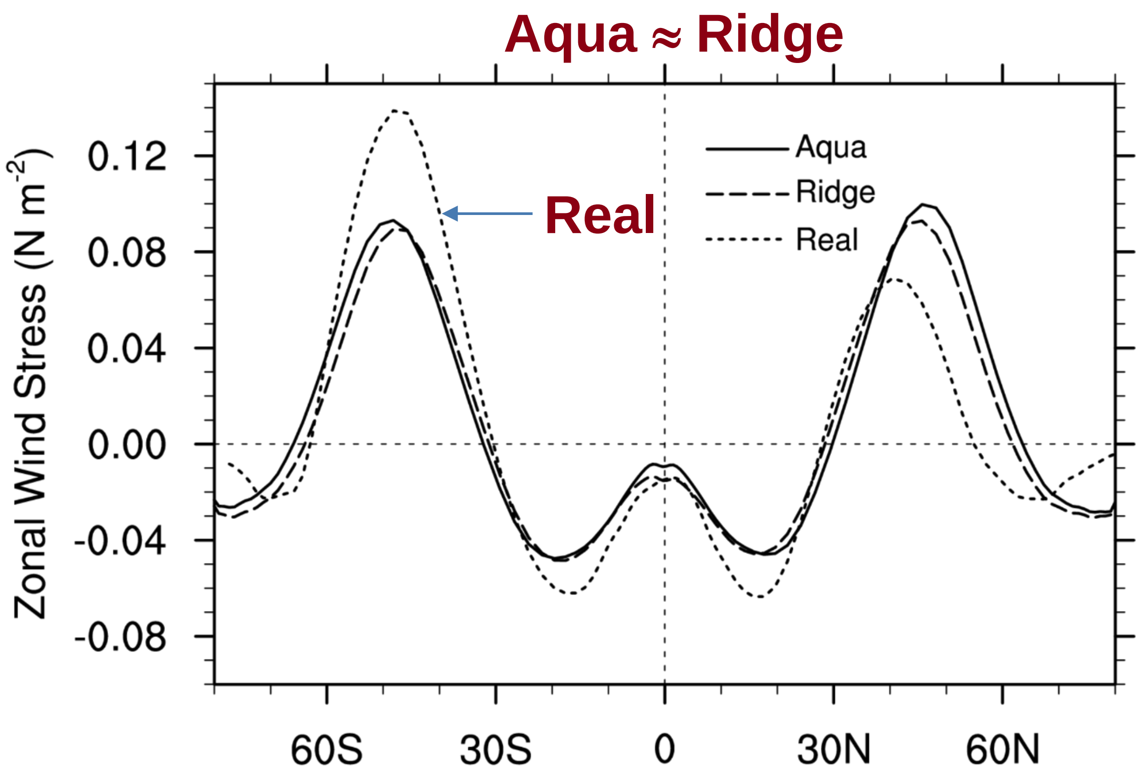
SAT: Warmer Climate

Aqua ≈ Ridge ≈ 24 ~ 25°C



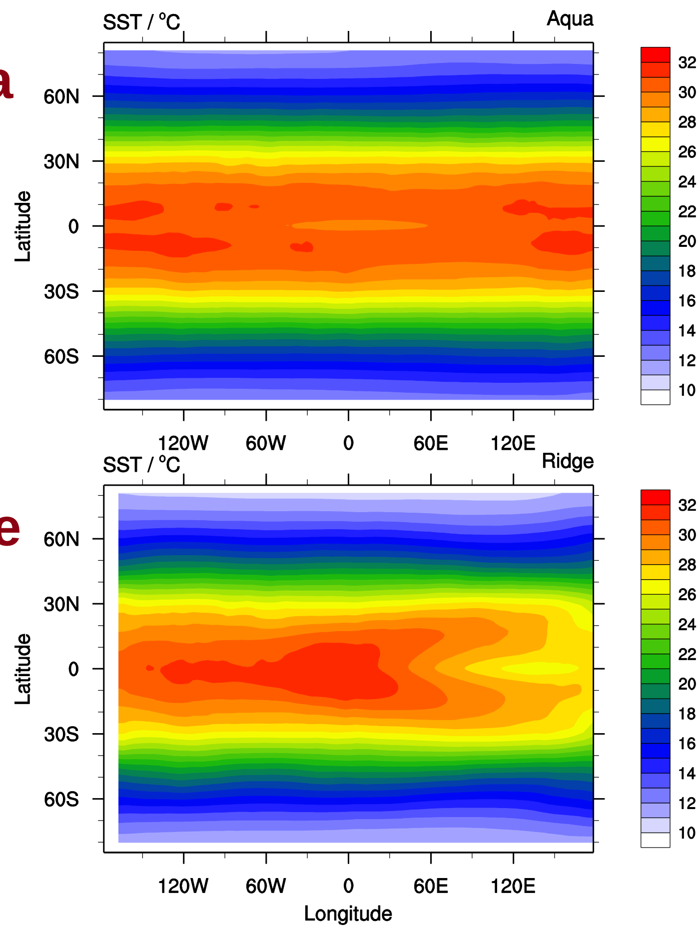
Symmetric Zonal Surface Wind Stress

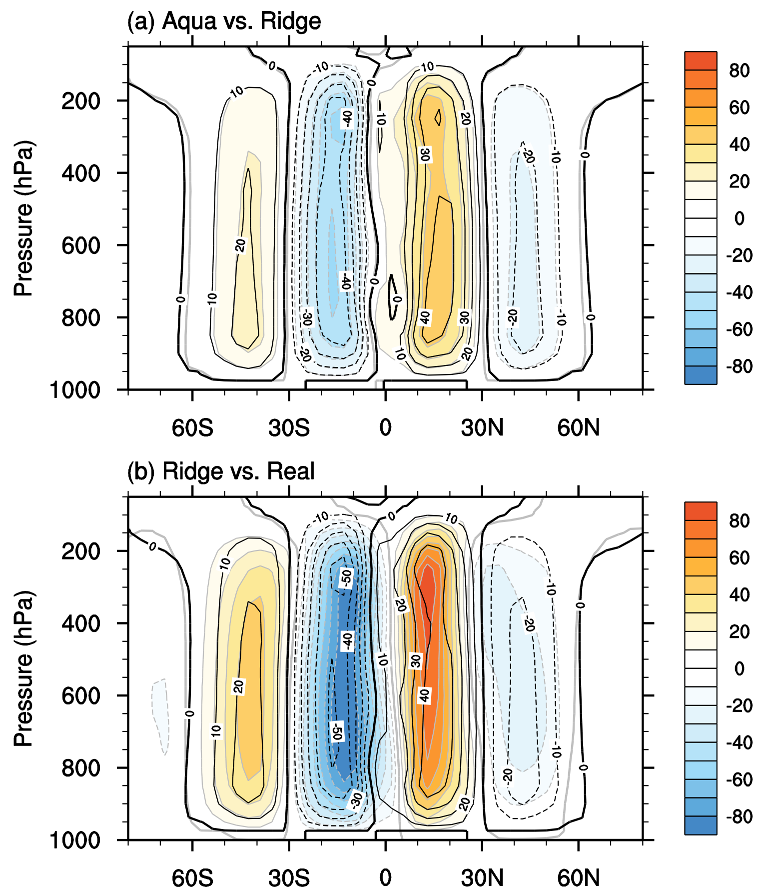


Symmetric SST and Hadley Cell

Aqua

Ridge



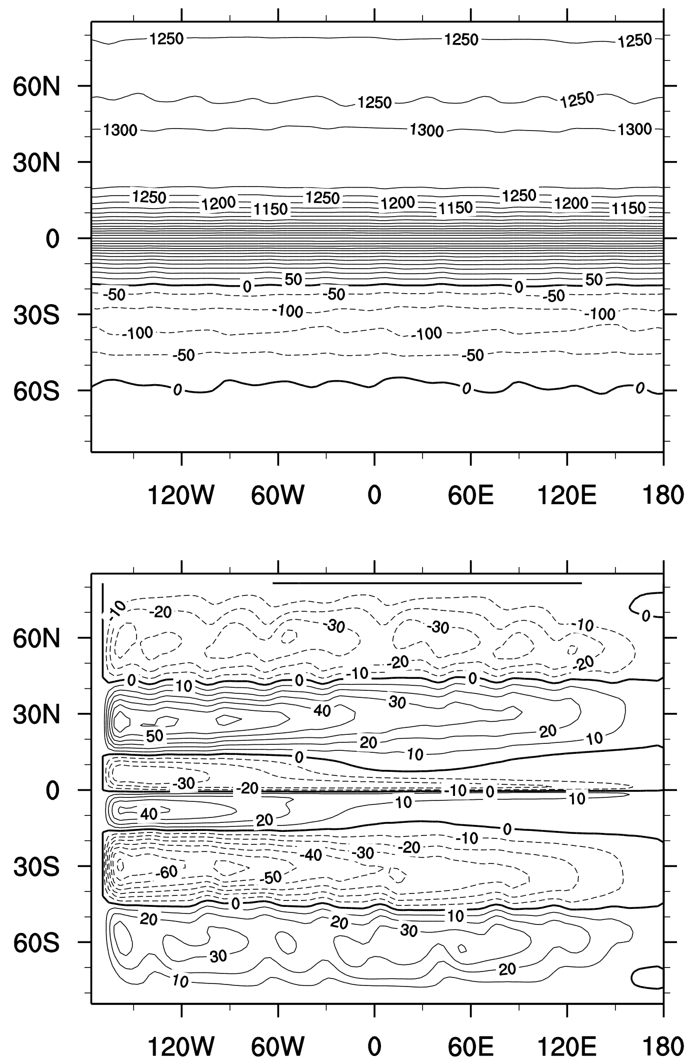


Symmetric Ocean Circulation

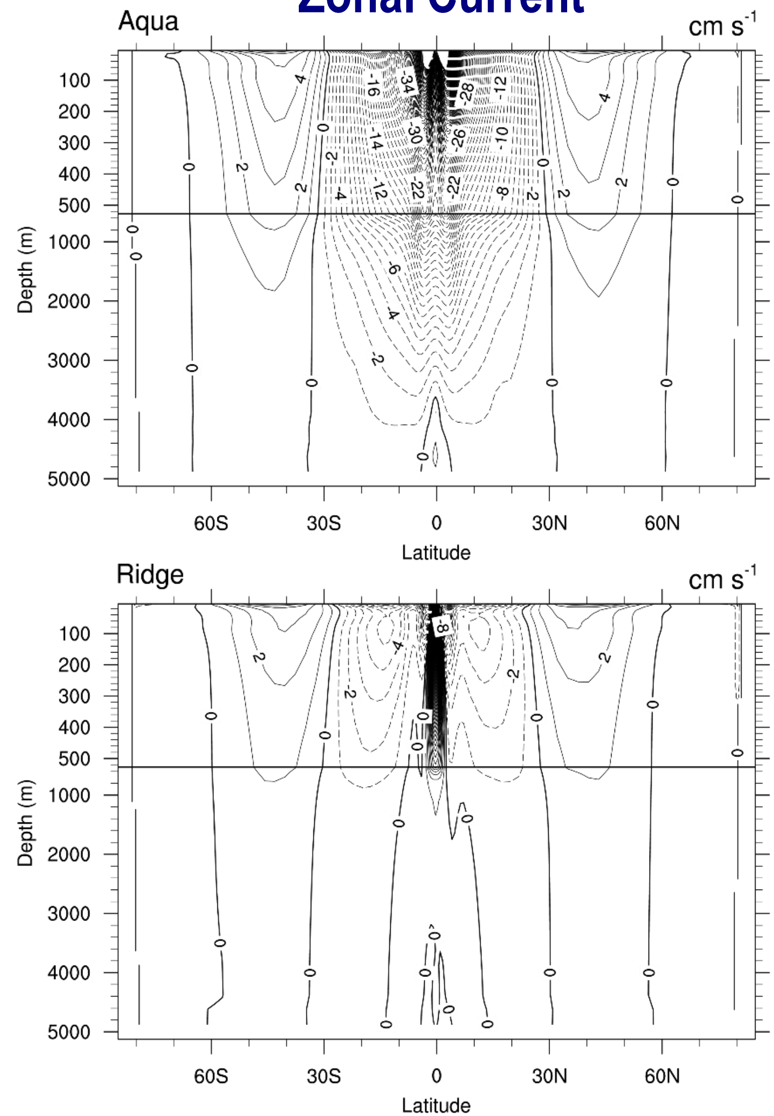
Barotropic Streamfunction

Aqua

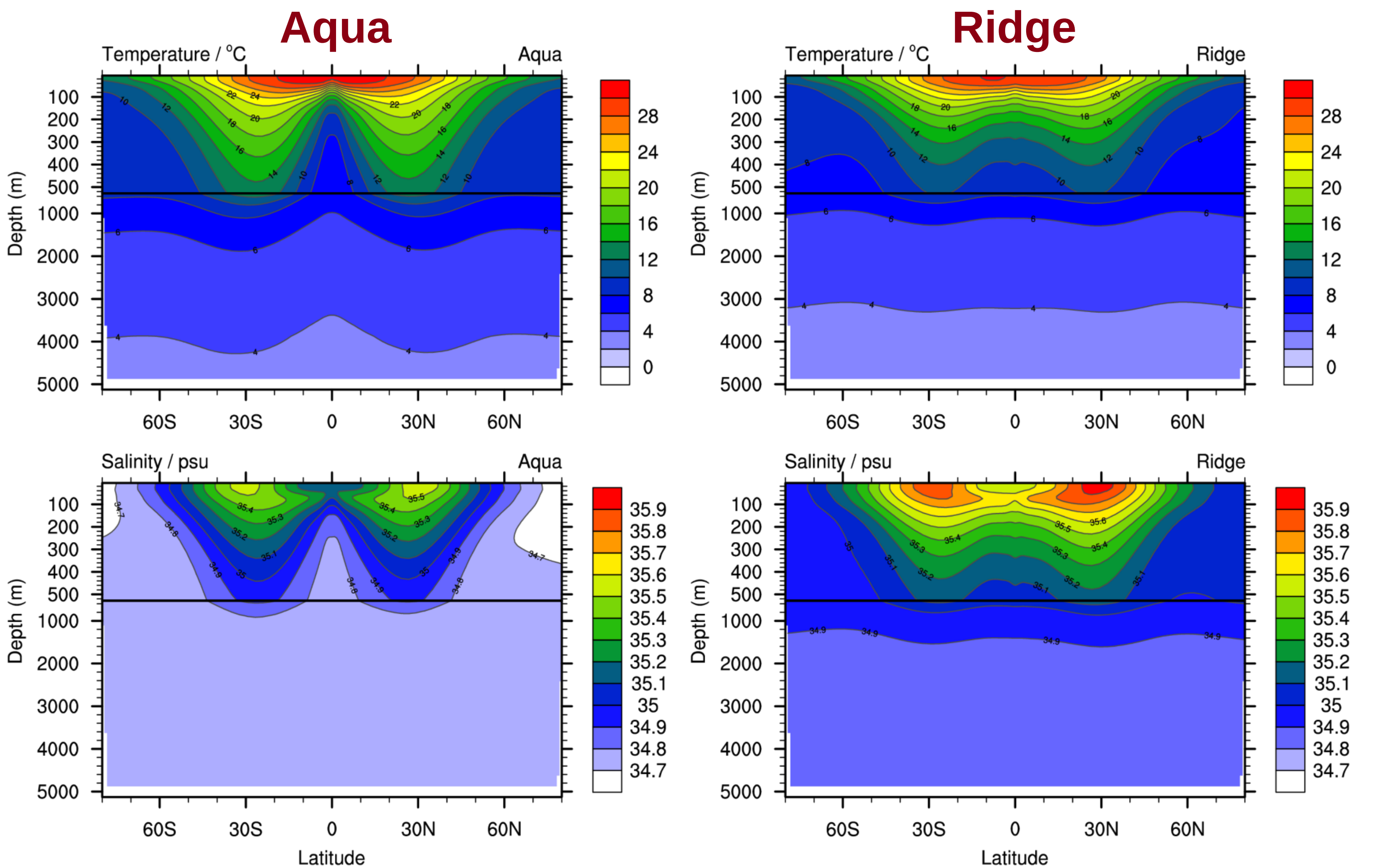
Ridge



Zonal Current

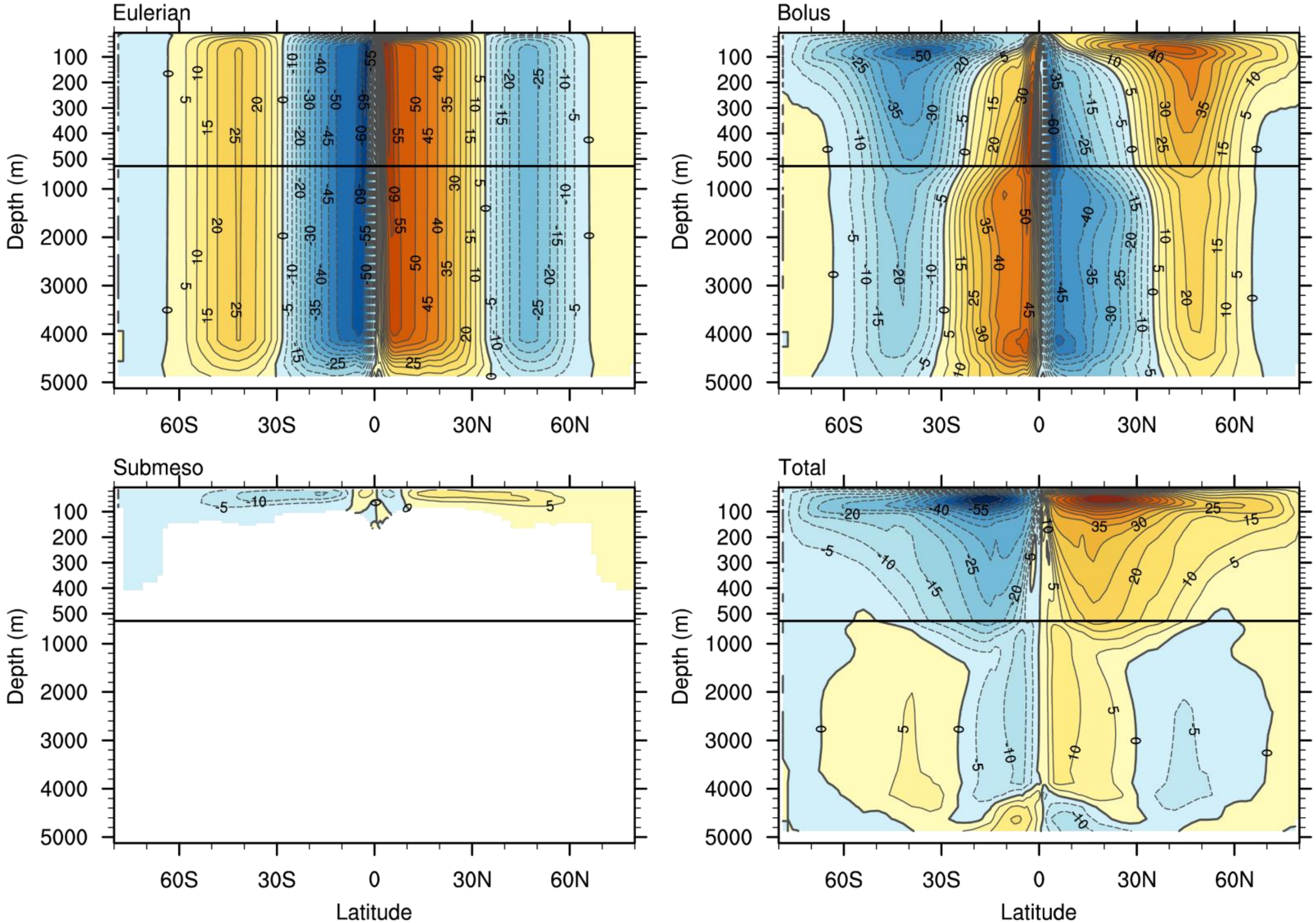


Symmetric Ocean Buoyancy



Symmetric Overturning Circulation

Aqua



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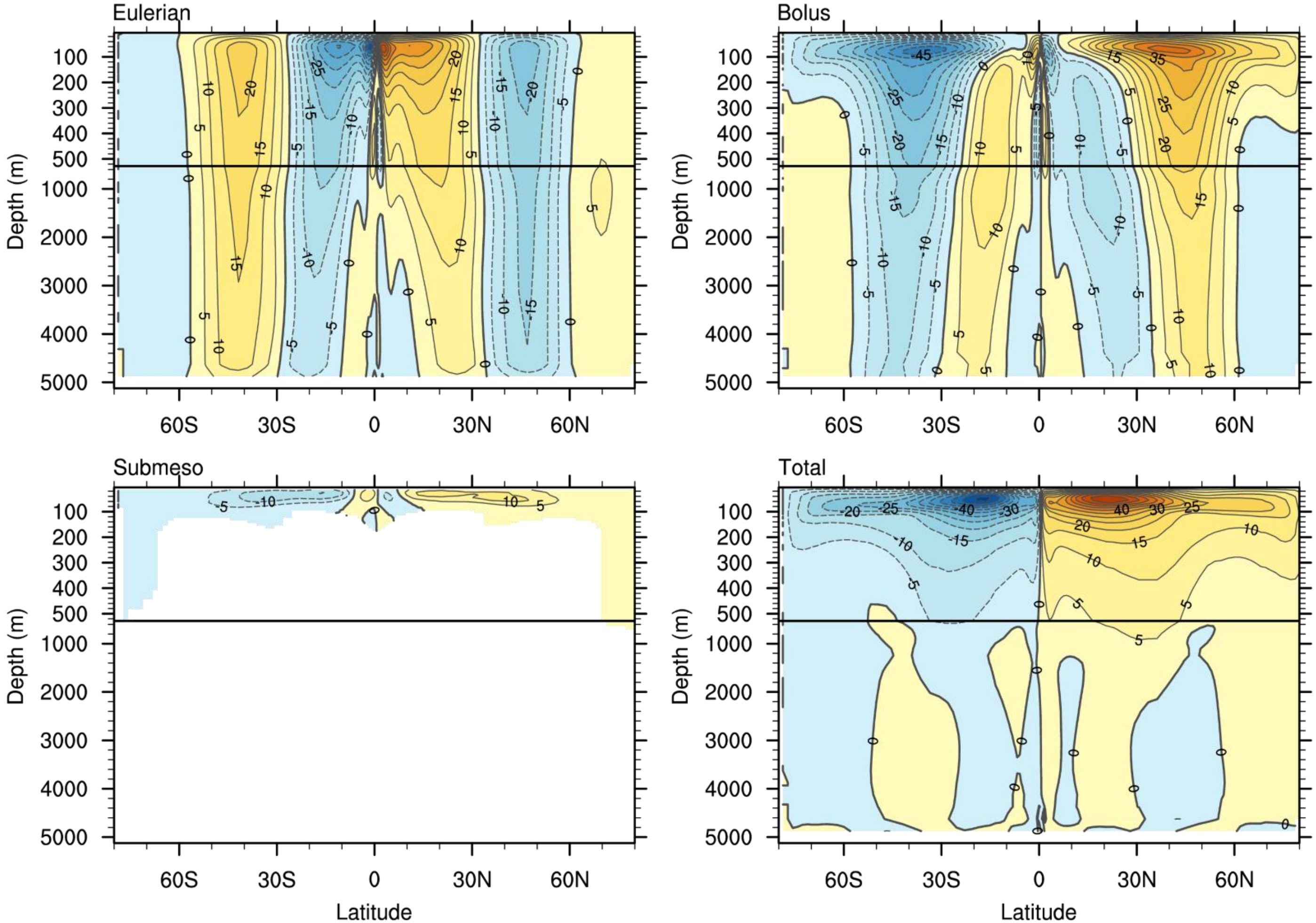
Symmetric Overturning Circulation

Ridge



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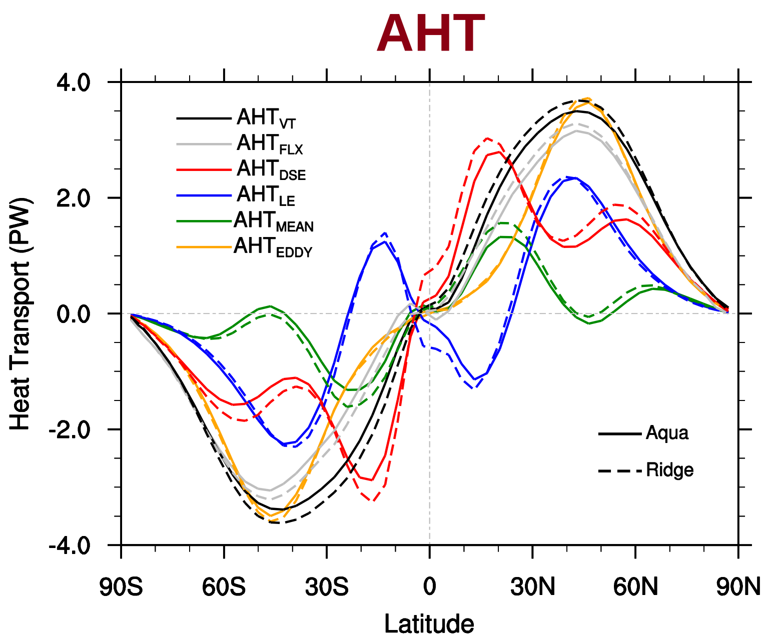
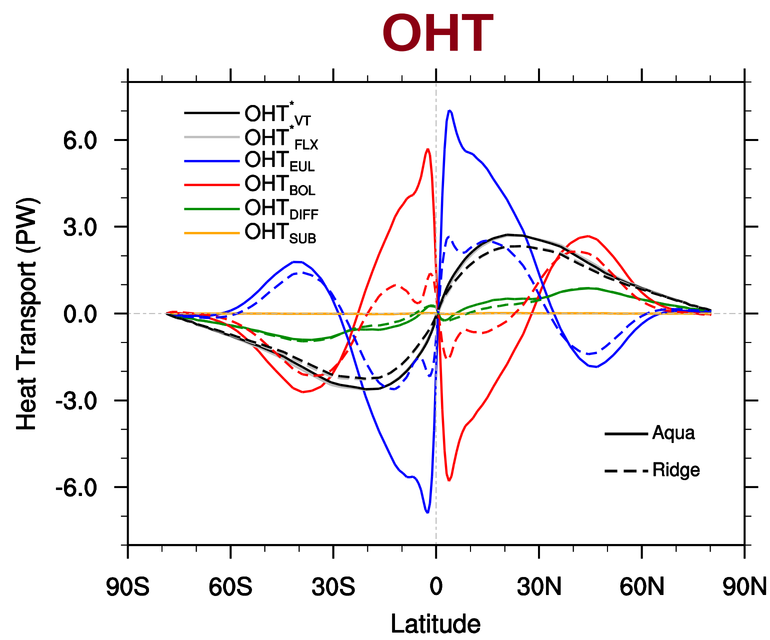
Symmetric AHT and OHT





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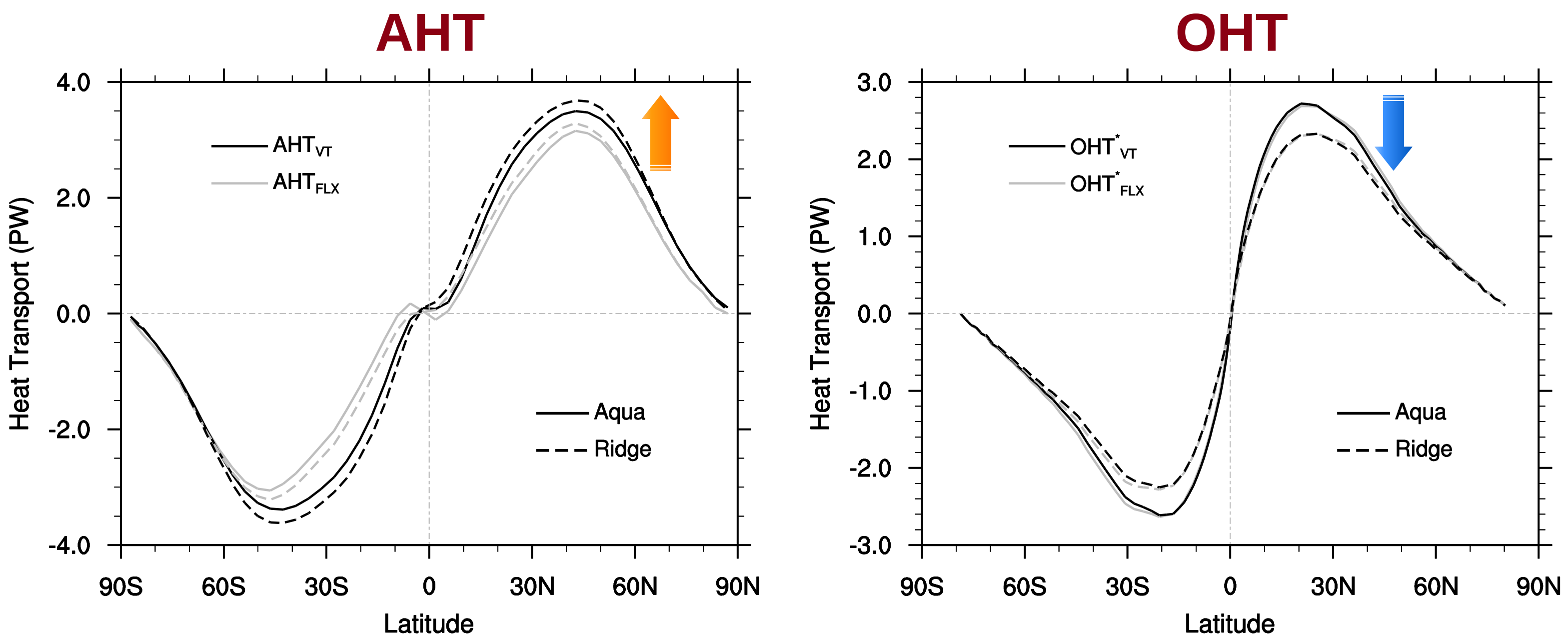
Symmetric AHT and OHT

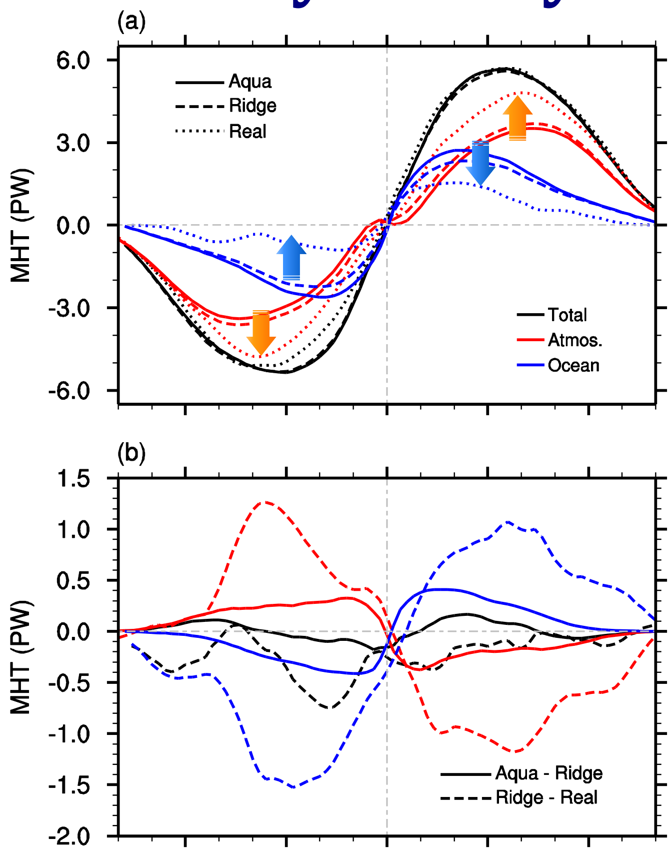


From *Aqua* to *Ridge* ➔ OHT ⇓ and AHT ⇑

Bjerknes Compensation

Aquaplanet → Real Earth

Why anti-symmetric MHT? Answered



From *Aqua* to *Real*

1. OHT ⇓ and AHT ⇑
2. OHT → asymmetric, NH>SH, AMOC + Weaker baroclinic
3. AHT → asymmetric, SH>NH, Stronger baroclinic dT/dy ⇑

→ **Total MHT *unchanged***

BJC maintains antisymmetric MHT!

Summary and Discussion

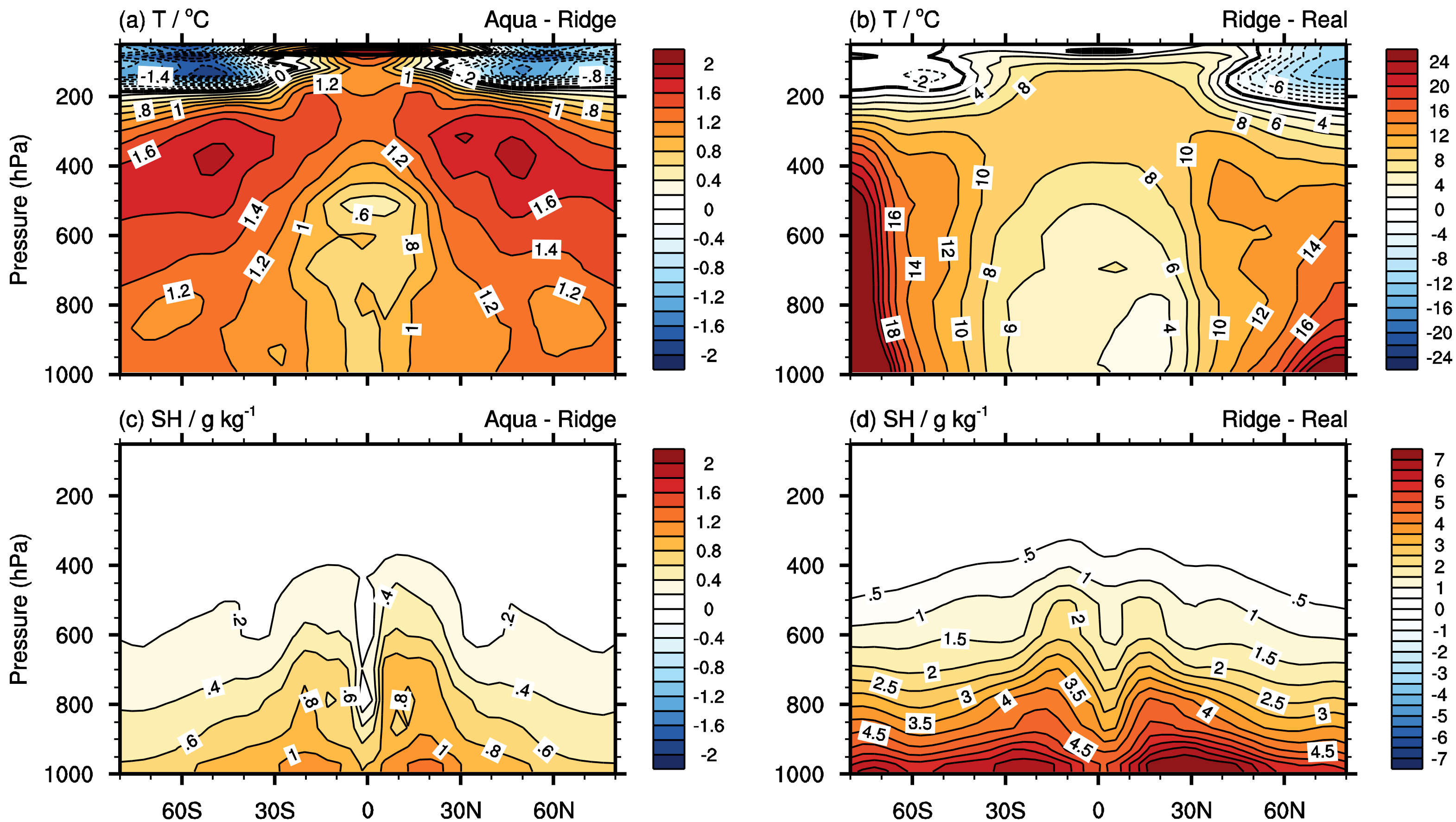
- Bjerknes compensation
  - Intrinsic mode
  - Atmospheric physics ⇔ Physical oceanography
    - Climate feedback ⇔ Thermohaline circulation
- Self-constraint mechanism
  - Climate didn't drift too much
- If feedback ➔ Reversibility of climate
  - Invisible hand (?)

LaCOAS
北京大学气候与海-气实验室

谢 谢

Aquaplanet → Real Earth

Weaker Baroclinic and More Moisture



Coupled Intrinsic Mode

Energy Compensation ⇔ Weight Management

Diet
Ocean

Motion
Atmos

Metabolism
Climate
Feedback

$$C_R = -\frac{1}{1-B}$$

Go to BJC for Climate Variability