

连接青藏高原与南北极：季节与路径
**Connecting Tibetan Plateau to Arctic and Antarctic
Seasonality and Passage**

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Tibetan Plateau (TP): the 3rd Pole

Total Area: 2.5 million km², Elevation: 4000 m



A Fundamental Question: *With / Without* TP



- Greenland Ice Melting → **+7m** ↑ global ocean



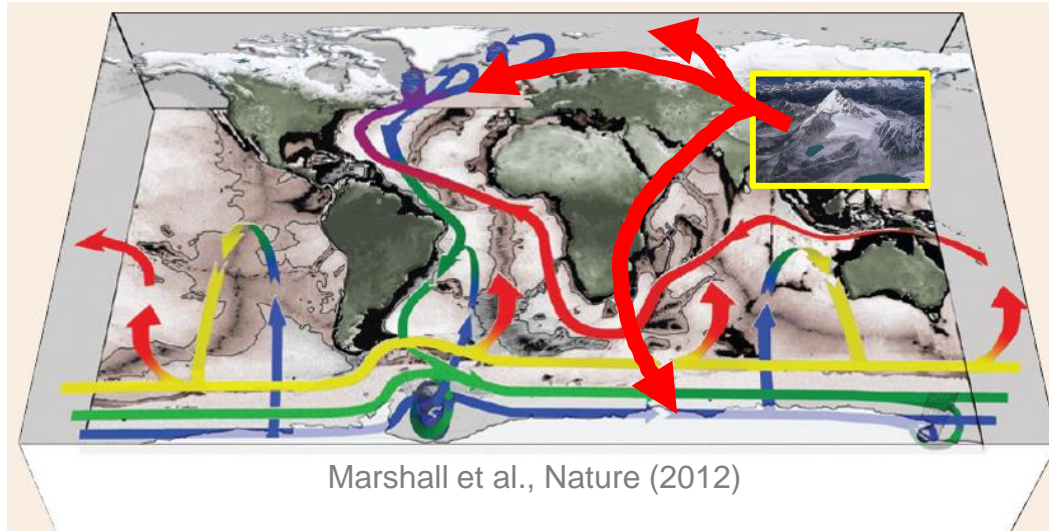
- Antarctic Ice: 70% FW, 90% Ice Melting → **+61m** ↑ global ocean



- ***With / Without*** TP: Sea level and fundamental climate differences?
25 billion m³ freshwater around TP

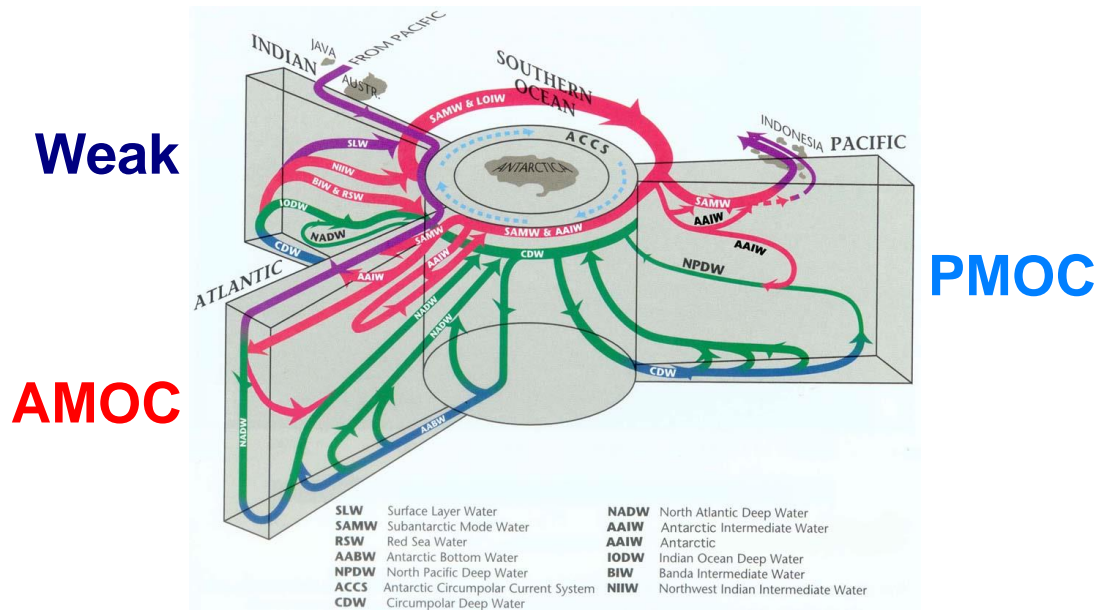
TP: A *Global* Perspective

How and to what extent?



Global Meridional Overturning Circulation

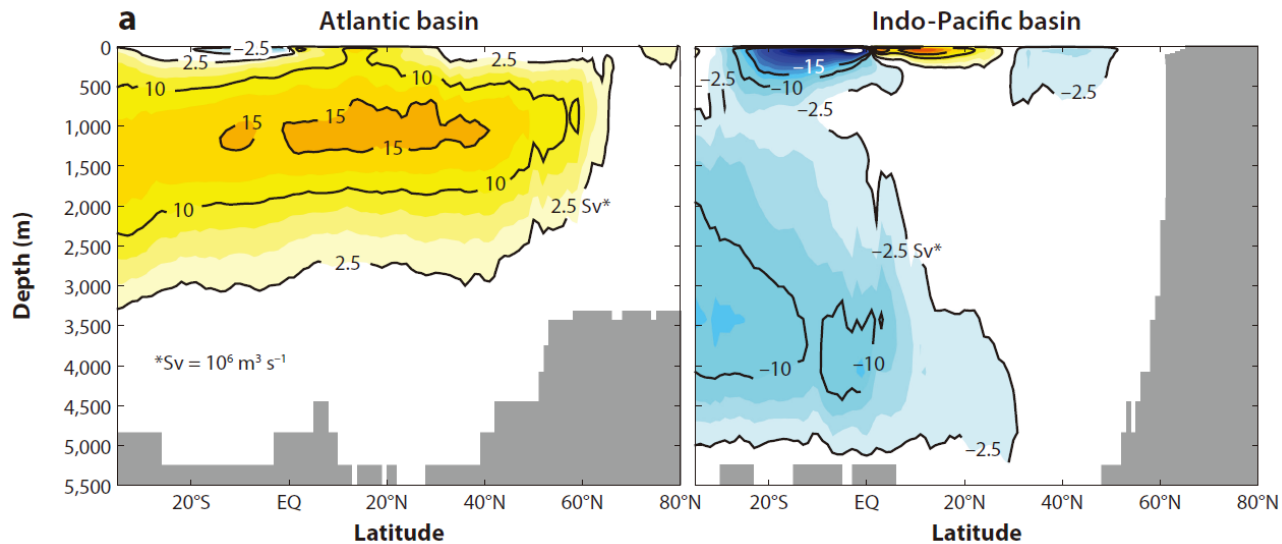
Energy and Freshwater Balance



Schmitz (1997) Overturning circulation: Southern Ocean View

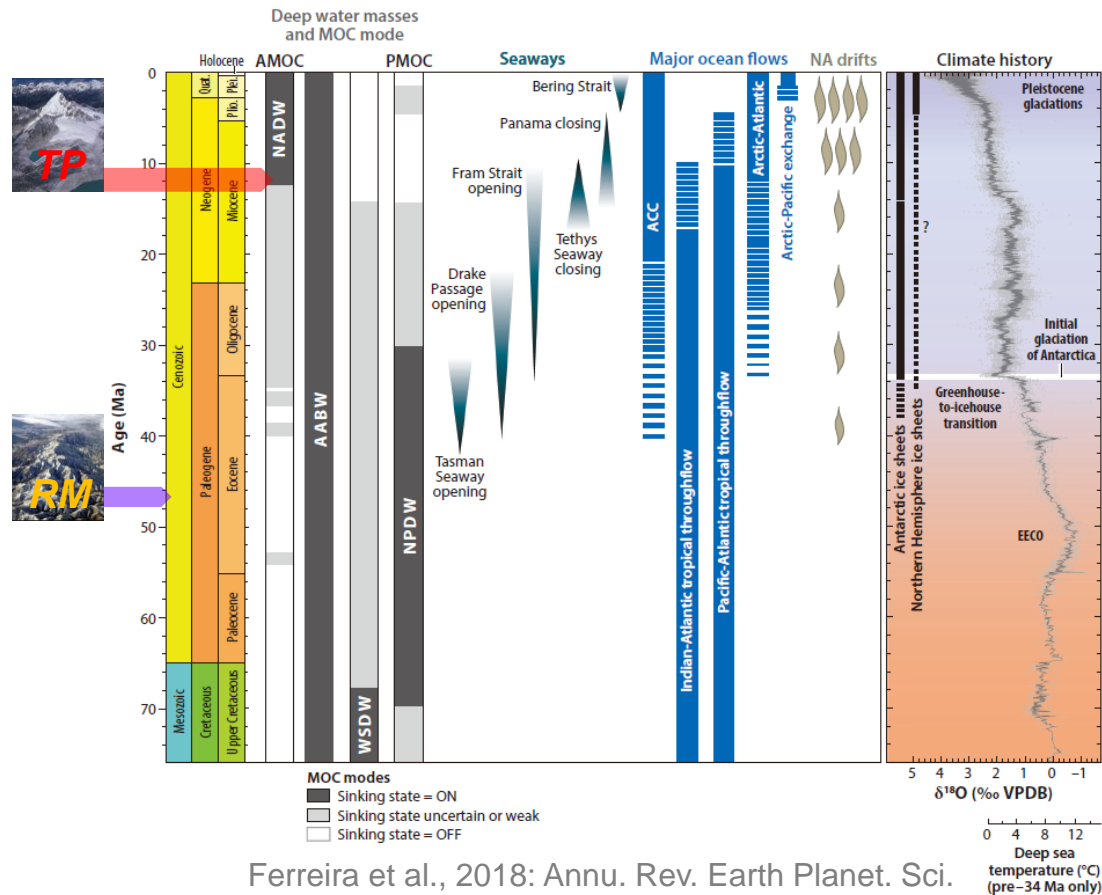
Strong AMOC

Weak PMOC



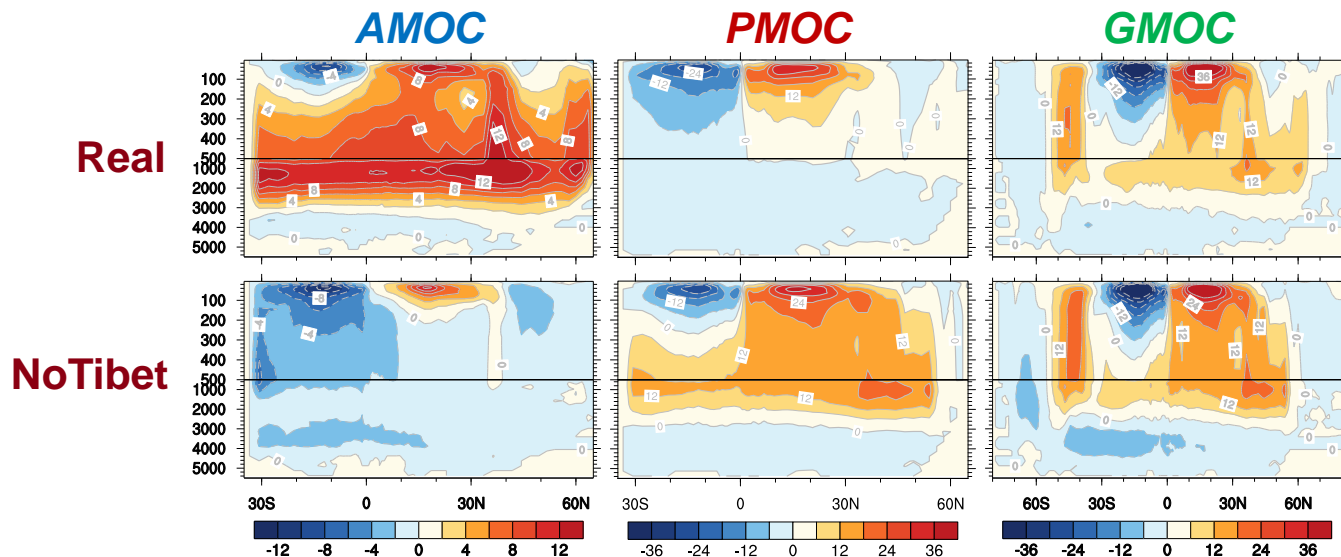
Ferreira et al., 2018: Annu. Rev. Earth Planet. Sci.

Geological History of *GMOC*

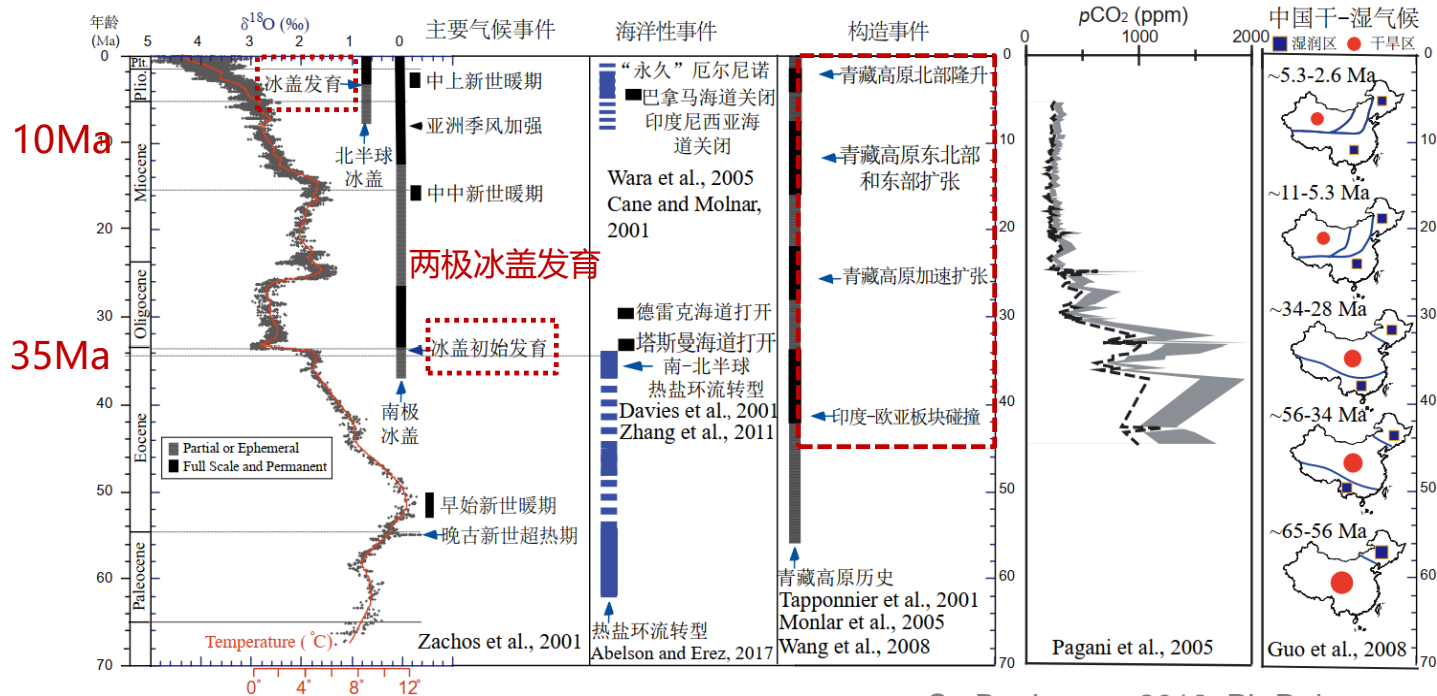


Ferreira et al., 2018: Annu. Rev. Earth Planet. Sci.

Global MOC



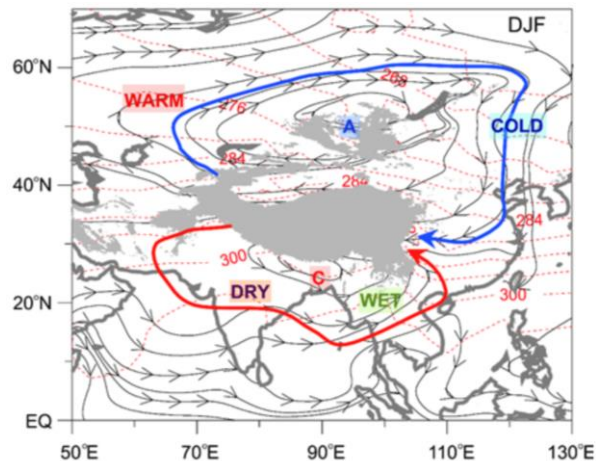
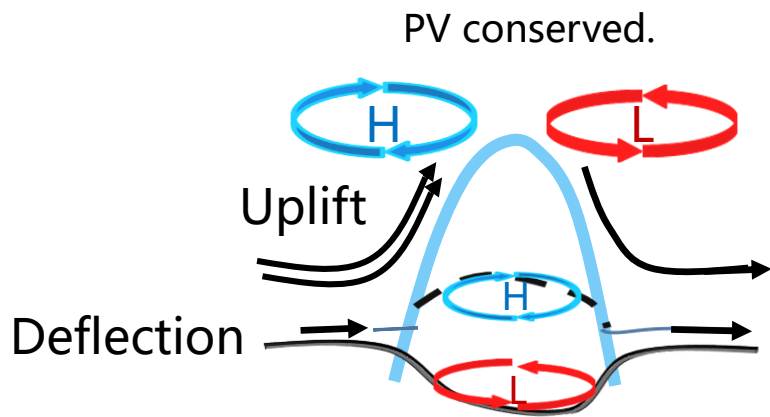
Geological History of Significant Events



Su Baohuang, 2018: Ph.D theses

- 新生代以来，全球气候发生了巨大的变化，山脉隆升被认为是长时间尺度气候变化的重要驱动因素
- 青藏高原的隆升对邻近区域气候的研究非常广泛，较少关注遥远的北极和南极

Dynamical Effect

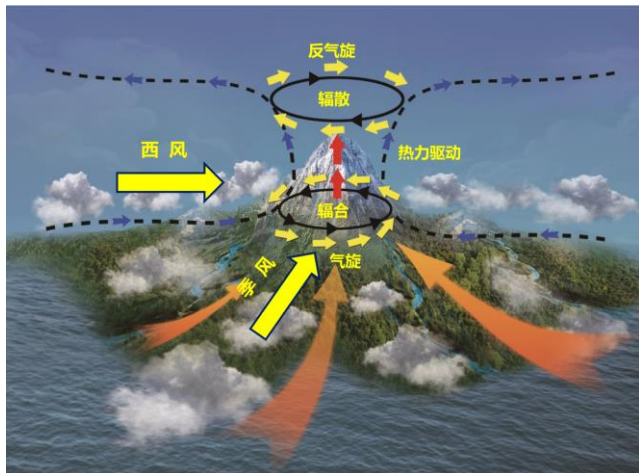


Wu et al. (2015)

低层绕流：
冷暖平流
干湿平流

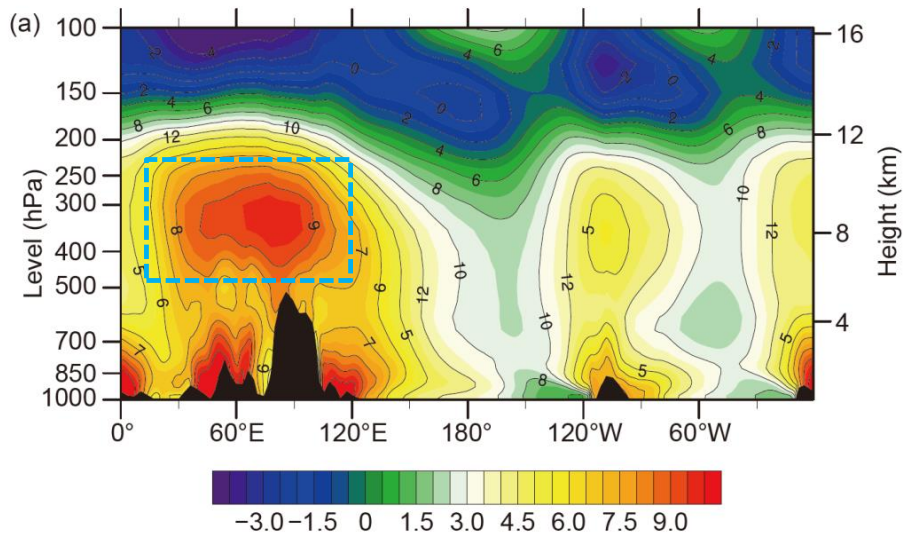
- 低层绕流：“北脊南槽”；中高层爬坡：迎风坡高压，背风坡低压
- 冬季：西风南移加强，TP位于西风带内；夏季：西风带北移，动力作用减弱
- 经过高原的背景气流有季节变化，高原对环流的影响有显著季节性差异

夏季热力作用示意图



- 夏季热源，冬季冷源
- 低层热低压，高层南亚高压

夏季27.5°~35°N东-西向气温纬向偏差垂直剖面图

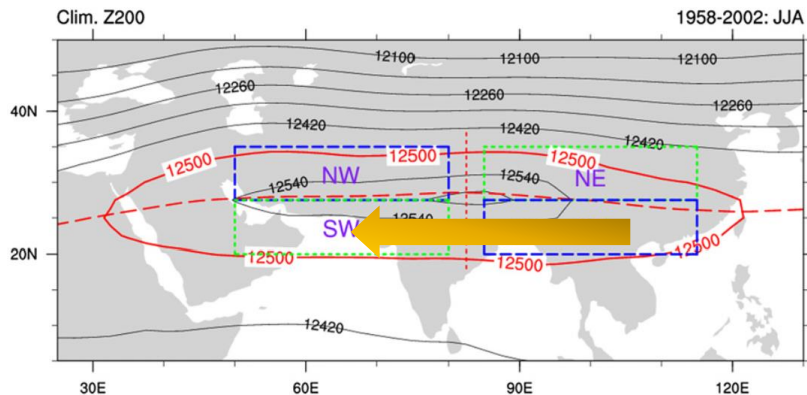


- 夏季热源可加热到高层自由大气

Xu et al. (2019)

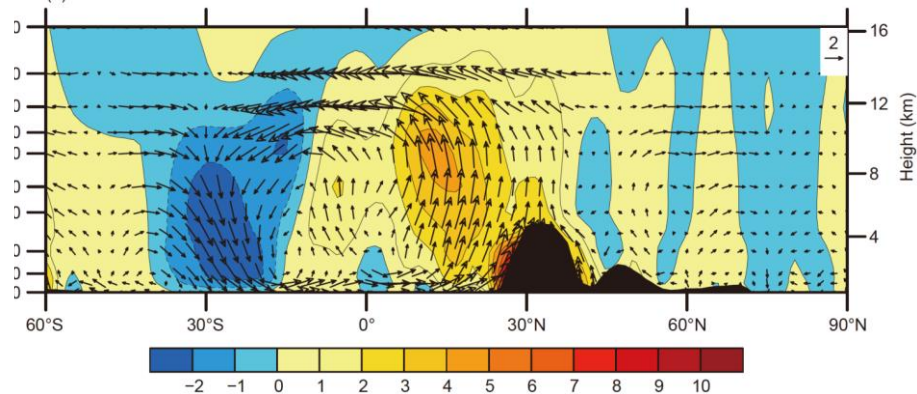
Thermal Effect of TP

1958-2022年观测的200hPa位势高度



- 高层南亚高压，热带东风急流

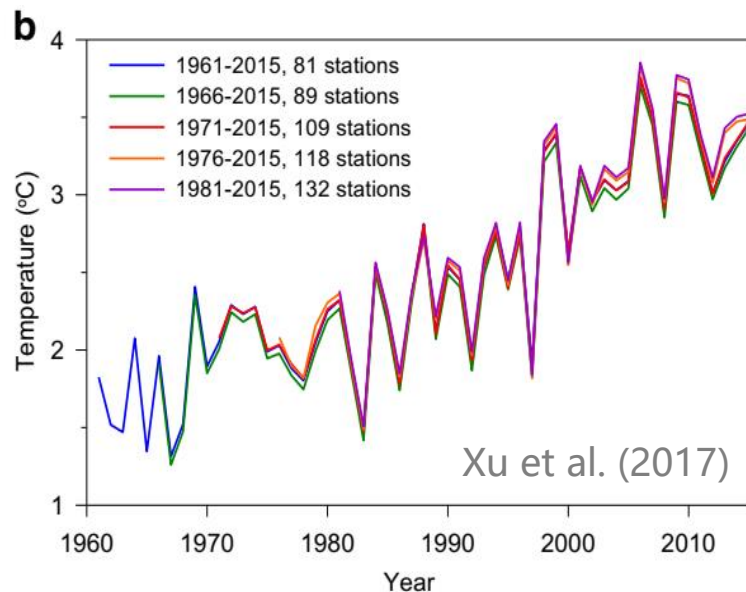
1948~2018年夏季80°~110°E平均经向风速 (vector, m/s)和垂直速度 (color)



- 经向反哈德莱环流，捕获水汽

Xu et al. (2019)

TP: Paleoclimatic and Modern Significance



0.35°C/decade in the recent 55 years

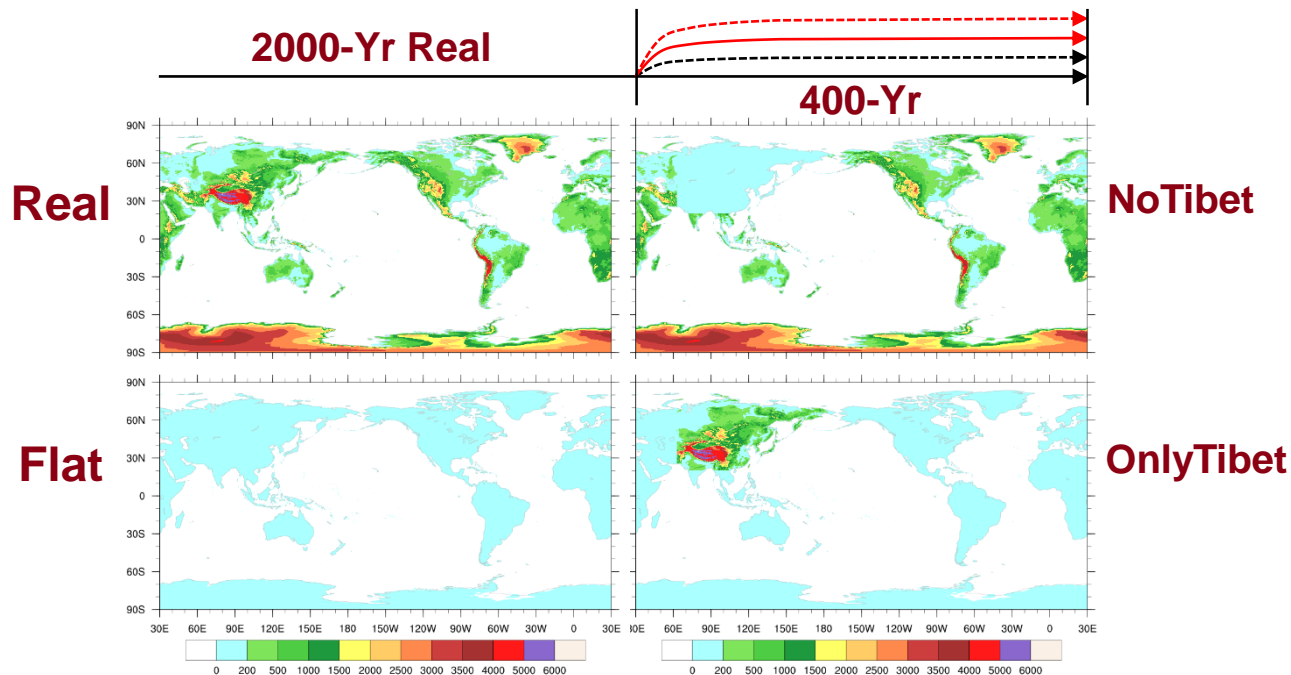
- 青藏高原区域对全球变暖非常敏感，远大于全球平均表面增温率，对全球气候产生强反馈
- 大气过程迅速持久，未来气候变化下，青藏高原的异常如何影响两极地区

Fascinating Question: NP ⇔ TP ⇔ SP



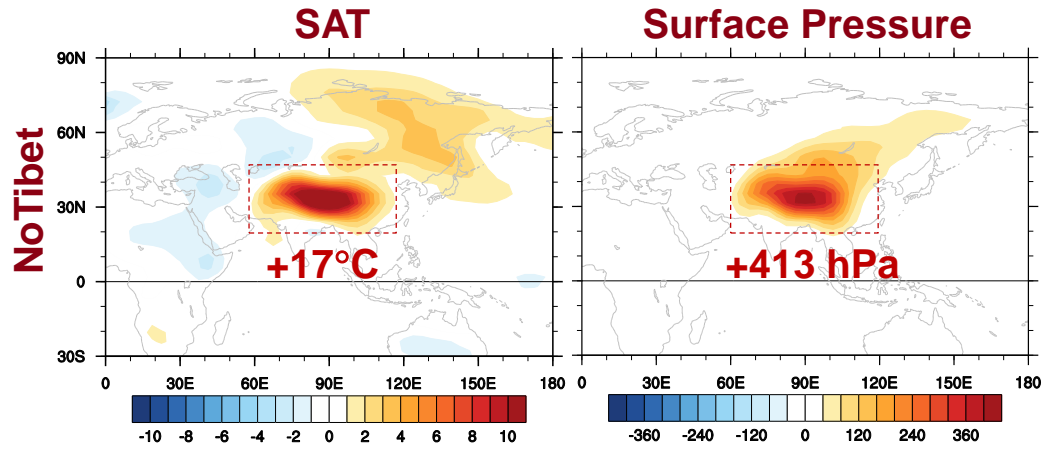
Seasonality, Passage and Mechanism ?

Coupled Model Experiments *With / Without* TP



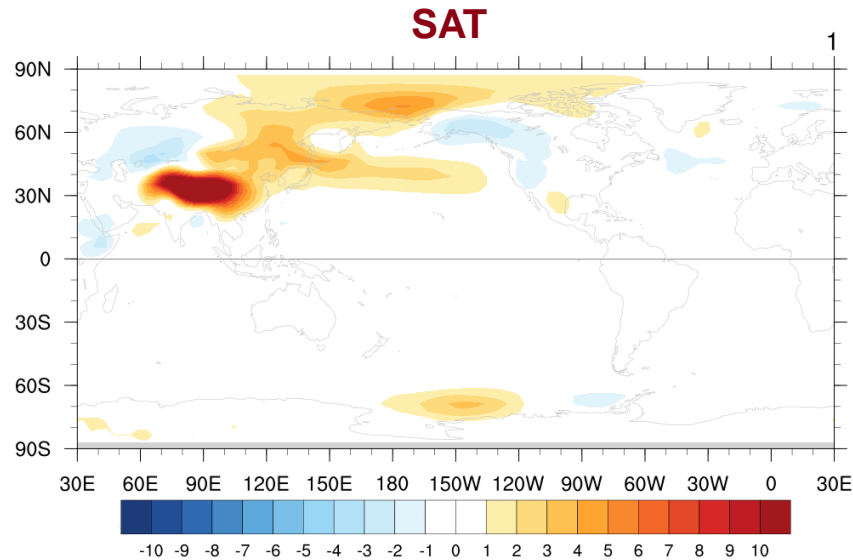
NCAR CESM1.0: CAM5 / POP2 / CLM4 / CICE4 / Glimmer-CISM

TP Forcing: *Thermal* and *Dynamical*

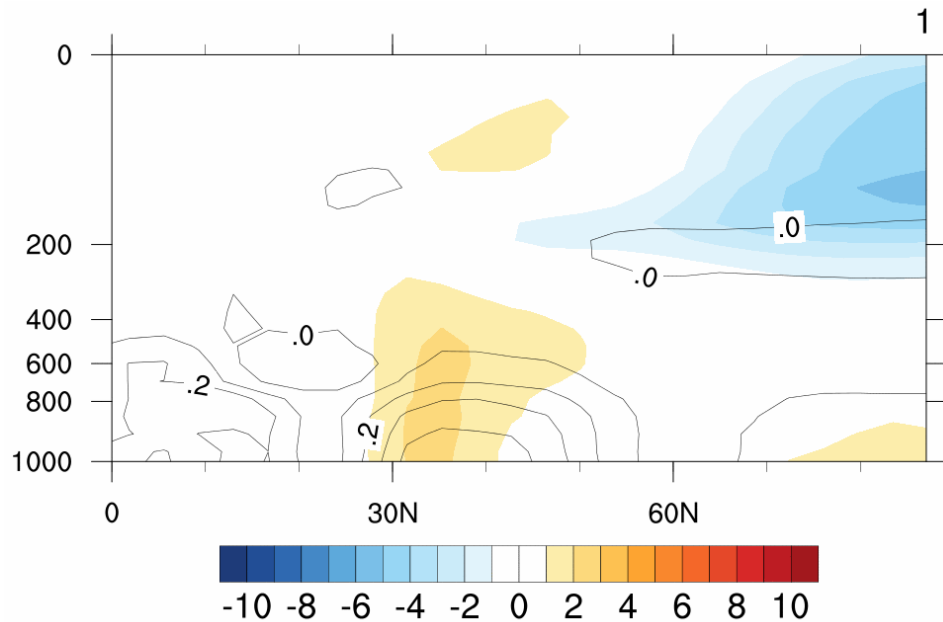


Lapse Rate $T \sim 4 \text{ km} \times 7 \sim 28^\circ\text{C}$

SAT Evolution w/o TP

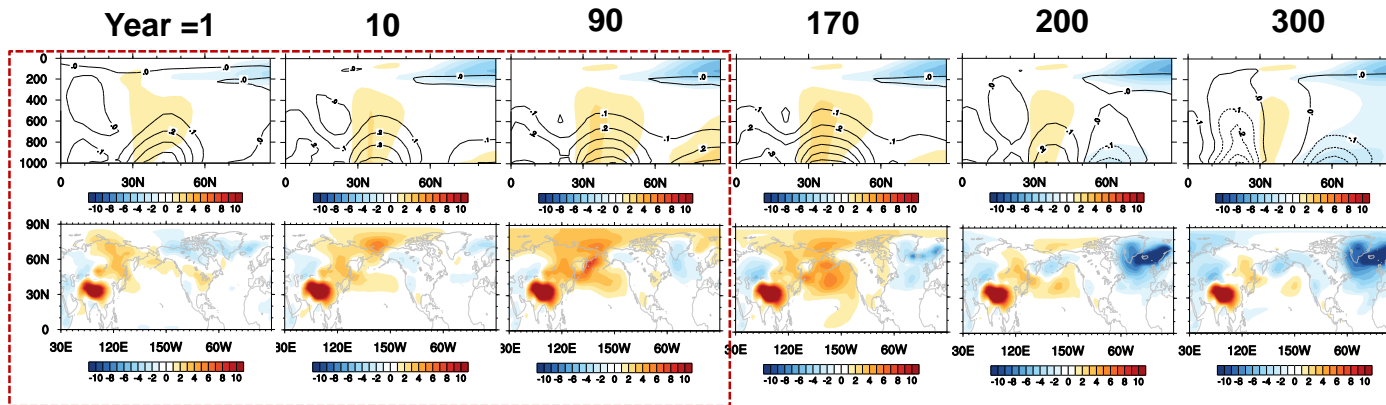


Atmosphere T and Moisture in NH



Transient Atmosphere T and Moisture in NH

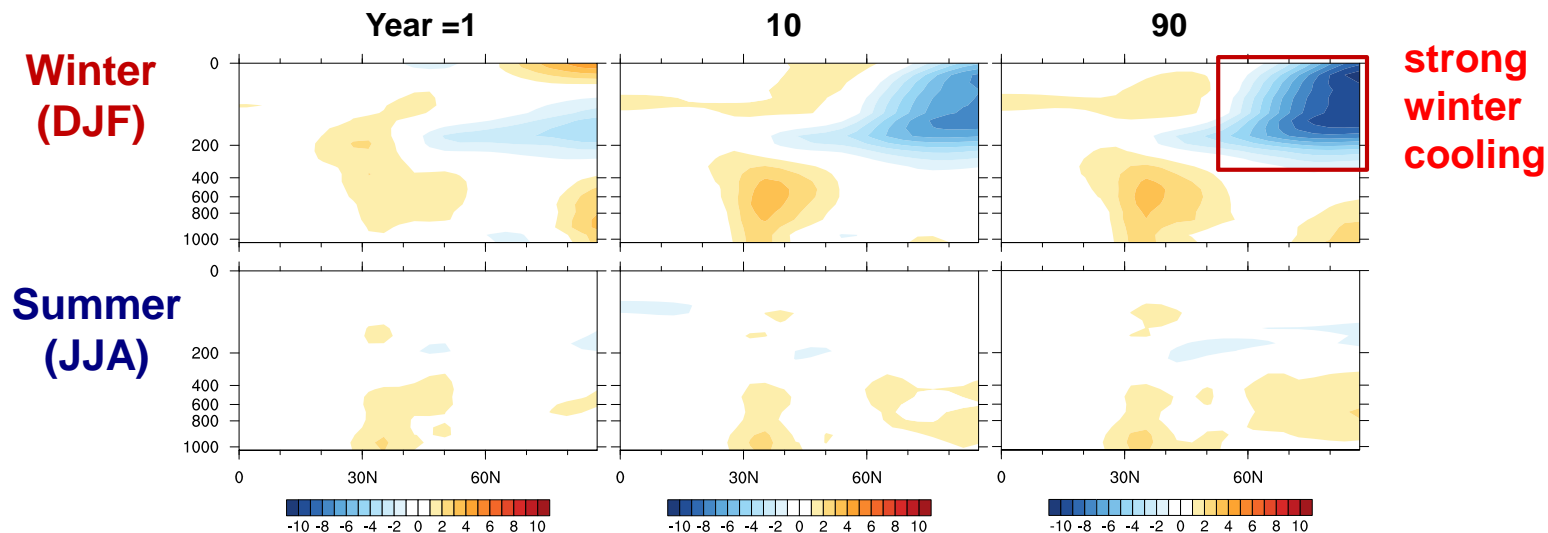
NoTibet → Arctic upper cooling lower warming



Atmospheric process

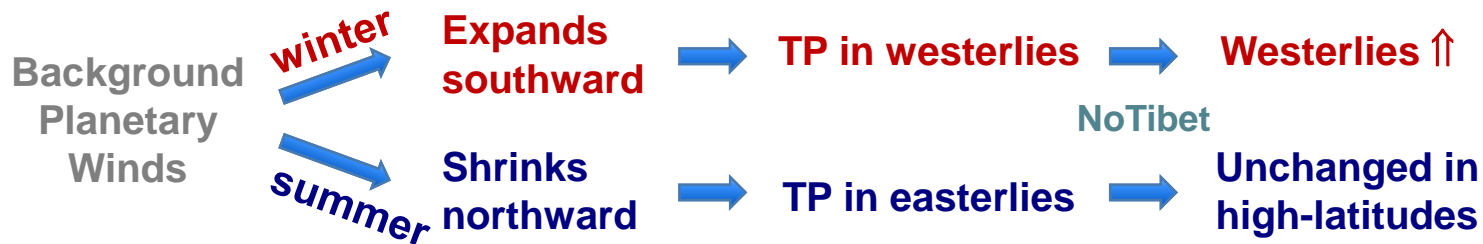
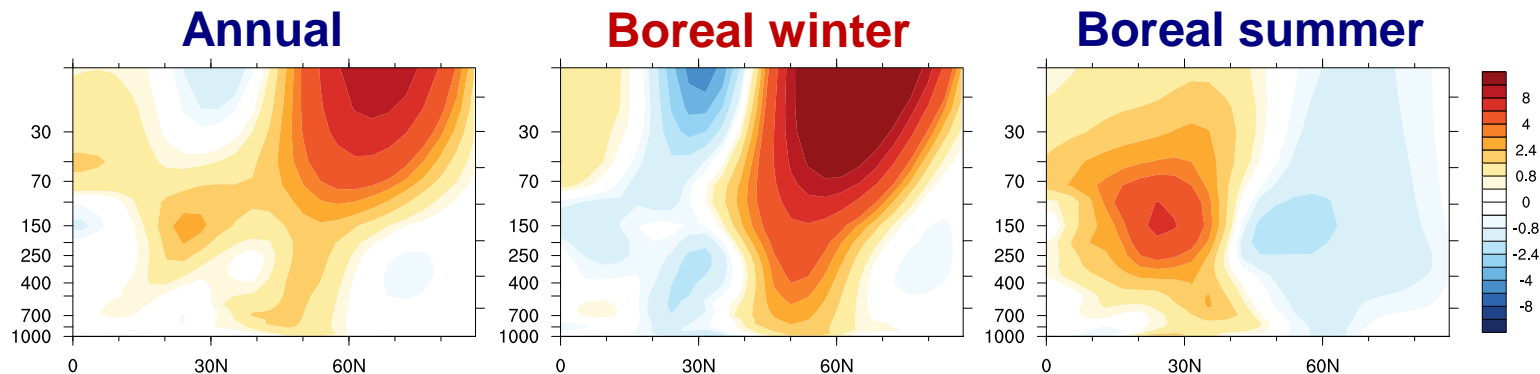
Ocean adjustment

Transient Atmosphere T and Moisture in NH

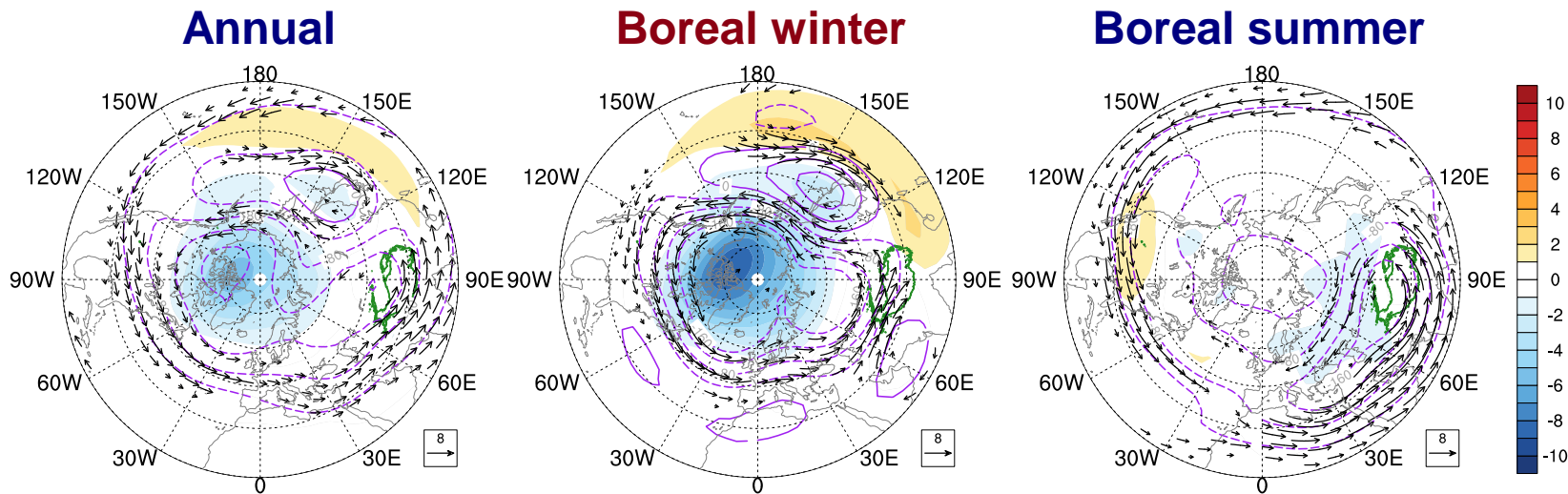


From TP to Arctic: **Seasonality**

Winds Change in NH



Upper Circulation Change in NH



NoTibet → Winter Westerlies ↑ → Polar Low ↑ → Upper-level cooling

Mechanism: *Wave Activity Flux*

NoTibet

Wave dynamics



Higher-latitude Circulation

$$\mathbf{W} = \frac{P \cos \Phi}{2|U|} \begin{bmatrix} \frac{u}{a^2 \cos^2 \Phi} \left[\left(\frac{\partial \Psi'}{\partial \lambda} \right)^2 - \Psi' \frac{\partial^2 \Psi'}{\partial \lambda^2} \right] + \frac{v}{a^2 \cos \Phi} \left[\frac{\partial \Psi'}{\partial \lambda} \frac{\partial \Psi'}{\partial \Phi} - \Psi' \frac{\partial^2 \Psi'}{\partial \lambda \partial \Phi} \right] \\ \frac{u}{a^2 \cos \Phi} \left[\frac{\partial \Psi'}{\partial \lambda} \frac{\partial \Psi'}{\partial \Phi} - \Psi' \frac{\partial^2 \Psi'}{\partial \lambda \partial \Phi} \right] + \frac{v}{a^2} \left[\left(\frac{\partial \Psi'}{\partial \Phi} \right)^2 - \Psi' \frac{\partial^2 \Psi'}{\partial \Phi^2} \right] \\ \frac{f_0^2}{N^2} \left\{ \frac{u}{a \cos \Phi} \left[\frac{\partial \Psi'}{\partial \lambda} \frac{\partial \Psi'}{\partial z} - \Psi' \frac{\partial^2 \Psi'}{\partial \lambda \partial z} \right] + \frac{v}{a} \left[\frac{\partial \Psi'}{\partial \Phi} \frac{\partial \Psi'}{\partial z} - \Psi' \frac{\partial^2 \Psi'}{\partial \Phi \partial z} \right] \right\} \end{bmatrix}$$

Takaya and Nakamura
(1997, 2001)

- Ψ' → perturbations created by the orography (**NoTibet minus Real**)
- **W directions** → directions of Rossby wave energy propagation $\sim C_g$

$$\frac{\partial}{\partial t} M + \nabla \cdot \mathbf{W} = D_T$$

- **Divergence** → wave perturbation energy is emitted and converted to mean kinetic energy

Mechanism: *Wave Activity Flux*

Wave Activity Flux

Takaya et al. (2001)

equal to

$$-[(e - v'v'), u'v', \zeta'P_x']$$

$$e = \frac{1}{2} \left[u'^2 + v'^2 + \left(\frac{f_0}{N^2} \frac{\partial \psi'}{\partial z} \right)^2 \right]$$

Takaya et al. (2015)

equal to 2-D

Applied to 2-D deep system

$$E_u = \left[\frac{1}{2} (v'^2 - u'^2), -u'v' \right] \cos \Phi$$

Hoskins et al. (1983)

$$\frac{\partial \bar{u}}{\partial t} = \cos \Phi \nabla \cdot E_u$$

Simmons et al 1983)

$$\frac{\partial}{\partial t} (\text{PKE}) = - \frac{\partial}{\partial t} (\text{KE}) = - \cos \Phi \cdot \bar{u} \nabla \cdot E_u$$

Passage: *Planetary Wave* in Upper Level

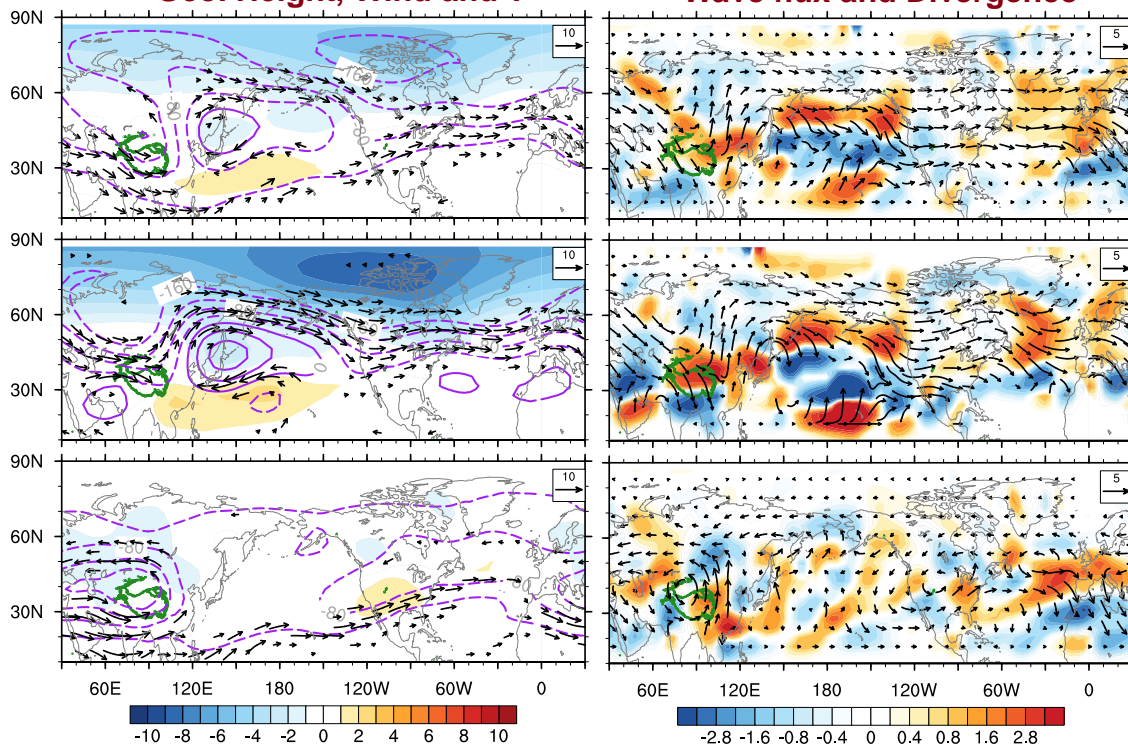
Annual

Boreal Winter

Boreal Summer

Geo. Height, Wind and T

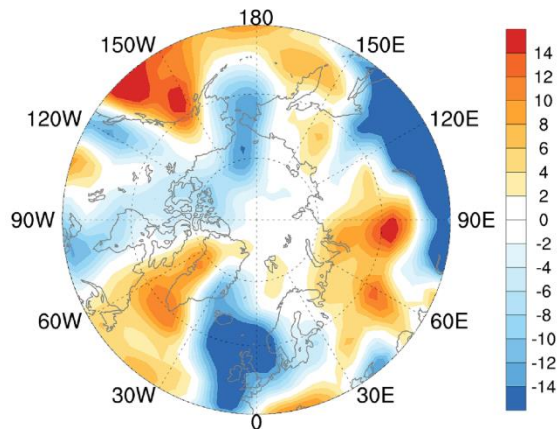
Wave flux and Divergence



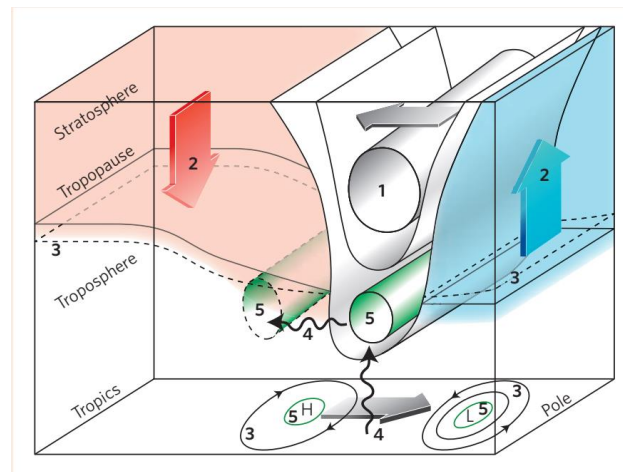
Northeastward
by
planetary flow

Locally trapped

Mechanism to *Polar Cooling* in Upper Level



Winter $\omega < 0$ ascending motion



Kidston et al. (2015) Review article
in Nature Geoscience

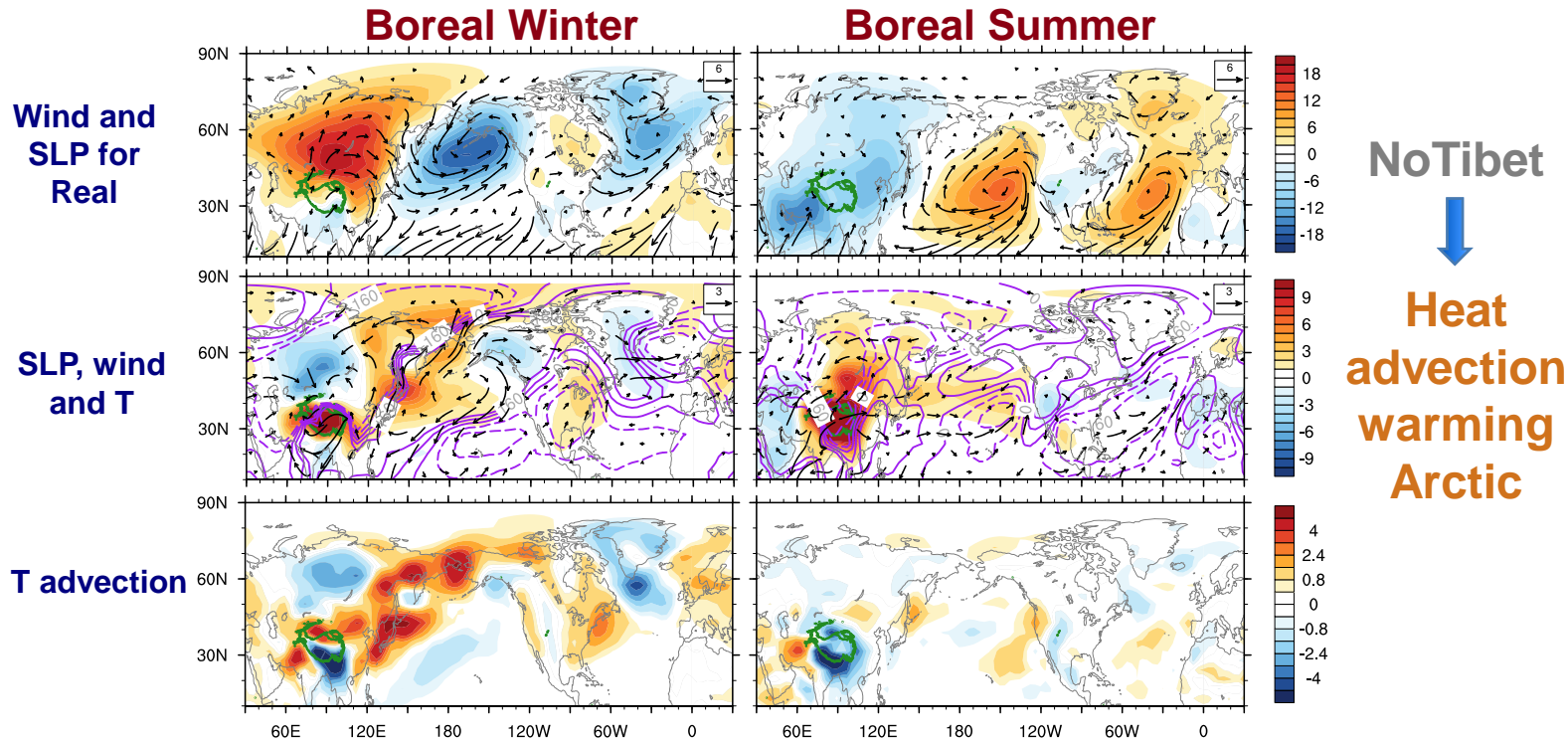
Stratosphere
Westerlies \uparrow

Polar
Ascending

Adiabatic
Cooling

Stratosphere
Cooling

Passage: *Monsoon* in Lower Level



East-northward dominated by monsoon

Seasonality and Passage: *Background* Winds

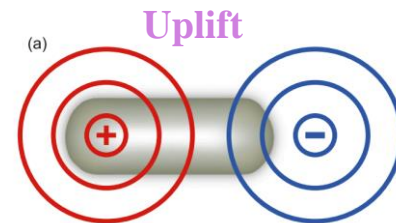
TP-Arctic connection **Robust** in boreal winter:
Seasonal variation of background winds

Upper Westerlies → Eastward wave energy propagation
Easterlies → Locally trapped

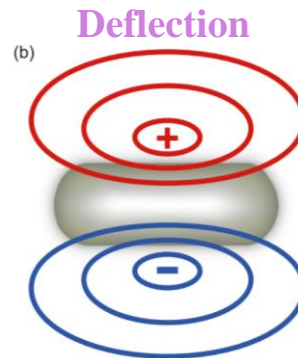
Forced topographic wave solution

Lower
$$\phi_n = \frac{f_0^2 h_n}{gH(K^2 - K_s^2 - irK^2/k\bar{u})}$$

Wave amplitude $\propto U$

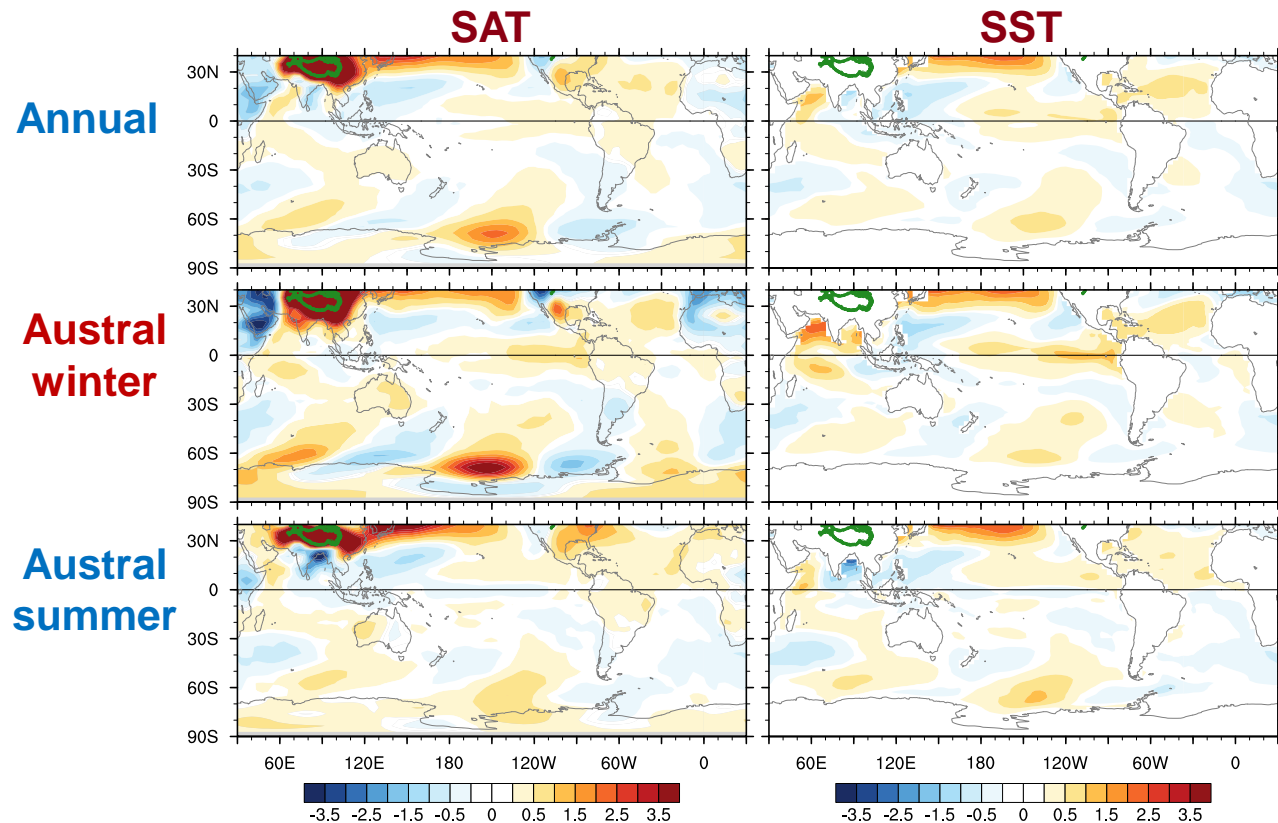


Holton (2004)



Held et al. (1983)

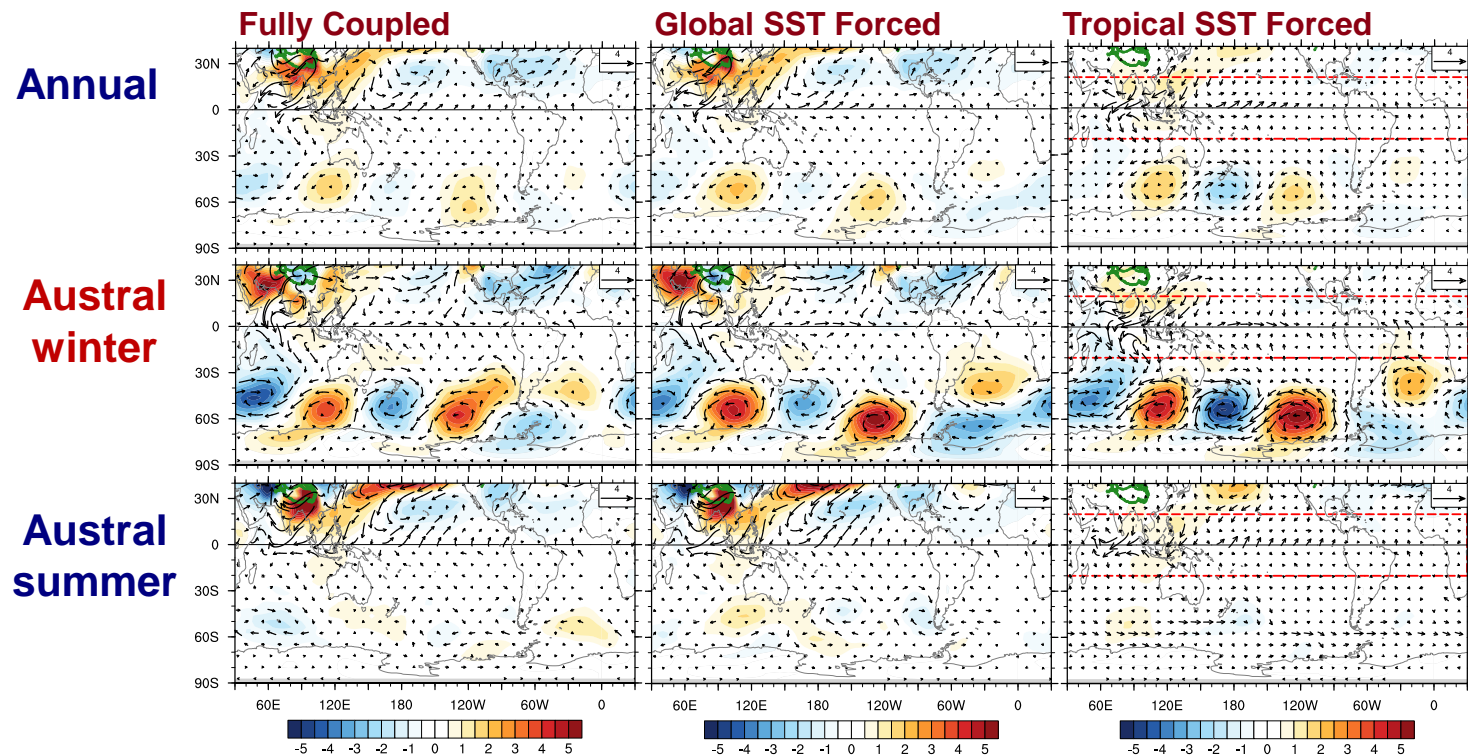
Temperature Change in the SH



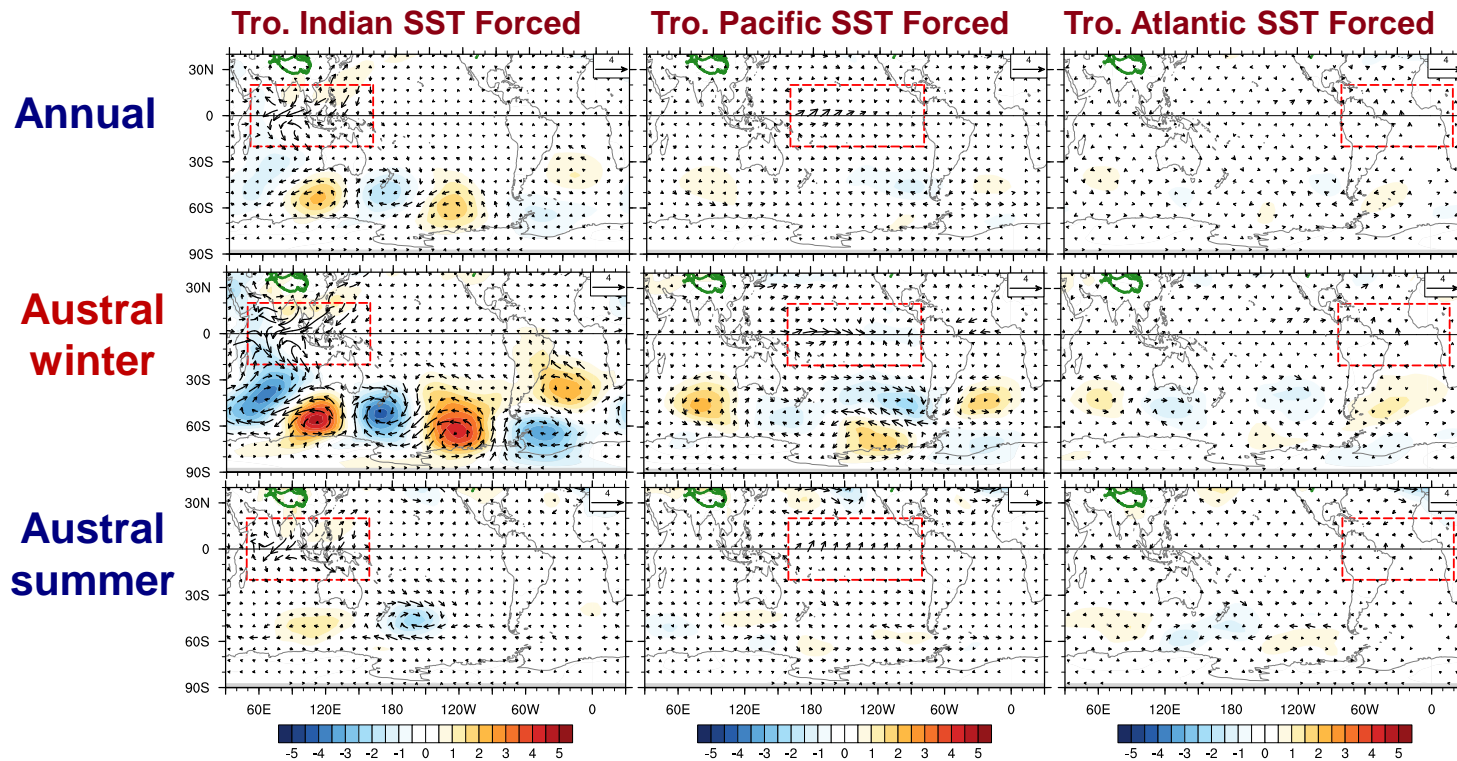
NoTibet
↓
Antarctic
Rossby-wave
like T change

Seasonality and Passage

NoTibet → Wave-5 pattern

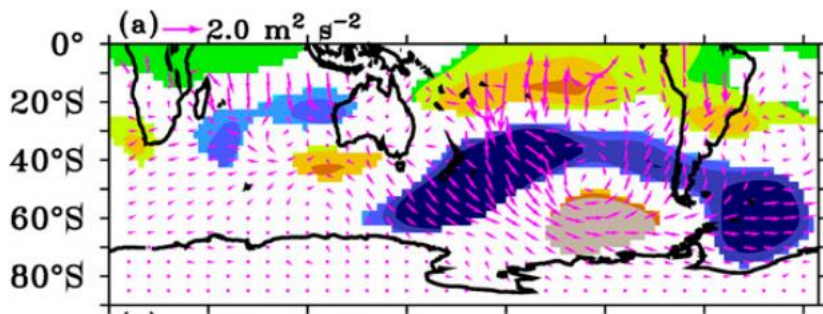
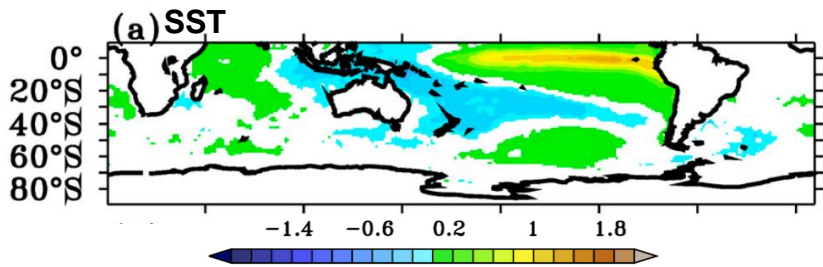


Seasonality and Passage: *Indian Ocean* Ocean

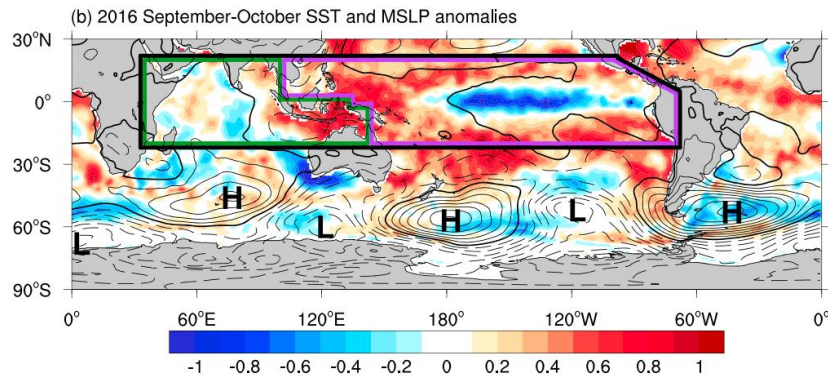


Tropical Indian Ocean act as a bridge !

Indian Ocean Impacts Antarctic: *Stationary Wave*

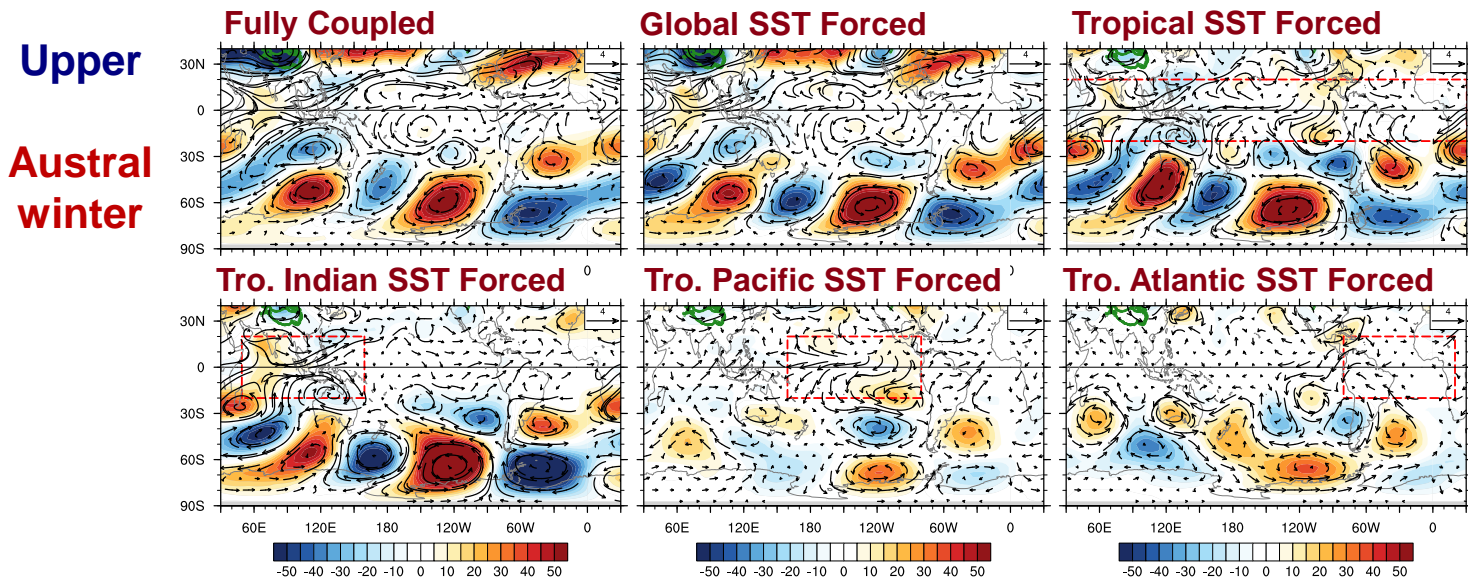


Nuncio and Yuan (2015)



Purich and England (2019)

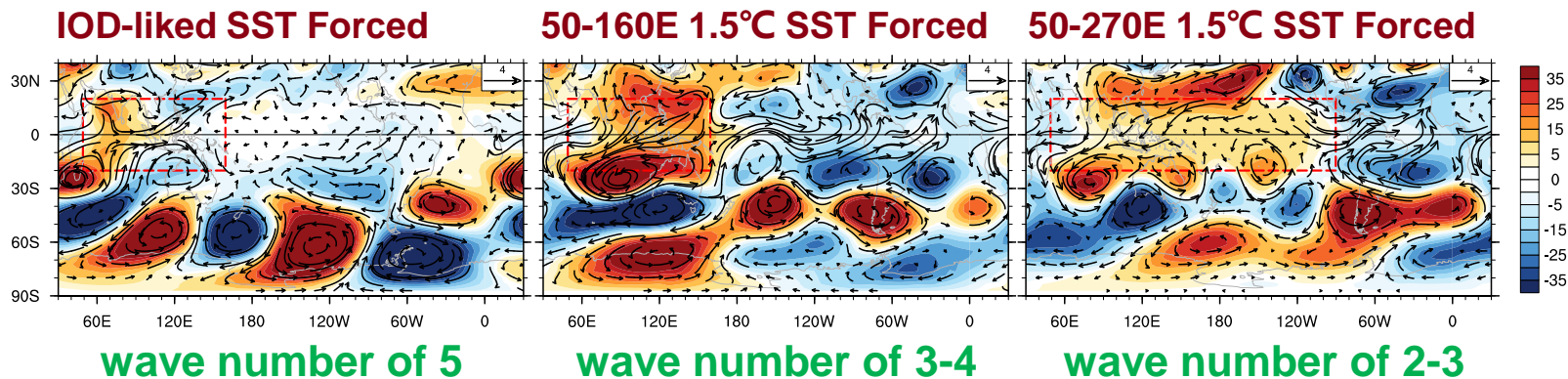
TP → Antarctic: *Stationary* Wave by *Indian* Ocean



Southward and eastward $\frac{C_{gx}}{C_{gy}} \sim \frac{L_y}{L_x}$



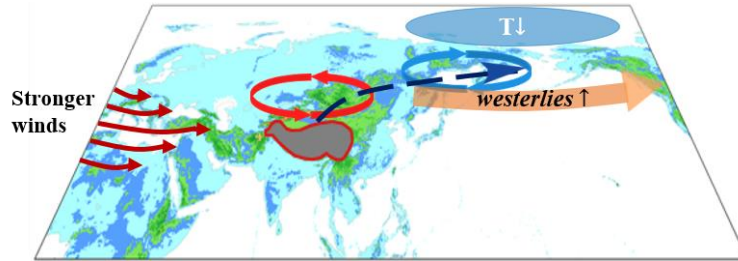
Wave *Number* Related to Forcing Size



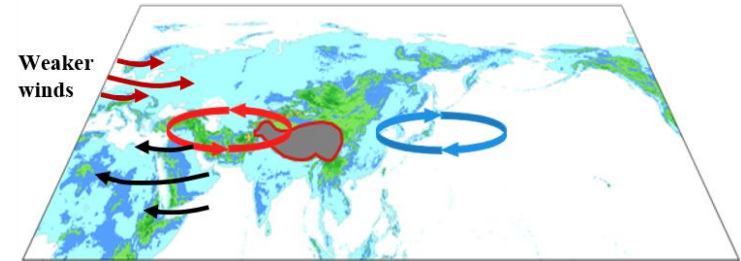
Size and **sign** of the forcing determine the response wave number

From TP to *Arctic*

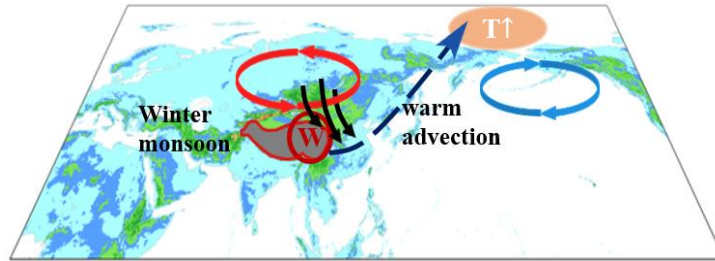
Boreal winter 200hPa



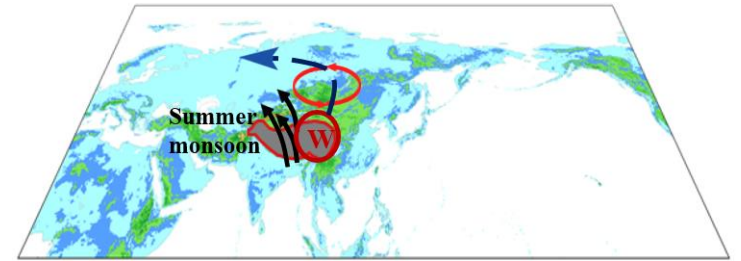
Boreal summer 200hPa



Boreal winter surface

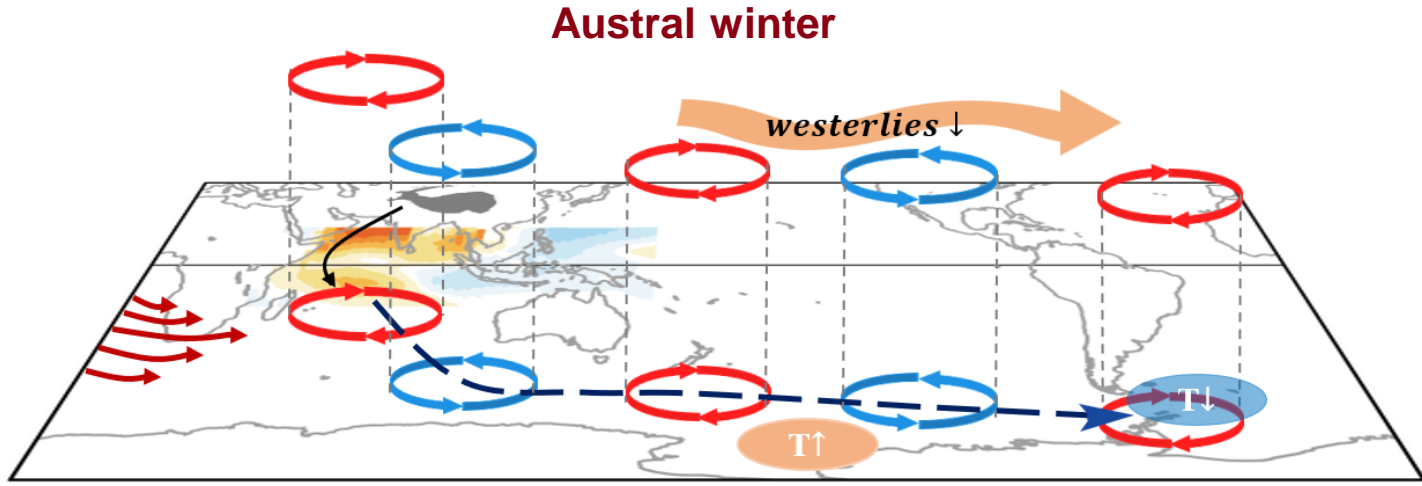


Boreal summer surface



Directly by topography-forced stationary waves

From TP to *Antarctic*



Indirectly by *Indian Ocean* SST-forced stationary waves

Summary

- Connection from TP to Arctic and Antarctic: ***Robust***
- Seasonality and Passage: ***Winter***
- Mechanism: ***Forced Stationary Wave***
- Background: ***Planetary Winds and Monsoon***



LaCOAS
北京大学气候与海-气实验室

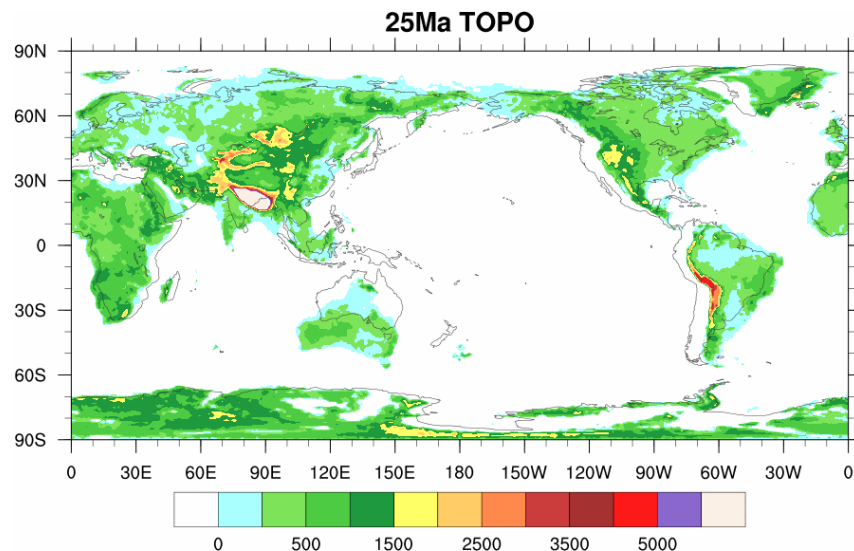
谢谢

TP: *determining* the *Modern GMOC*

Tibetan Plateau



Necessary to Modern AMOC



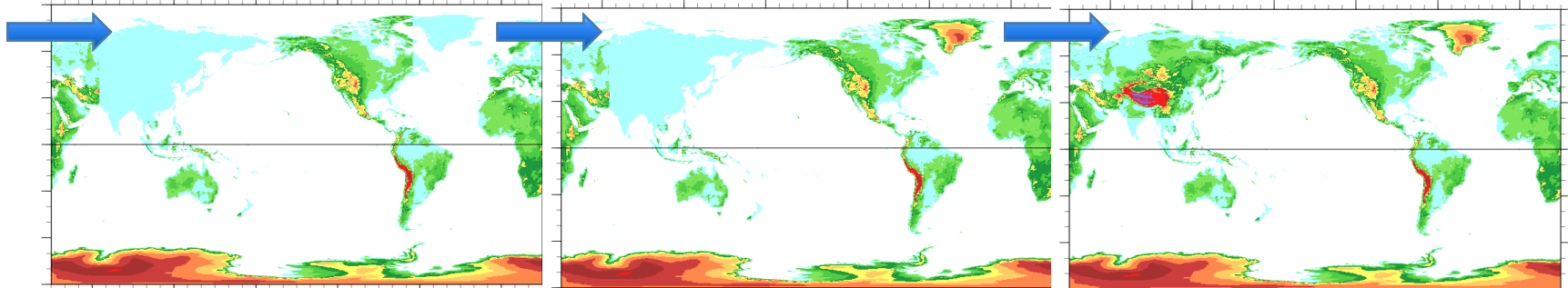
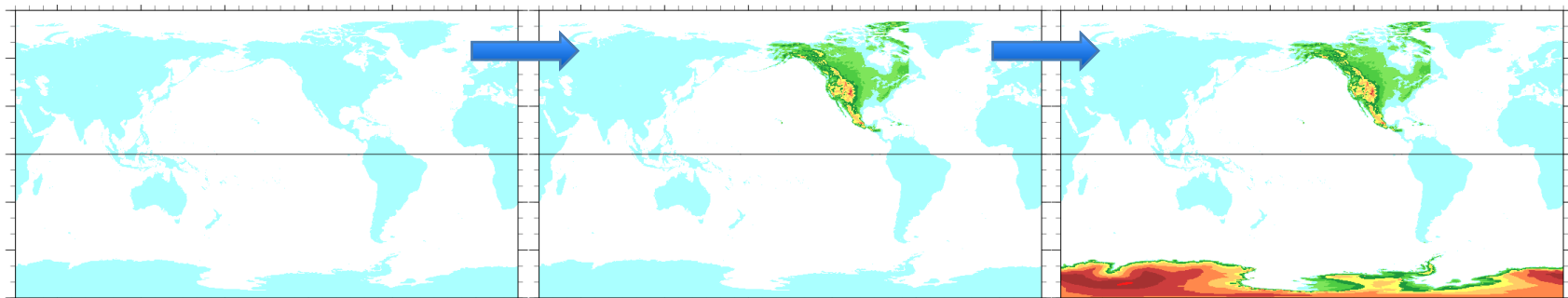
Jiang R. and H. Yang, 2021: Tibetan Plateau: A necessary for the modern AMOC. In *Writing*.

From *Flat* to *Real*

Flat

Rocky 45Ma

... + Antarctic 25Ma

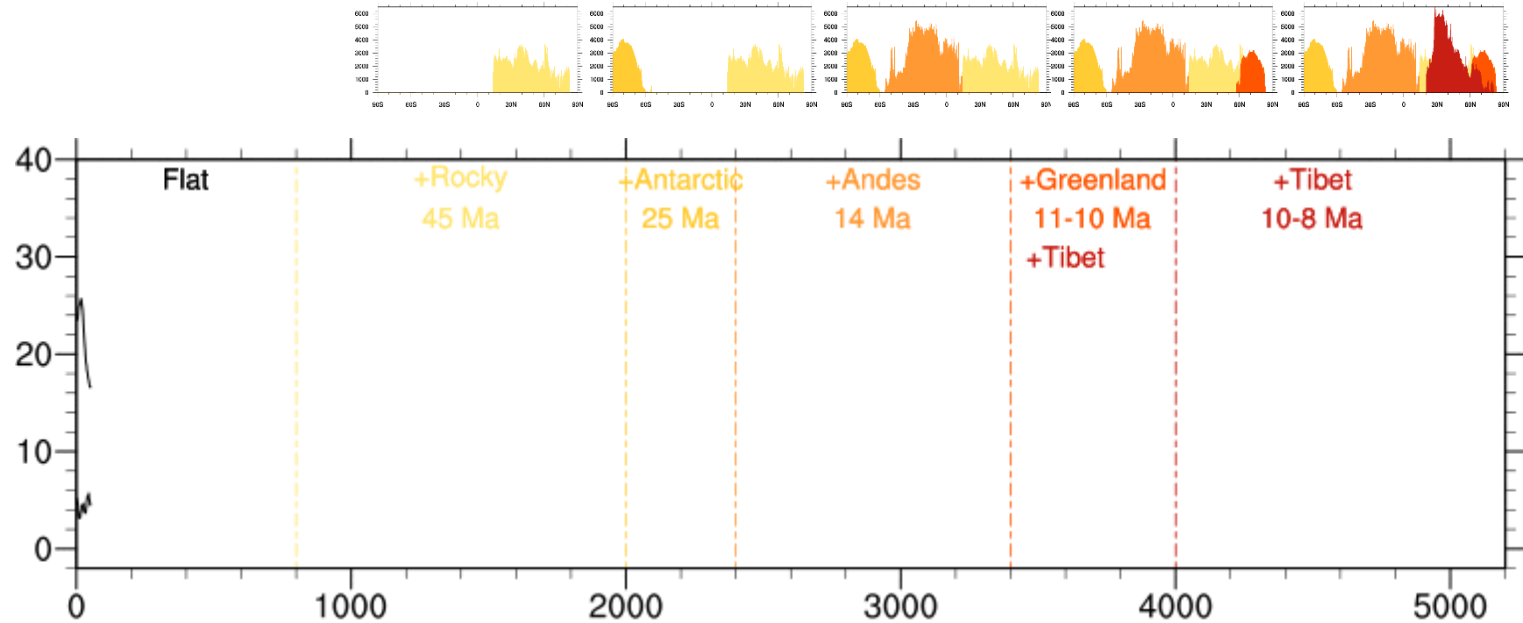


... + Andes 14Ma

... + Greenland 11Ma

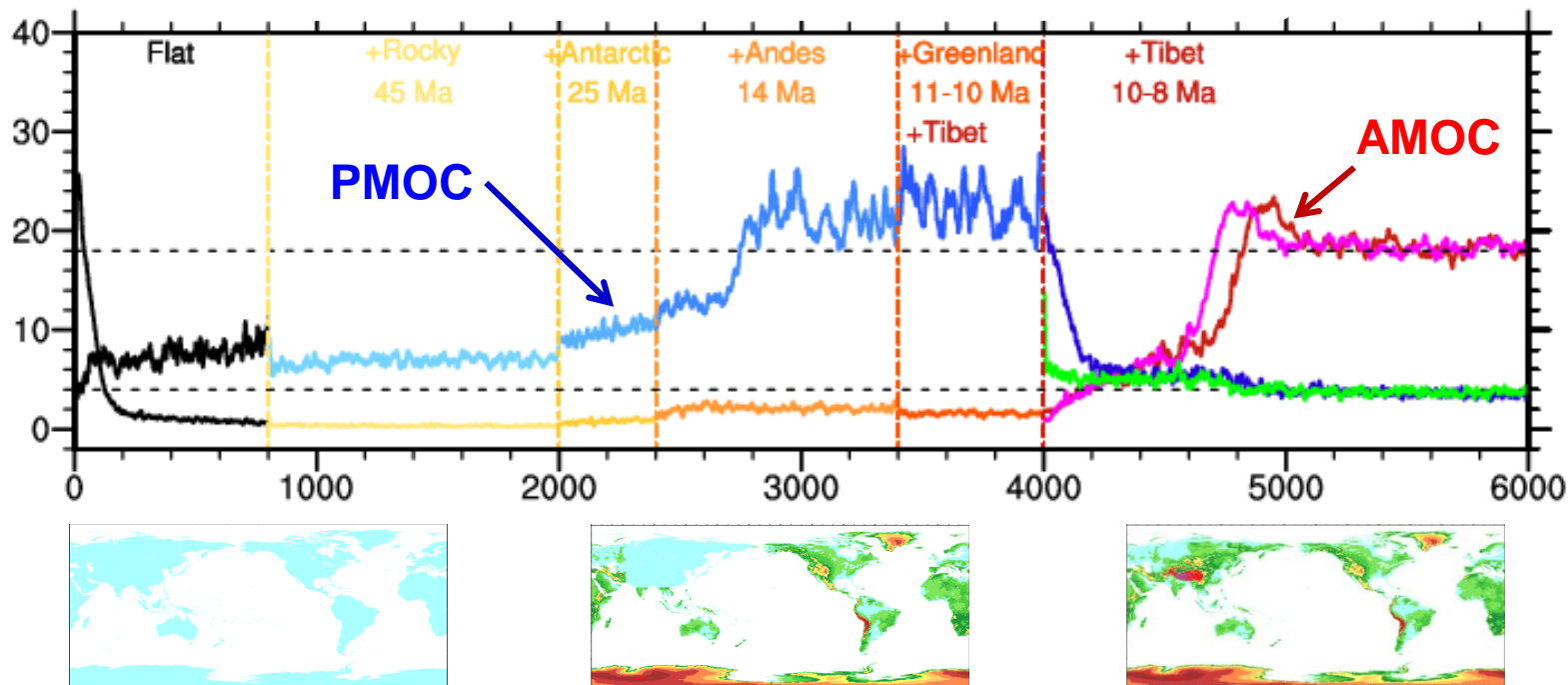
... + Tibet 10Ma = **Real**

From *Flat* to *Real*

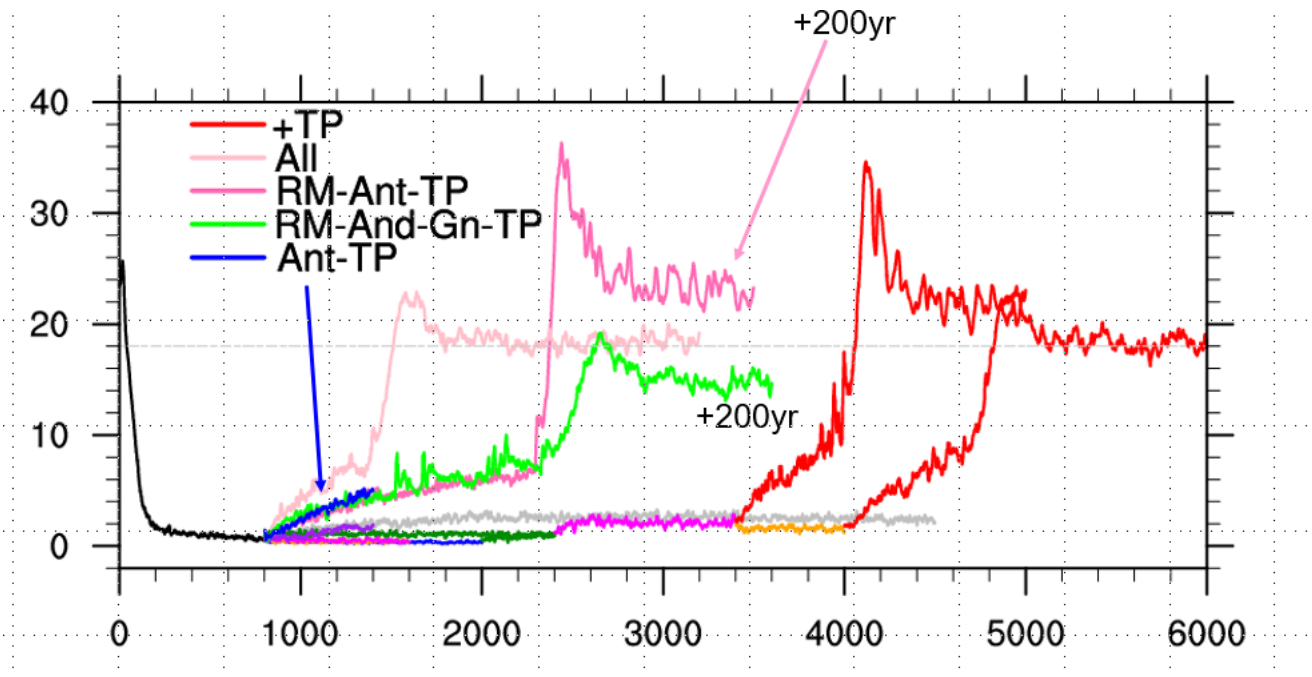


TP: *determining* the *Modern GMOC*

Total 6000 Years Simulations



TP: *determining the Modern GMOC*



Experiments: *Topography Combination*

单个地形

试验时间

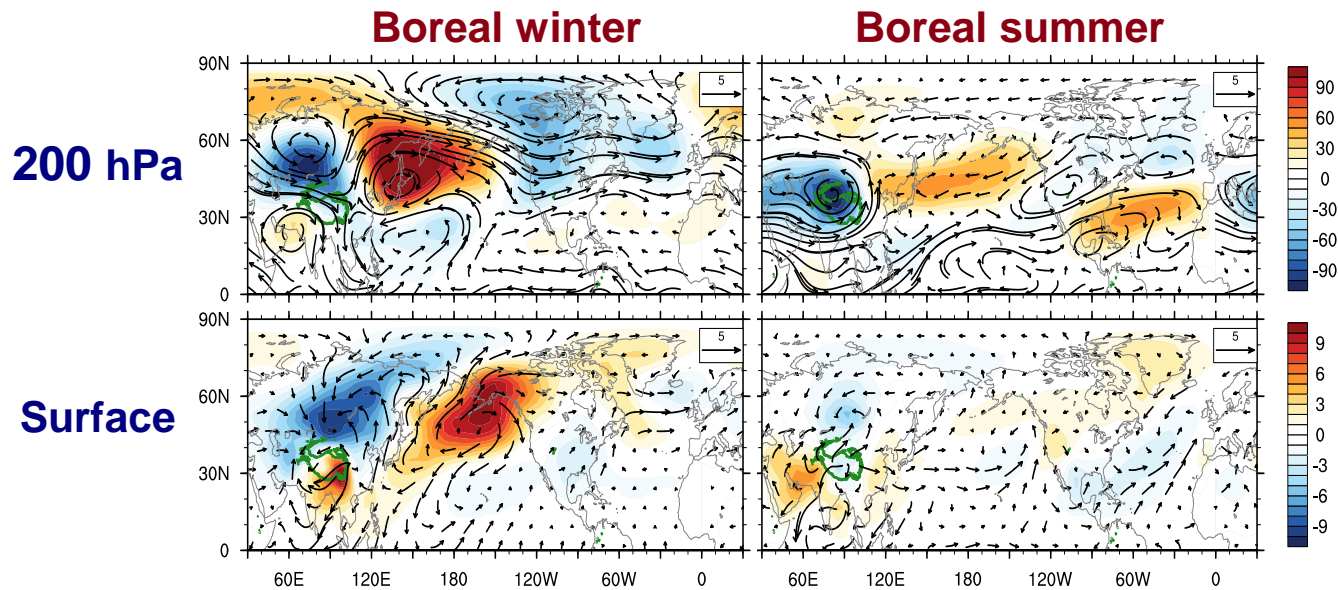
Flat	1-1600	平板
Only _Rocky	801-2000	45Ma
Only _Ant	801-1600	25Ma
Only _Andes	801-1600	14Ma
Only _Greenland	801-1600	11-10Ma
Only _Tibet	801-4700	10-8Ma
Only _All	801-3200	全地形

组合地形试验

试验时间

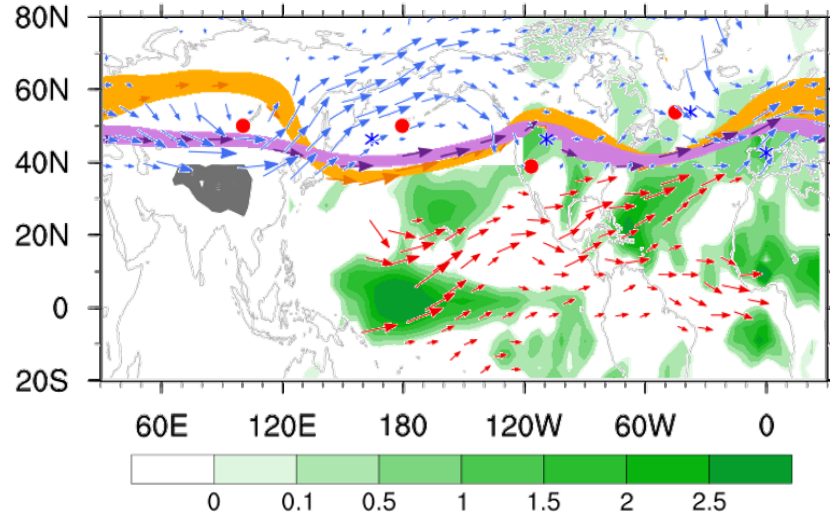
Only _Tibet _Rocky	801-1400
Only _Tibet _Rocky _Ant	801-3400
Rocky-Ant	2001-2400
Rocky-Ant-Andes	2401-3400
Rocky-Ant-Andes-Green	3401-4000
Rocky-Ant-Andes-TP	3401-5000
Rocky-Ant-Andes-Gn-TP	4001-6000

Seasonality: *Stationary Wave* Structure

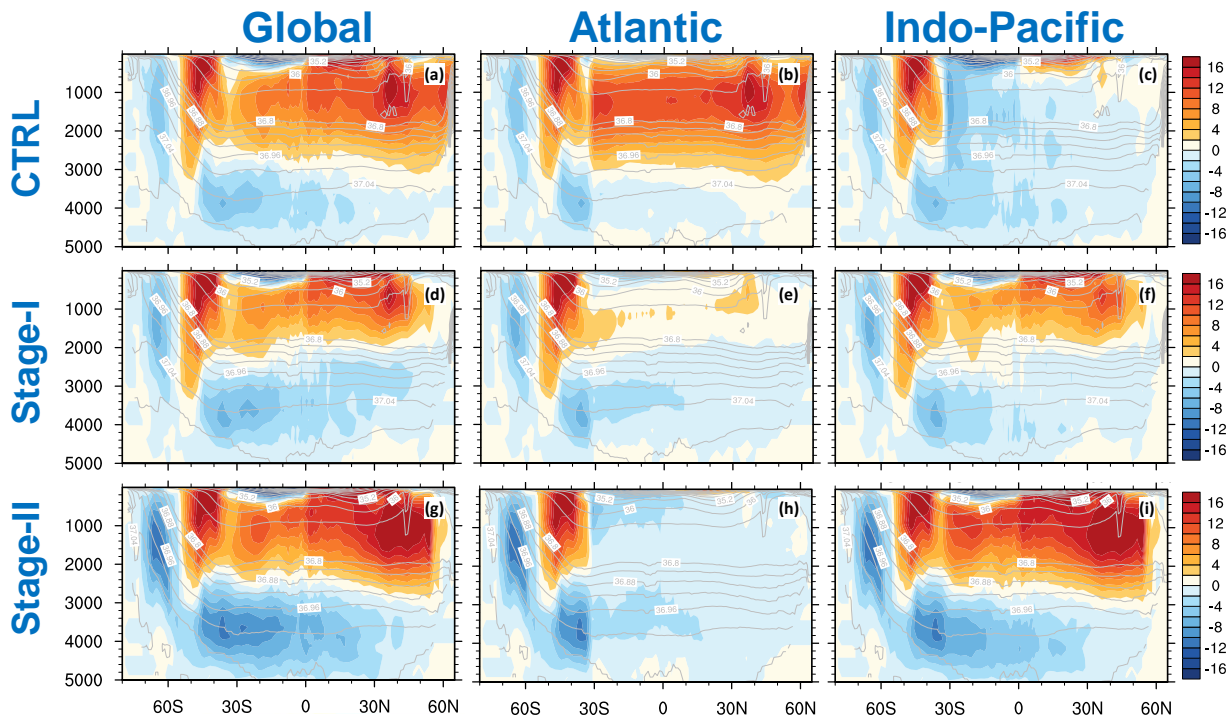


TP located in winter strong westerlies

Stationary Waves with Tibetan Plateau



NoTibet: AABW Enhanced



In Stage-I, the AABW is enhanced in Atlantic basin
In Stage-II, the AABW is also enhanced in Pacific basin

TP effect on AABW: Atmospheric Change

